

2018-01-01

Comparison Of Mesoscopic Fracture Analysis Techniques: A Case Study Near Onion Creek Diapir, Paradox Basin, Utah

Muhammed Tarik Iraz

University of Texas at El Paso, mtiraz@miners.utep.edu

Follow this and additional works at: https://digitalcommons.utep.edu/open_etd



Part of the [Geology Commons](#)

Recommended Citation

Iraz, Muhammed Tarik, "Comparison Of Mesoscopic Fracture Analysis Techniques: A Case Study Near Onion Creek Diapir, Paradox Basin, Utah" (2018). *Open Access Theses & Dissertations*. 94.
https://digitalcommons.utep.edu/open_etd/94

This is brought to you for free and open access by DigitalCommons@UTEP. It has been accepted for inclusion in Open Access Theses & Dissertations by an authorized administrator of DigitalCommons@UTEP. For more information, please contact lweber@utep.edu.

COMPARISON OF MESOSCOPIC FRACTURE ANALYSIS
TECHNIQUES: A CASE STUDY NEAR ONION
CREEK DIAPIR, PARADOX BASIN, UTAH

MUHAMMED TARIK IRAZ

Master's Program in Geological Sciences

APPROVED:

Terry Pavlis, Ph.D., Chair

Katherine Giles, Ph.D.

Mark P. Fischer, Ph.D.

Charles Ambler, Ph.D.
Dean of the Graduate School

Copyright ©

by

Muhammed Tarik Iraz

2018

Dedication

To my family for encouraging me and making me trust myself and my fiancée for being patient for years despite the peer pressures because of our long-distance relationship.

COMPARISON OF MESOSCOPIC FRACTURE ANALYSIS

TECHNIQUES: A CASE STUDY NEAR ONION

CREEK DIAPIR, PARADOX BASIN, UTAH

by

MUHAMMED TARIK IRAZ, B.E.

THESIS

Presented to the Faculty of the Graduate School of

The University of Texas at El Paso

in Partial Fulfillment

of the Requirements

for the Degree of

MASTER OF SCIENCE

Department of Geological Sciences

THE UNIVERSITY OF TEXAS AT EL PASO

December 2018

Acknowledgements

First of all, I would like to thank Turkish Petroleum Corporation for their financial support, and for providing me with this incredible opportunity to learn English which helped me to accomplish my master's degree in the USA.

I would like to thank my committee members, Dr. Terry Pavlis, Dr. Katherine Giles, and especially Dr. Mark P. Fischer for supporting me from Northern Illinois University by emails and video calls, as well as Dr. Rick Allmendinger for the great free software, Stereonet 10, and responding to my question right away.

I would also like to mention my family and fiancée by being patient and encouraging me.

Finally, I am grateful to those who contributed in my academic and personal life in the USA, including I.T.S research group, specifically Nila Matsler; ExxonMobil for supporting me financially during my field work; David Lankford, Kate Grisi, Garrett Goff and Alex Eddy for assisting me in the field work; Galen Kaip for teaching me how to use surveying DGPS (digital Global Positioning System) equipment and both Kinematic and RTK collection methods in the field work; and Dr. Annette Veilleux for answering my questions stoically since I got the acceptance from UTEP.

Abstract

Faults and fractures are abundant along the diapir contact in the Onion Creek Diapir in eastern Utah and relate to the mechanisms of diapir formation. Western Onion Creek exhumes two spectacular exposures of the diapir contact and adjacent wall rocks providing two nearly orthogonal transects oriented ~E-W and S-N. In this study, I used field measurements from natural outcrops and a 3D model developed using Structure-from-Motion/Multiview Stereo photogrammetry to evaluate fracture reservoir development along diapir margins. The 3D-Model was developed using Agisoft Photoscan from ~3200 ground-based images and was georeferenced using 76 ground control points located with a GPS unit capable of cm level accuracy. Simultaneous measurements of ground control points with a handheld WAAS enabled GPS allowed a comparison of model accuracy with this lower resolution GPS. Although horizontal positions of the handheld GPS were in line with known errors in differential GPS, the vertical estimates from the handheld produced a systematic error. The origin of this error is uncertain but likely is related to local terrain affecting the handheld GPS. The 3-D models were interpreted using Maptek I-site Studio, allowing 3D mapping of fractures and bedding as well as measurements of strike and dip from fractures visible on the model. This allowed direct comparison of model based measurements to field measurements at focus sites. This comparison of model results with field data showed comparable results at all sites where fractures formed clear surfaces on the model and arguably the model results were superior at one site where a second fracture set was revealed by allowing measurements at the top of a cliff, above the site. At complexly deformed sites near the diapir contact, however, measurements made on the model were unsatisfactory, producing measurements that were inconsistent with the field data due to

insufficient 3D control and resolution of the model at this site. This problem could have been eliminated had we taken high resolution images of this site as a virtual outcrop, allowing sufficient point density for resolving surfaces on the model. Future studies should consider this approach of developing higher resolution models within larger models where this kind of assessment is needed.

Table of Contents

Dedication	iii
Acknowledgements	v
Abstract	vi
Table of Contents	viii
List of Tables	x
List of Figures	xi
1. Introduction	1
2. Background: Previous Work and Importance of Fracture Network Systems	4
2.1. The Importance of the Area	4
2.2. Fracture Network	4
2.3. Geology of the Paradox basin	5
3. Methodology and Results	10
3.1. Data acquisition	13
3.2. Results	16
3.2.1. DGPS vs WAAS corrected GPS positions on Ground Control Points	16
3.2.2. Geology of the Study Area	19
3.2.3. Analysis of Fracture Orientations	32

3.2.4. Fracture systems	38
4. Discussion	43
4.1. Mapping of fracture systems on virtual outcrop models	43
4.2. Use of 3D outcrop models for remote sensing of fracture orientations	44
4.3. Fracture patterns near the Onion Creek diapir	47
4.4. Use of a Smart Phone for collecting orientation data	50
5. Conclusion	55
Suggestion for Future work	56
References	57
Appendix	65
Appendix 1	65
Appendix 2	75
Appendix 3	81
Appendix 4	84
Vita	87

List of Tables

Table 1: Comparison of GPS data points of survey and handheld (Elevation did not apply)	18
Table 2: Comparison of GPS data points of survey and handheld according Ellipsoidal Height	18
Table 3: Comparison of GPS data points of survey and handheld according Orthometric Height	19

List of Figures

Figure 1: Salt dome trap (Green, 2014)	2
Figure 2: Location of Onion Creek from Google Earth.....	3
Figure 3: A structural map of the Paradox Basin. The map illustrates the salt domes and Uncompahgre Front. (modelled from Trudgill, 2011; after Doelling, 2001)	3
Figure 4: The stratigraphic column of the study area (Modified from Venus et al., 2015)	7
Figure 5: Left-hand figure shows the location of the area and the stations of collected fracture data from Google Earth Pro and right-hand figure shows the geological map of the area (Modified from Doelling 2002, Trudgill, 2011 and Kenroy 2013)	8
Figure 6: Cross section of the eastern part of the Onion Creek Diapir. The location and legend of the cross section is in the Figure 5. A is South and A' is North (from Kenroy (2013) after Doelling (2002)).....	9
Figure 7: A view from the base station of Topcon GB1000 GPS equipment to diapir contact. ..	11
Figure 8: Intermediate and Diapir Sections from Google Earth (the image is oblique view)	15
Figure 9: The stereonet illustrates bedding and fracture's strike and dip raw data from the field. n shows the number of the data point in the stereonets. (Figure 10 shows the location of the stations).....	16
Figure 10: The study area from Google Earth (red dots show the intact fault locations while green show the location of the stations).....	20
Figure 11: The figure shows the contact of the salt diapir with Domain 0. The orange line joints the two fracture stations “On the Contact” to the right and “Closest to the Contact” on the left, a distance of 19m	21

Figure 12: Diapir contact focus site with view zoomed in to the exact area of the field fracture measurements. Left figure shows point cloud with measurements where the right figure shows only patches for orientation estimates. Note the relative sparsity of the point cloud at this zoom level. Figure shows a visualization of remotely sensed fracture orientations developed from 3D model as well as mapped traces of fractures (blue) and faults (red). The stereoplots show the data collected from 3D model (Left) and in situ (right) (see Appendix 1-2 for the data). 23

Figure 13: Closest to the Contact focus site (second fracture station from the diapir, about 20m). Center and lower thirds of the figure are a visualization of ~20m of the model near the fracture focus site outlined by the red circle. Central figure shows the model with interpreted traces of bedding (green), fractures (blue) and faults (red, with fractures only mapped in the focus site (red circle). The lower 1/3 of the figure is the same view as the central figure with the point cloud visualization removed and showing patches where fractures were measured on the model. Upper third of the figure is a zoomed in view of the area of the field measurements with the point cloud include on the left and interpretation only on the right (color scheme and green patches are the same as the lower figure). The stereoplot insets show the fracture data collected from 3D model (left) and in situ (right). All views are looking to the north along an approximately horizontal line. (see Appendix 1-2 for the data) 24

Figure 14: Intermediate-Diapir side focus site (third station from the diapir). The color is the same as figures 12 and 13 with fracture traces (blue) and beds (green). The red circle shows the approximate fracture station that collected data in situ, but Photoscan left as a gap because of deep shadow (see text for discussion)..... 25

Figure 15: Intermediate Cutler Side focus site (the fourth station from the diapir). Upper pair of figures (a and b) show zoomed out view of ~30m of the outcrop with bedding (green) and

fractures (blue) traced on the outcrop (fractures only mapped at the field in-situ site, outlined as the red circle). Upper figure (a) is a visualization of the point cloud with interpretations on it and the lower figure (b) shows the same area with the point cloud turned off. Lower two figures (c and d) show an enlarged view of the area of the field measurements corresponding to the red circle in the upper figures. Color scheme of lower figure is the same as upper figure. Green patches in all figures are sites of fractures measured on the model using multipoint tools in I-site studio. The stereoplot inserts in a show the fracture orientation data collected from 3D model (left) and in situ (right). (see Appendix 1-2 for the data)..... 27

Figure 16: Cutler-Second Farthest focus site (the fifth station from the diapir on the mouth of the Canyon). Upper pair of figures (a and b) show zoomed out view of ~10m of the outcrop with bedding (green) and fractures (blue) traced on the outcrop. Red circle shows the area where field data were acquired and enlarged in the lower figures. Upper figure (a) is a visualization of the point cloud with interpretations on it and the lower figure (b) shows the same area with the point cloud turned off. Lower two figures (c and d) show an enlarged view of the area of the field measurements corresponding to the red circle in the upper figures. Color scheme of lower figure is the same as upper figure. Green patches in all figures are sites of fractures measured on the model using multipoint tools in I-site studio . The stereoplot inserts in a show the fracture orientation data collected from 3D model (left) and in situ (right). (see Appendix 1-2 for the data)..... 29

Figure 17: Cutler-Farthest focus site (the sixth station in the canyon on the cliff from the diapir). Upper Figure shows a visualization of the point cloud with interpretations drawn on it and lower figure shows the interpretation without the point cloud. Red circle shows the position of the in-situ field measurements at this site. Stereograms show orientations measured in situ only. This

site could not be properly georeferenced for direct comparison between field and model data. (see text for discussion) (see Appendix 1 for the data).....	31
Figure 18: The fault filled with Sulphur oxidation (see Figure 8).....	32
Figure 19: The comparison of surface and 3-D model On the Contact focus site (see Appendix 4 for the data).....	35
Figure 20: Diapir contact focus site (the first fracture station from the diapir). The images show the strike and dip measurements of the fractures (green) on the 3D model. When zoomed in, there is a pixel issue on the Maptek. Note that the pixels are larger than original (see Appendix 4 for the data).....	36
Figure 21: Maptek surface (Upper figure shows Figure 25's original figure while lower figure shows the fracture tracing map) (see Appendix 4 for the data)	37
Figure 22: Kamb contoured equal area stereonet plot of poles to fractures with bedding at each site shown as heavy great circle. (The arrows between stereoplots show distance from diapir contact toward the canyon) (see Appendix 1 for the data)	40
Figure 23: Kamb contours after bedding rotation to horizontal (The arrows between stereoplots show distance from diapir contact toward the canyon) (see Appendix 1 for the data).....	41
Figure 24: Offset around the Cutler-Second Farthest site	42
Figure 25: Closest to the Contact focus site. Stereoplot from field data (Left), from the model (Right).....	45
Figure 26: Intermediate-Cutler Side focus site. Stereoplot from field data (Left), from the model (Right).....	46
Figure 27: Cutler-Second Farthest focus site. Stereoplot from field data (Left), from the model (Right).....	46

Figure 28: Polygonal fault zone demonstrates the geometry of a hoop stress field around a salt diapir. The blue dashed circle shows the limit of the hoop stress. The red and blue arrows respectively illustrate the ratio of the maximum and minimum horizontal principal stresses (Carruthers, 2012).	49
Figure 29: The figure shows some examples of radial and concentric faults around salt diapirs. Interpreted by oilcompanies such as Conoco, Ranger Oil, Enterprise Oil, and BP Amoco (Davison, 2000).	50
Figure 30: The comparison of phone data and compass data (the declination of the compass was set up for El Paso which is 8.1° although the phone was set up for Utah that is 10.1°) (Blue circles show the overall fracture orientation while dots show fracture) (see Appendix 3 for the data).	51
Figure 31: Cyan-magenta stereo view of the data station of Cutler side in the intermediate section of the walls from Photoscan stereo view (Y shows north, X shows east)	52
Figure 32: Southern west section of the walls in cyan-magenta Photoscan stereo view (Black arrow shows north)	52
Figure 33: Northern west section of the walls in cyan-magenta Photoscan stereo view (Black arrow shows north)	53
Figure 34: Canyon in cyan-magenta Photoscan stereo view (Black arrow shows north)	53
Figure 35: Southern east side of the walls in cyan-magenta Photoscan stereo view (Y shows north, X shows east).	54
Figure 36: The data stations of the diapir section of the wall in cyan-magenta Photoscan stereo view (Y shows north, X shows east).	54

1. Introduction

The understanding of near salt stratigraphy in salt diapir systems is a key exploration issue for the oil and gas industry (Cook, 2017). Sediments that are affected by salt diapirs can have spectacular structural traps (Figure 1), stratigraphic traps, or both. Unfortunately, deformation at the salt interface is complex and fracture networks in the vicinity of the salt contact can complicate local permeability near the salt margin, confounding industry attempts to model subsurface petroleum systems with potentially expensive results when these models are in error. As a part of this effort to evaluate this type of problem, this paper concentrates on western Onion Creek in eastern Utah where a salt diapir contacts the Permian Cutler Formation with spectacular exposures of the deformed rocks along the contact. The Cutler Formation in the Paradox Basin has been widely studied for more than five decades (eg. Baars, 1966; Baars et al, 1981; Barbeau, 2003; Bromley, 1991; Condon, 1997; Elston et al, 1960, 1962, 1963; Friedman et al, 1994; Goldhammer et al, 1994; Hazel, 1994; Hite et al, 1981; Hudec, 1995; Kluth et al, 2009; Lawton et al, 2006; Moore et al, 2008; Nuccio et al, 1996; Ohlen, 1965; Rasmussen, 2006; Rasmussen et al, 2009; Raup et al 1992; Trudgill et al, 2009; Trudgill, 2011; Wengerd, 1962; White et al, 1983). The Onion Creek diapir (Figure 2) is one of many NW-SE salt walls in Paradox Basin (Figure 3) (Doelling, 1985; Trudgill, 2011). The details of what happens at the salt sediment interface at Onion Creek, however, are poorly understood and it is this problem that is the subject of this study.

Fracture is defined by Narr et al (2006) as discontinuity or parting of a material that is a direct product of brittle failure. In this study, fractures along the contact of the Salt Diapir and Cutler Fm were investigated to see how fracturing changes near the diapir margin and to

determine if any variations are related to salt movement. To analyze a fracture network, it is required to have many fractures which might be grouped based on orientations, types, and ages and there is a need to assess how fracture branches interact with each other (Peacock et al., 2018). The characterization of fracture interactions is listed as Fracture type (opening, closing, shearing or mixed-mode displacements), geometry and topology, chronology, kinematics (Peacock, 2018).

In this study I examine a well-exposed contact zone between the Onion Creek diapir and adjacent Cutler Formation. The focus of the study was on use of new technology, primarily 3D Structure-from Motion/Multiview Stereo modeling of the outcrops but also rapid field measurement of orientations using an Iphone, to assess the fracture systems within the outcrop.

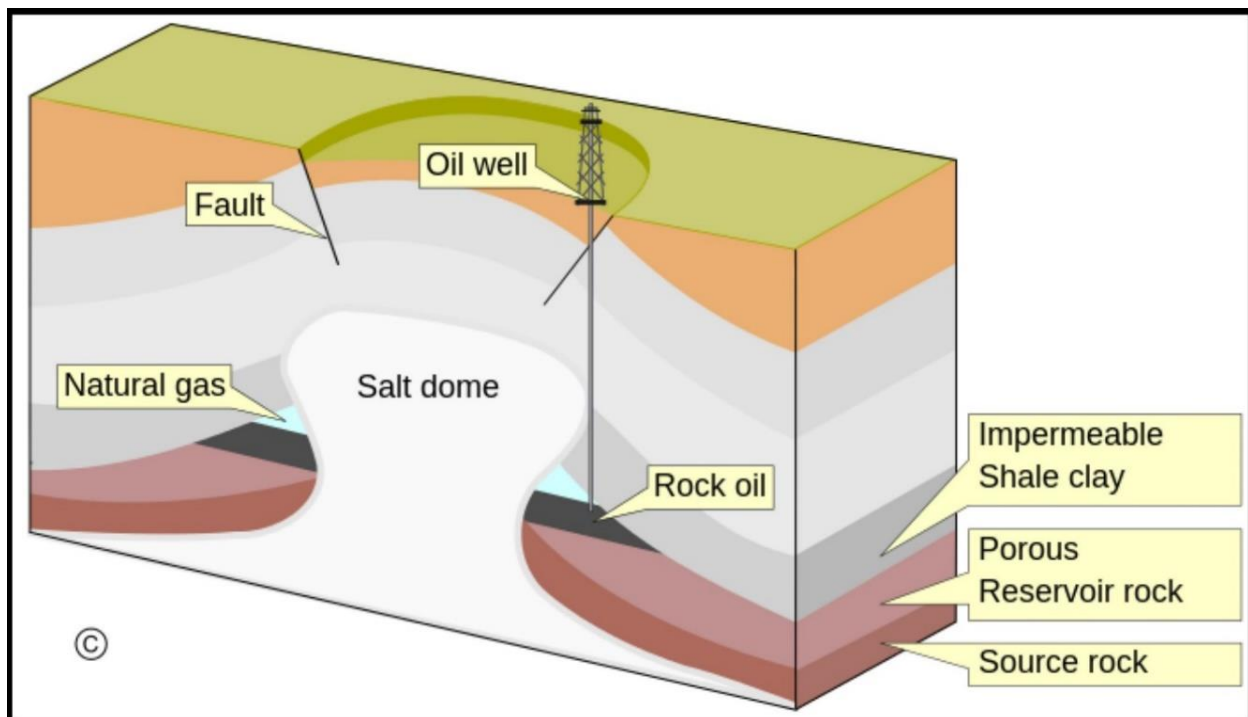


Figure 1: Salt dome trap (Green, 2014)

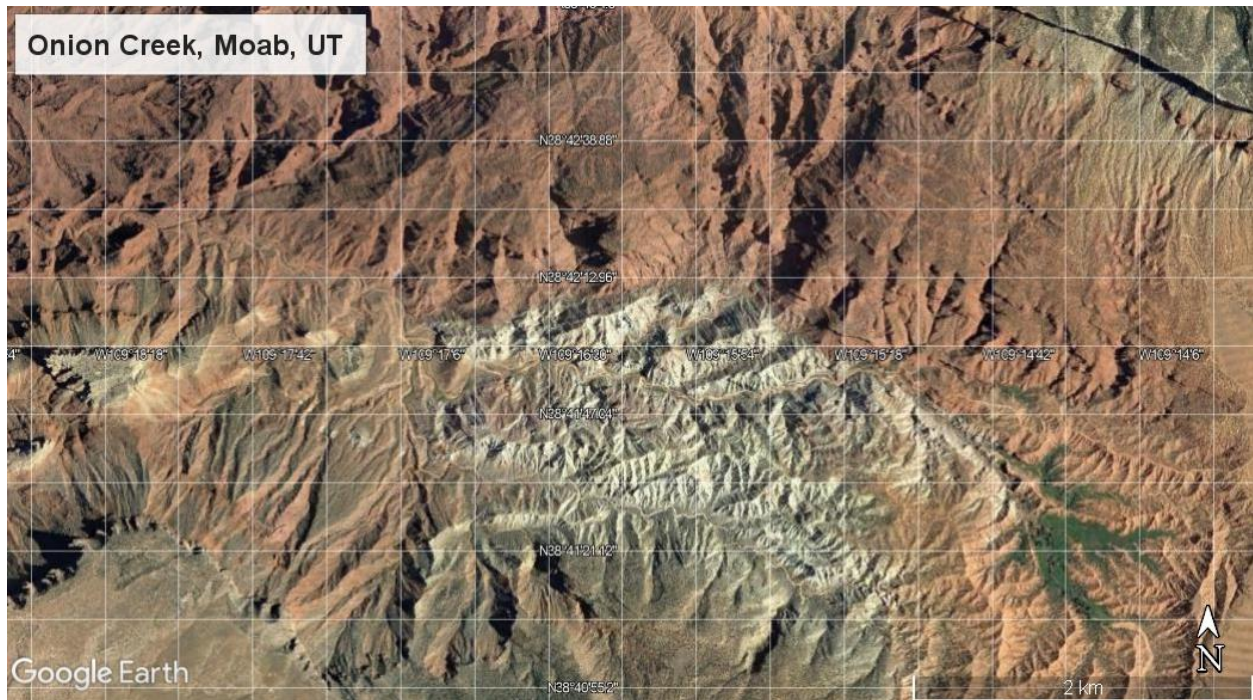


Figure 2: Location of Onion Creek from Google Earth

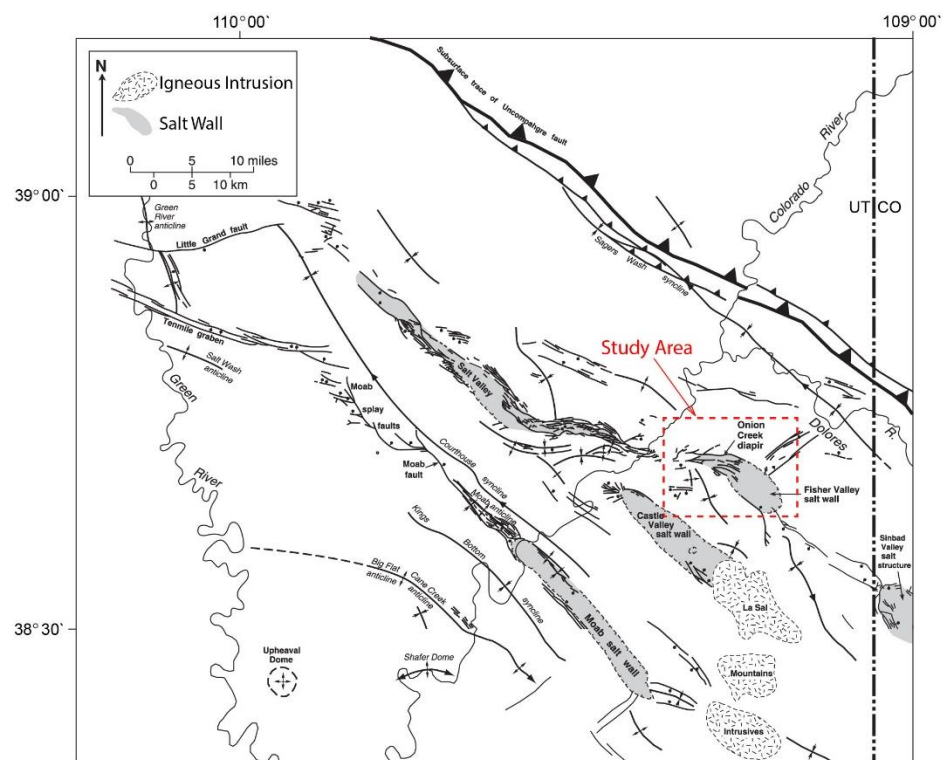


Figure 3: A structural map of the Paradox Basin. The map illustrates the salt domes and Uncompahgre Front. (modelled from Trudgill, 2011; after Doelling, 2001)

2. Background: Previous Work and Importance of Fracture Network Systems

2.1. The Importance of the Area

In the Paradox Basin, petroleum has been extracted from stratigraphic traps in rocks of Devonian and Carboniferous (Mississippian and Pennsylvanian) ages (Baars and Stevenson, 1982). A massive sequence (~up to 4300m (Trudgill, 2011)) of Pennsylvanian rocks with organic-rich black shales form petroleum source rocks of the basin with younger Pennsylvanian-Permian clastic strata forming the primary reservoir rocks. The black shales generated over 400 MMbbl (million barrels of oil) and 1 TCF gas (trillion cubic feet) within the basin (Baars and Stevenson, 1982; Hite et al., 1984; Nuccio and Condon, 1996; Grisi, 2018). Thus, understanding this petroleum system is important for the oil and gas industry (Cook, 2017).

2.2. Fracture Network

The significance of fractures is understood in engineering, geotechnical, and hydrogeological applications because they can play a role as hydraulic conductors, providing pathways for fluid flow or a barrier which can inhibit flow toward them. Many reservoirs of petroleum, gas, geothermal, and water are developed in fractured rocks (NRC, 1996). Fractures form in a variety of modes and can record a stress history within rock bodies, but separating fracture generation and fracture type within individual outcrops can be challenging in complexly deformed rocks. Here I attempt to evaluate this type of problem using a new approach in which limited field time is supplemented by extensive study of the outcrop post-field work, using a 3D model of the outcrop and remotely sensed fracture measurements.

The way to identify, locate and characterize the fractures that are important hydraulic conductors or barriers is through determination of their orientation and connectivity. Rock properties have a significant function in both the creation of fractures and in the structure of fault zones. Within the inner structures of the fractures as well as the fracture zones in various lithologic sections, there exist distinct differences (NRC, 1996).

There are numerous ways to determine fracture properties such as surface methods, borehole-borehole/borehole-surface methods, single-hole methods, or fluid flow monitoring using geophysical methods. Subsurface methods, however, are always a poor estimator of fracture systems because of limited sampling. Hence, excellent natural exposures provide critical information to assess subsurface data when only those data are available. In this study neither well data nor seismic data are available, but excellent surface exposures allow careful assessment of fracture systems in situ. Specifically, after field reconnaissance we recognized that a river cut provided two nearly orthogonal transects oriented ~E-W and S-N that produced exceptional exposures of fractured rocks along the Onion Creek diapir contact. Therefore, I used these natural outcrops to evaluate fracture reservoir development along diapir margins.

2.3. Geology of the Paradox basin

The Paradox Basin is a large asymmetric basin. Onion Creek exposes one of the salt diapirs in the basin. The Paradox Basin developed from middle of Pennsylvanian to Permian time as part of the Ancestral Rocky Mountain (ARM) orogenic event (Kelley, 1958; Doelling, 1988; Barbeau, 2003). The basin exhibits similarities to intraforeland flexural basins developed by basement-cored thrusting which formed the Uncompahgre Uplift. The basin is elongate in a NW-SE direction and it is bounded to the northeast by the Uncompahgre Thrust. Evaporites, black shale, and halite of the Pennsylvanian Paradox Formation were mobilized by sediment

loading and displaced upward through salt diapirism (Trudgill, 2011). The Onion Creek diapir in particular is located within 16 km of the Uncompahgre Thrust (Table 1 from Trudgill, 2011, modified from Banbury, 2005) (Figure 3). The diapir exposes evaporites of Paradox Fm and large carbonate blocks of Honaker Trail Fm which were distorted whilst Undifferentiated Permian Cutler Fm was deposited across the basin as progradational alluvial to fluvial fans (Cook, 2017). Diapirism continued through deposition of younger rocks as well, at least through Late Tertiary and Quaternary time during deposition of Upper Triassic of the Chinle Formation, Jurassic Morrison Formation, Tertiary burial, and Late Tertiary-Quaternary, erosion and dissolution of salt (Doelling, 1988; Hazel, 1994; Doelling, 2002; Trudgill, 2011).

Deposition of the Cutler Group had a well know effect on formation of the salt domes and has been studied for some time (Barbeau, 2003; Condon, 1997; Elston et al, 1962; Hazel, 1994; Nuccio et al, 1996; Trudgill et al, 2009; Trudgill, 2011 etc.). Before salt movement, Honaker Trail Fm (Pennsylvanian) was deposited and therefore is the mostly strongly deformed rock unit above the Paradox Formation (Trudgill and Paz, 2009; Trudgill, 2011).

The Onion Creek salt diapir is one of a linear trend of salt structures. Specifically, it creates a structural high within the Fisher Valley salt wall (Figure 3), and has the best exposures of the evaporites in the Paradox Basin. At the surface more than 4 km² of evaporites are revealed in the core of the diapir and with outcrops of the Paradox Formation composed primarily of gypsum, with minor anhydrite and clastic rocks (Trudgill, 2011).

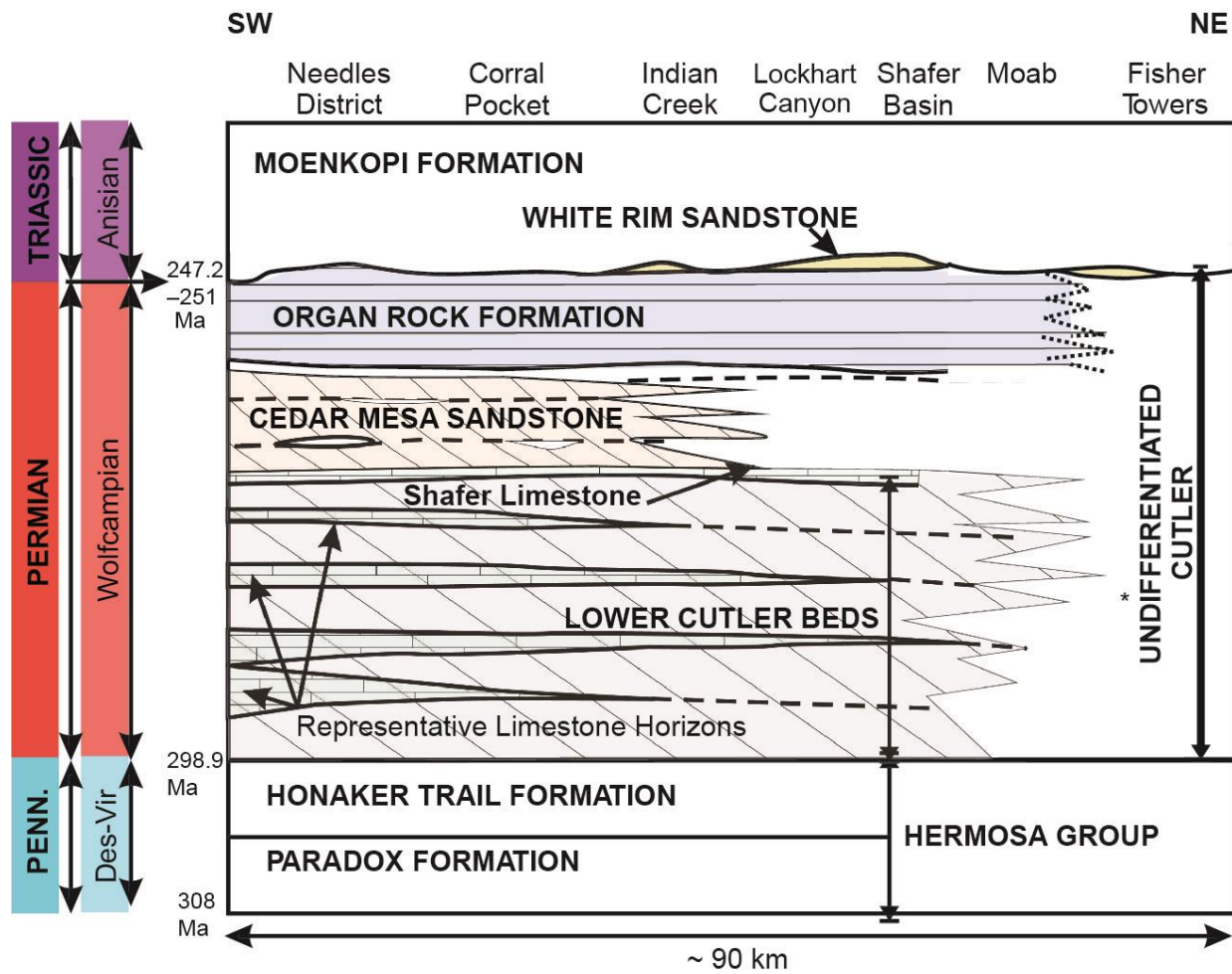


Figure 4: The stratigraphic column of the study area (Modified from Venus et al., 2015)

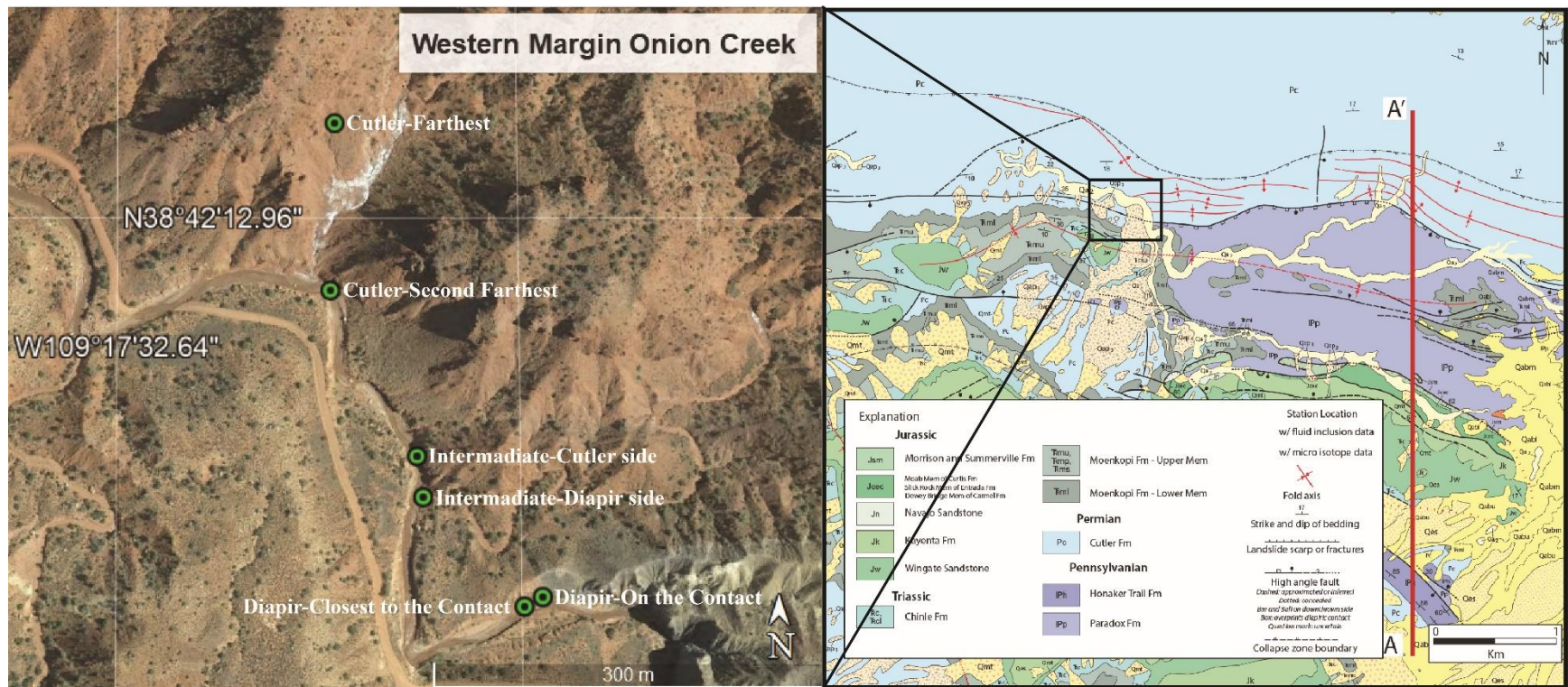


Figure 5: Left-hand figure shows the location of the area and the stations of collected fracture data from Google Earth Pro and right-hand figure shows the geological map of the area (Modified from Doelling 2002, Trudgill, 2011 and Kenroy 2013)

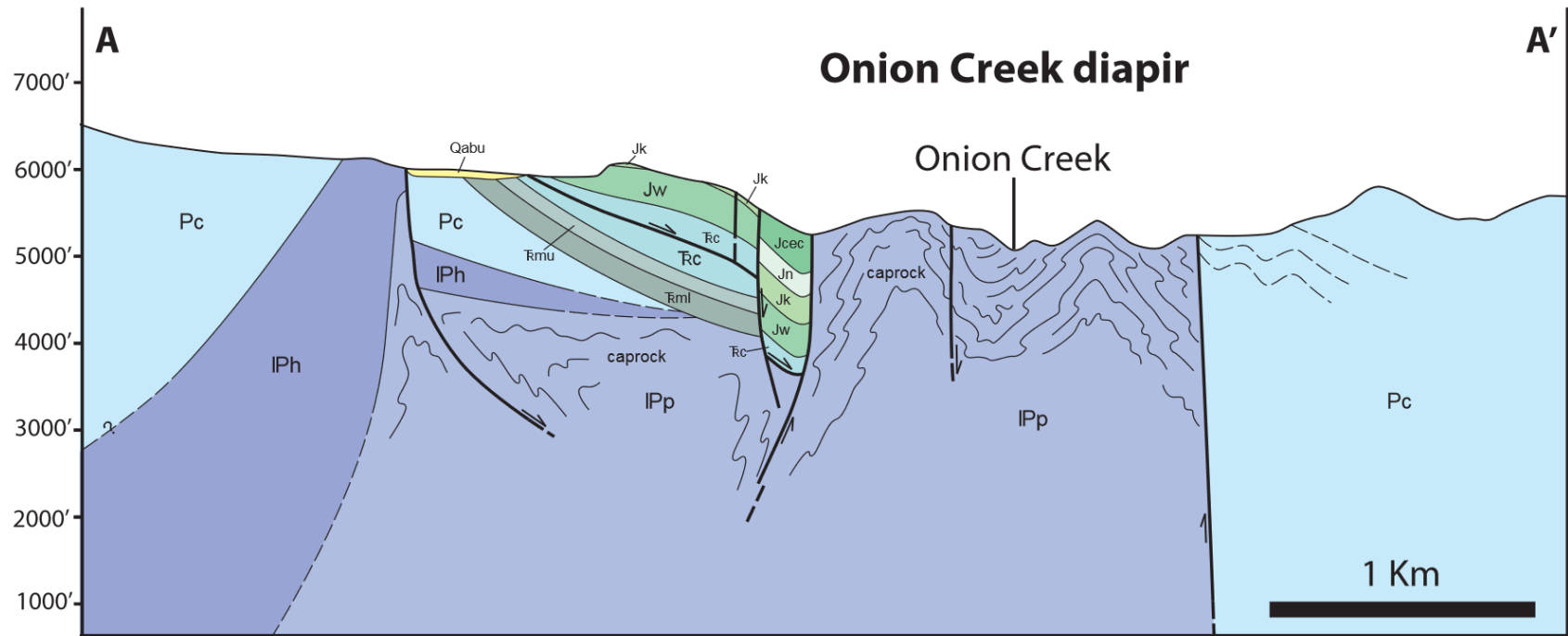


Figure 6: Cross section of the eastern part of the Onion Creek Diapir. The location and legend of the cross section is in the Figure 5. A is South and A' is North (from Kenroy (2013) after Doelling (2002)).

3. Methodology and Results

For the field work I emphasized two complementary approaches: 1) a study of fracture systems in situ such as measuring the orientation of fractures along some segments of focus study outcrops while at the same time acquiring GPS data and photos for photogrammetric modeling, 2) obtaining data for a 3D model using photogrammetry and using this model as a base for macroscopic fracture analysis. For the in-situ fracture studies I measured fracture systems, focusing on details in 6 stations, chosen from sites ranging from at the diapir contact across complexly deformed rocks of the Cutler Fm into more intact Cutler Formation approximately 500m from the diapir. Whilst making measurements, I used three different sized circular scan lines that were located in the same center. This allowed me to spatially reference fracture measurements to estimate other fracture parameters instead of measuring every single fracture. In this method the number of fractures intersecting the circumference (n) of the circle and the number of fracture traces terminating in the circle interior (m) give a quantitative estimate of fracture parameters such as density, size and intensity, (Reiss, 1982; Nelson, 1985; Dershowitz and LaPointe, 1994; Narr, 1996; Rohrbaugh Jr, 2002). These parameters are important for deduction of patterns that would assist in the evaluation of hydrocarbon flow and storage in fractured reservoirs. Through this I would know whether the fractures cross the circle lines which would give the estimated lengths. Unfortunately, I did not count n nor m in the field under an assumption that I could use the 3D model to estimate these parameters, thus, increasing field time for other studies. Ultimately, technology limitations prevented clear application of this method.

Because of the need to collect many strike and dip measurements in a limited time, I chose to use a digital compass. To determine if my device (iPhone 7 Plus) was comparable with an analog compass I collected comparison data sets (Figure 30).

For the macroscopic fracture studies, I began by developing 3D models of the outcrop using photogrammetry methods of Structure from Motion/Multiview Stereo (SfM/MvS) methods (Carrivick et al, 2016). For these studies I began by placing markers throughout the study area and used a high-resolution DGPS Survey (Digital Global Positioning System) (Figure 7) to accurately locate the markers –which has an estimated accuracy of 1 cm or better. At the same time I used a handheld GPS to provide approximate locations for later reference, which provided an opportunity to compare the accuracy of the handheld unit relative to the DGPS. To get a ground control data from DGPS, I needed to make tripod stable and push the bottom. In the beginning, I set it for 30 sec to let communicate with satellite for each data. While I was waiting



Figure 7: A view from the base station of Topcon GB1000 GPS equipment to diapir contact.

for DGPS. I let the handheld GPS to run 15-20 sec, so I could write down latitude, longitude and elevation. The handheld GPS was always running next to the tripod, it could expect to have maximum 1 m difference on X-Y axes of DGPS because of the cable of DGPS with its antenna. However, I would not expect to see meters difference on the elevation axis.

I then used the department's camera to acquire ground-based images of the exposure, I scanned across the exposure from as many camera positions as possible. For example, images were acquired at distances of a maximum of 4m. These data formed the basis for 3D model development, using the SfM/MvS software Photoscan from Agisoft Ltd.

In the office I used the photos acquired in the field and the DGPS data to process the imagery using Agisoft Photoscan, in a computer Intel Xeon 12 CPU E5, 1,.60 GHz, and 64 GB RAM. Processing times for this effort was at least a week for each wall of the focus outcrops in medium resolution cloud density. I used the workflow of Matthews et al (2016) to insure high model accuracy, and through use of the high-resolution DGPS positioning of the markers, the resultant model should be accurate to the mm level. Later, once the model was developed, I used a combination of Maptek i-site Studio and CloudCompare to interpret the fracture systems. This included superimposing the field fracture data onto the model as well as digitizing features on the model, including fractures visible on the model. For fractures visible on the model, I use multipoint methods for estimating fracture orientations, together with fracture mapping to analyze the fracture systems.

To interpret the fracture data collected in the field I used the traditional method of orientation analysis on a stereonet. After exporting my phone data to the computer, I used Stereonet 10.0 and Adobe Illustrator CC 2018 to compare the six fracture stations to each other. All figures refer to the sequence of stations as you walk in the river cut from the diapir contact to

northern side (Figure 5): *On the Contact/First station, Closest to the Contact, Intermediate Diapir Side, Intermediate Cutler Side, Cutler Mouth Canyon, and Cutler Farthest Canyon* respectively. None of the stations had open fracture aperture, nor mineralization, but partly sulfuric oxidation.

3.1. Data acquisition

For the photos, the camera used in this study to obtain ground-based images of the exposure was a NIKON D5300. This camera is a DSLR with a 24.2 MP (megapixel) sensor. I used a 35mm, fixed focal length Nikkor lens (Online Reference 1).

For the ground control on the 3D models I placed surveying flags at GCP points and used the following tools. For high resolution positioning I used the DGPS Topcon GB1000 (the differential kinematic GPS set). It is a dual frequency and GPS+ receiver. It is one of the most advanced receivers and can combine real time kinematic measurements (RTK) accurately (Online Reference 2). Also, a Garmin eTrex10 was used as a handheld GPS which is common device with only limited differential correction from a WAAS signal.

I collected 76 ground control points using GPS equipment obtained in a kinematic survey with Topcon GB 1000, and with Garmin eTrex10 to compare the results and reference the data to a map. Topcon GB1000's base station was set up on the top of the North and East transects' hill (Figure 5). The data shown in Tables are in WGS84 UTM 12N coordinate system.

The original intent of the using the handheld GPS was as a backup GCP solution in the event of data loss as well as a tool to get real time positions to aid in finding markers on the model without the data processing needs of the DGPS data. This allowed, however, an evaluation of the handheld GPS accuracy in this site relative to the high- resolution kinematic survey (Table 1), which is discussed below.

For the photogrammetric models I took 3290 photos of the study area. The outcrop geometry made it advantageous to process the data in four “chunks” (Agisoft terminology for related pieces of a model that are more easily processed independently); 1) the Canyon chunk used 485 images - this site contains the fracture station called *Farthest/Canyon* Figure 22 (Figure 17), 2) the northern and eastern chunk 1428 images – this area has the rest of the fracture stations, 3) the western chunk used 608 images, and the southern chunk has two sets one taken in the early morning and another taken in the afternoon with 388, 381 images respectively – the two imaging times produced different shadow patterns that affected the result. Once the individual chunks were processed and georeferenced, the individual chunks were merged into a single model. For interpretation, however, visualizing the entire model at once challenged most computers, thus, the model was ultimately subdivided into segments using tools in CloudCompare.

For fracture measurements I chose six stations in advance from the diapir contact across complexly deformed rocks of Cutler Fm into intact Cutler Formation almost half a km (map distance on Google Earth) away from the diapir. At each station I measured strike and dip for all fractures visible at the site. For the fracture measurements, I followed Rohrbaugh et al. (2002)’s measurement methods. Specifically, I had three different sizes of circular scanlines that were located in the same center -radius 25cm, 50cm, and 75cm- instead of measuring every single fracture’s length. Then I collected 119 measurements, in total, from the same location with a Brunton compass and the smartphone to compare these two devices (Novakova and Pavlis, 2017; Allmendinger et al, 2017) (Figure 30). After the result of this test was positive, I started using the smartphone for the remainder of the data collection. A total of 412 fractures and bedding data were measured in the field (Figure 9 and Figure 10). These fracture measurement stations served

as reference sites for direct comparison of field measurements to measurements made on the 3D model. For the 3D model, however, the measurement areas were expanded beyond the limited radii of the specific stations because resolution was not always sufficient to make 1:1 comparisons.

Upon returning to the office, I processed the DGPS data and compared these data with handheld GPS (Tables 1-3). I used the photos obtained in the field and the DGPS data to process the imagery using Agisoft Photoscan by using the workflow of Matthews et al (2016). Once I made the models, I exported dense points of the models as “.las” format to import them to Maptek I-site Studio to digitize, collect same fracture measurements, and analyze from the same stations via remote analysis of the point cloud, using the geotech utilities in I-site studio. I imported the measurements from the phone to Stereonet 10 for further analysis.



Figure 8: Intermediate and Diapir Sections from Google Earth (the image is oblique view)

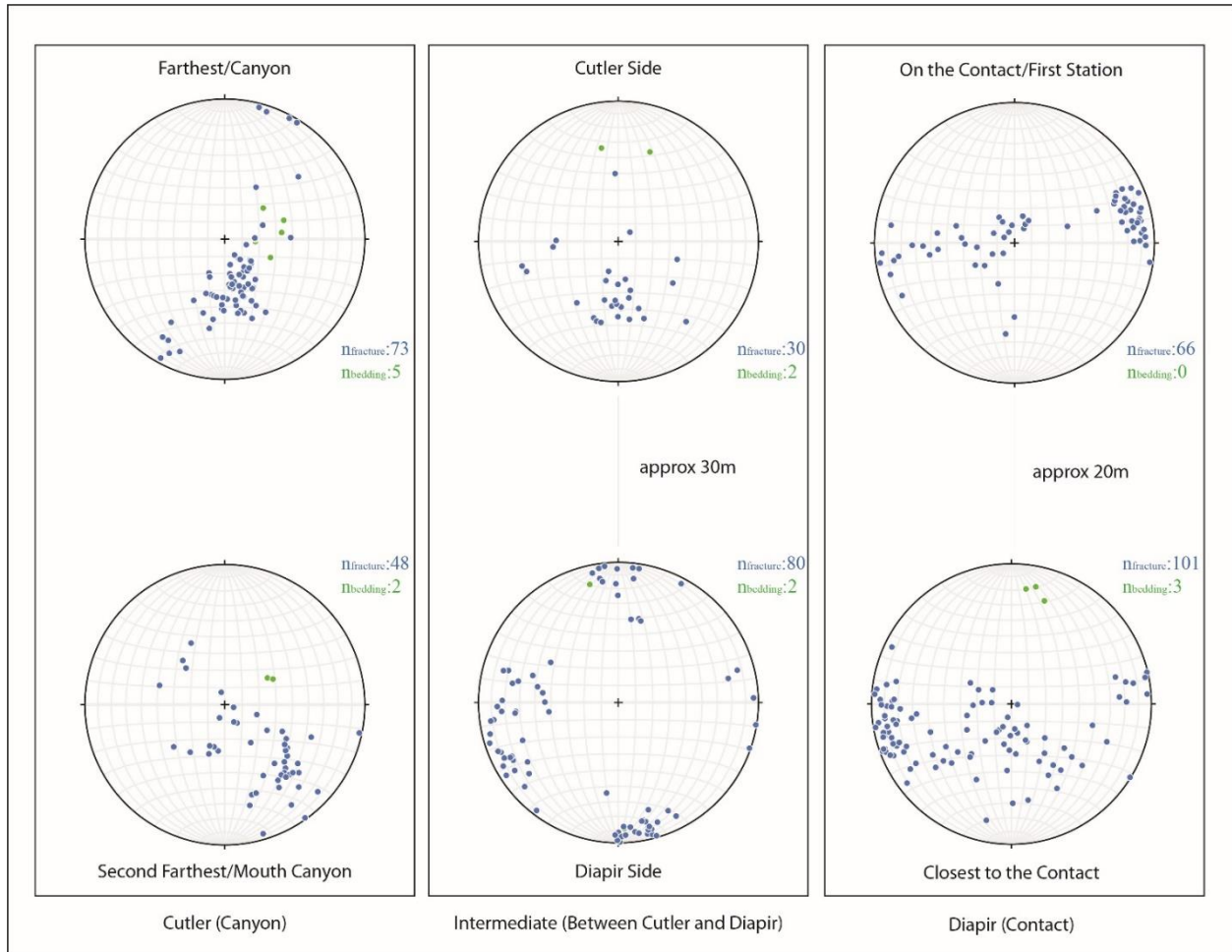


Figure 9: The stereonets illustrates bedding and fracture's strike and dip raw data from the field. n shows the number of the data point in the stereonets. (Figure 10 shows the location of the stations)

3.2. Results

3.2.1. DGPS vs WAAS corrected GPS positions on Ground Control Points

Looking at horizontal errors the UTM Northing and Easting errors can be used with Pythagorean theorem to estimate total horizontal error. The average error (Table 1) is 1.80m with a S.D. of 1.03m; which is consistent with the generally estimated error for WAAS corrected GPS (Online Reference 4). When I took elevation into account to estimate the total error, the error increases to 8-11m (Table 2-3). This is consistent with well-known errors in vertical vs

horizontal positioning with GPS, but these data revealed a surprising result. The reason there are two different tables for the elevation is because I compared the handheld GPS elevations to two reference elevations in the DGPS survey: ellipsoidal height (Table 2), which is calculated based on the theoretical ellipsoidal earth shape reference (Online Reference 5), and orthometric height (Table 3), which is calculated based on height above geoid, which is determined by the earth's gravity and approximates Mean Sea Level (MSL) at a given site (Online Reference 5), from the DGPS survey. The handheld GPS was set to record a WGS ellipsoidal height, and therefore should directly compare to Table 2. This comparison produces a systematic error of elevation in both cases. In Table 3, the error is consistent with a vertical datum mismatch with the orthometric height consistently above the handheld's height estimated for an ellipsoid; that is, regional low gravity over the continents produces a geoid high and thus, orthometric height should be above ellipsoid height as seen in the height estimates for the DGPS data between Tables 2 and 3. What is surprising is that in Table 2, where both heights are estimated relative to an ellipsoid, there is a systematic error between the handheld and DGPS with the handheld consistently above the DGPS estimate by an average of 8.7m. The origin of this discrepancy is unknown but is greater than generally accepted estimate for WAAS corrected GPS values.

Table 1: Comparison of GPS data points of survey and handheld (Elevation did not apply)

Name	Survey GPS		Hand held GPS		Northing defer	Easting defer	Distance defer	Avarage	Stand dev	STD error
	Northing	Easting	Northing	Easting						
OCT100	4284972.541	648754.5011	4284971.4	648754.7	1.140682	-0.198882	1.157890125	1.804941727	1.029220175	0.19807352
OCT101	4284967.101	648755.1168	4284968	648753.9	-0.898646	1.2168092	1.512676014			
OCT102	4284949.546	648732.5809	4284949.9	648732.5	-0.354208	0.0808675	0.363321655			
OCT103	4285000.04	648721.7588	4284999.6	648721.1	0.4403371	0.6587753	0.792389856			
OCT104	4285033.919	648728.5245	4285034.1	648725.7	-0.18055	2.8244746	2.830239398			
OCT105	4285051.448	648724.7789	4285052.9	648724.5	-1.452404	0.2789321	1.478945546			
OCT106	4285012.751	648718.2036	4285014	648717.4	-1.248997	0.8036248	1.48519568			
OCT107	4285022.346	648710.3281	4285023.8	648711.1	-1.453895	-0.771855	1.646077281			
OCT108	4285022.968	648710.6072	4285023.8	648711.1	-0.831672	-0.492787	0.966704174			
OCT109	4285065.595	648719.7684	4285067.2	648717.3	-1.604523	2.4684384	2.944093005			
OCT110	4284969.646	648712.2678	4284968.4	648713	1.2459687	-0.73224	1.445203418			
OCT111	4284948.441	648720.0655	4284950.7	648719.5	-2.259174	0.5655008	2.328874775			
OCT112	4284944.163	648713.0959	4284943.9	648713.5	0.2625641	-0.404064	0.481879164			
OCT113	4284956.985	648760.8691	4284957	648759.4	-0.015044	1.4691371	1.469214106			
OCT114	4284952.842	648783.3488	4284954.1	648781.2	-1.257876	2.1488322	2.489926176			
OCT115	4284971.811	648798.8434	4284972.2	648798.2	-0.388505	0.6433691	0.751571765			
OCT116	4284984.104	648828.3711	4284982.7	648827.6	1.4038431	0.7710928	1.601674048			
OCT117	4284968.034	648858.1232	4284966.6	648859.2	1.434269	-1.07684	1.793519157			
OCT118	4284964.778	648857.8477	4284963.3	648858.4	1.4784232	-0.552314	1.578222225			
OCT119	4284952.032	648783.8294	4284950.8	648781.2	1.2319212	2.629429	2.903709103			
OCT120	4284973.552	648786.0534	4284973	648784.3	0.5516777	1.7533818	1.838122936			
OCT200	4284970.287	648755.4923	4284972.5	648755.6	-2.213097	-0.107679	2.215715492			
OCT201	4284997.075	648810.3205	4284999	648810.8	-1.924713	-0.479545	1.983553407			
OCT202	4284949.104	648707.4598	4284950.5	648706.4	-1.395973	1.0598212	1.752701506			
OCT203	4284980.548	648708.2958	4284981.6	648708.4	-1.051765	-0.104167	1.056910821			
OCT204	4285035.49	648710.9313	4285040.4	648708.2	-4.91003	2.7312575	5.618555206			
OCT205	4285066.161	648715.7119	4285068.3	648716.4	-2.138565	-0.688101	2.246540598			

Table 2: Comparison of GPS data points of survey and handheld according Ellipsoidal Height

Name	Survey GPS			Hand held GPS			Northing defer	Easting defer	Elevation defer	Distance defer	Avarage	Stand dev	STD error
	Northing	Easting	Ellipsoidal Ht.	Northing	Easting	Elevation							
OCT100	4284972.541	648754.5011	1440.34	4284971.4	648754.7	1449	1.14068201	-0.19888211	-8.66	8.73706527	8.71884	2.39669	0.46124
OCT101	4284967.101	648755.1168	1439.965	4284968	648753.9	1450	-0.89864574	1.21680917	-10.035	10.14837			
OCT102	4284949.546	648732.5809	1438.089	4284949.9	648732.5	1445	-0.35420767	0.08086749	-6.911	6.92054359			
OCT103	4285000.04	648721.7588	1436.151	4284999.6	648721.1	1443	0.44033707	0.65877534	-6.849	6.8946851			
OCT104	4285033.919	648728.5245	1435.472	4285034.1	648725.7	1448	-0.18055026	2.82447458	-12.528	12.8437159			
OCT105	4285051.448	648724.7789	1430.803	4285052.9	648724.5	1442	-1.452403801	0.27893212	-11.197	11.2942503			
OCT106	4285012.751	648718.2036	1430.85	4285014	648717.4	1440	-1.24899696	0.80362479	-9.15	9.26975222			
OCT107	4285022.346	648710.3281	1430.654	4285023.8	648711.1	1438	-1.45389498	-0.77185478	-7.346	7.52816621			
OCT108	4285022.968	648710.6072	1430.632	4285023.8	648711.1	1438	-0.83167174	-0.49278705	-7.368	7.43114668			
OCT109	4285065.595	648719.7684	1428.975	4285067.2	648717.3	1437	-1.60452346	2.46843839	-8.025	8.54800027			
OCT110	4284969.646	648712.2678	1432.436	4284968.4	648713	1444	1.245968729	-0.73223961	-11.564	11.6539568			
OCT111	4284948.441	648720.0655	1432.58	4284950.7	648719.5	1442	-2.25917386	0.56550083	-9.42	9.70361055			
OCT112	4284944.163	648713.0959	1432.539	4284943.9	648713.5	1443	0.26256409	-0.40406389	-10.461	10.4720928			
OCT113	4284956.985	648760.8691	1434.831	4284957	648759.4	1445	-0.01504426	1.46913708	-10.169	10.2745876			
OCT114	4284952.842	648783.3488	1434.085	4284954.1	648781.2	1442	-1.257876219	2.14883219	-7.915	8.29740666			
OCT115	4284971.811	648798.8434	1434.934	4284972.2	648798.2	1442	-0.38850522	0.64336911	-7.066	7.10585787			
OCT116	4284984.104	648828.3711	1436.484	4284982.7	648827.6	1448	1.40384315	0.77109284	-11.516	11.6268489			
OCT117	4284968.034	648858.1232	1436.101	4284966.6	648859.2	1440	1.43426895	-1.07683961	-3.899	4.29172599			
OCT118	4284964.778	648857.8477	1441.039	4284963.3	648858.4	1440	1.47842318	-0.55231358	1.039	1.88952544			
OCT119	4284952.032	648783.8294	1436.324	4284950.8	648781.2	1445	1.23192118	2.62942898	-8.676	9.14901648			
OCT120	4284973.552	648786.0534	1441.276	4284973	648784.3	1450	0.55167771	1.75338177	-8.724	8.91554103			
OCT200	4284970.287	648755.4923	1440.29	4284972.5	648755.6	1449	-2.21309745	-0.10767924	-8.71	8.98740759			
OCT201	4284997.075	648810.3205	1454.902	4284999	648810.8	1464	-1.92471306	-0.47954536	-9.098	9.31171779			
OCT202	4284949.104	648707.4598	1441.353	4284950.5	648706.4	1449	-1.39597338	1.05982116	-7.647	7.84528977			
OCT203	4284980.548	648708.2958	1437.071	4284981.6	648708.4	1447	-1.05176509	-0.10416659	-9.929	9.98509396			
OCT204	4285035.49	648710.9313	1436.432	4285040.4	648708.2	1446	-4.910030031	2.73125753	-9.568	11.09571103			
OCT205	4285066.161	648715.7119	1434.324	4285068.3	648716.4	1439	-2.13856542	-0.68810072	-4.676	5.18767006			

Table 3: Comparison of GPS data points of survey and handheld according Orthometric Height

Name	Survey GPS			Hand held GPS			Northing defer	Easting defer	Elevation defer	Distance defer	Avarage	Stand dev	STD error
	Northing	Easting	Orthometric Ht.	Northing	Easting	Elevation							
OCT100	4284972.541	648754.5011	1459.796	4284971.4	648754.7	1449	1.14068201	-0.19888211	10.796	10.8579153	11.2859	2.679	0.51557
OCT101	4284967.101	648755.1168	1459.421	4284968	648753.9	1450	-0.89864574	1.21680917	9.421	9.54166808			
OCT102	4284949.546	648732.5809	1457.546	4284949.9	648732.5	1445	-0.35420767	0.08086749	12.546	12.5512596			
OCT103	4285000.04	648721.7588	1455.61	4284999.6	648721.1	1443	0.44033707	0.65877534	12.61	12.6348717			
OCT104	4285033.919	648728.5245	1454.931	4285034.1	648725.7	1448	-0.18055026	2.82447458	6.931	7.48658908			
OCT105	4285051.448	648724.7789	1450.262	4285052.9	648724.5	1442	-1.452403801	0.27893212	8.262	8.39332615			
OCT106	4285012.751	648718.2036	1450.309	4285014	648717.4	1440	-1.24899696	0.80362479	10.309	10.415435			
OCT107	4285022.346	648710.3281	1450.113	4285023.8	648711.1	1438	-1.45389498	-0.77185478	12.113	12.2243339			
OCT108	4285022.968	648710.6072	1450.091	4285023.8	648711.1	1438	-0.83167174	-0.49278705	12.091	12.1295836			
OCT109	4285065.595	648719.7684	1448.435	4285067.2	648717.3	1437	-1.60452346	2.46843839	11.435	11.8079172			
OCT110	4284969.646	648712.2678	1451.894	4284968.4	648713	1444	1.245968729	-0.73223961	7.894	8.02520086			
OCT111	4284948.441	648720.0655	1452.037	4284950.7	648719.5	1442	-2.25917386	0.56550083	10.037	10.3036414			
OCT112	4284944.163	648713.0959	1451.997	4284943.9	648713.5	1443	0.26256409	-0.40406389	8.997	9.00989548			
OCT113	4284956.985	648760.8691	1454.286	4284957	648759.4	1445	-0.01504426	1.46913708	9.286	9.40150978			
OCT114	4284952.842	648783.3488	1453.539	4284954.1	648781.2	1442	-1.257876219	2.14883219	11.539	11.8045861			
OCT115	4284971.811	648798.8434	1454.387	4284972.2	648798.2	1442	-0.38850522	0.64336911	12.387	12.4097796			
OCT116	4284984.104	648828.3711	1455.936	4284982.7	648827.6	1448	1.40384315	0.77109284	7.936	8.09601481			
OCT117	4284968.034	648858.1232	1455.551	4284966.6	648859.2	1440	1.43426895	-1.07683961	15.551	15.6540829			
OCT118	4284964.778	648857.8477	1460.489	4284963.3	648858.4	1440	1.47842318	-0.55231358	20.489	20.5496936			
OCT119	4284952.032	648783.8294	1455.778	4284950.8	648781.2	1445	1.23192118	2.62942898	10.778	11.1622941			
OCT120	4284973.552	648786.0534	1460.73	4284973	648784.3	1450	0.55167771	1.75338177	10.73	10.8863031			
OCT200	4284970.287	648755.4923	1459.746	4284972.5	648755.6	1449	-2.21309745	-0.10767924	10.746	10.9720514			
OCT201	4284997.075	648810.3205	1474.356	4284999	648810.8	1464	-1.92471306	-0.47954536	10.356	10.5442506			
OCT202	4284949.104	648707.4598	1460.811	4284950.5	648706.4	1449	-1.39597338	1.05982116	11.811	11.9403385			
OCT203	4284980.548	648708.2958	1456.529	4284981.6	648708.4	1447	-1.05176509	-0.10416659	9.529	9.58743456			
OCT204	4285035.49	648710.9313	1455.891	4285040.4	648708.2	1446	-4.910030031	2.73125753	9.891	11.375414			
OCT205	4285066.161	648715.7119	1453.784	4285068.3	648716.4	1439	-2.13856542	-0.68810072	14.784	14.9537153			

3.2.2. Geology of the Study Area

The Onion Creek area exhumed an elongate salt diapir with an ~W-E long axis. The study area is along the western edge of the Onion Creek salt diapir core exposures of gypsum, but structures related to the diapir continue well westward from the study site. To the north of the diapir is intact, but deformed Cutler Fm, which is distinctly folded near the diapir while overlying strata typically dip northward, away from the diapir (Figure 5). In the study area the Cutler Fm is in fault contact with complexly deformed rocks near the contact. In our work in this area, we informally referred to these complexly deformed rocks as Domain 0, which forms a mappable unit along the diapir contact. The rocks in the study area between gypsum exposures and the intact Cutler can all be considered part of Domain 0. This study was chosen because of excellent rock exposures in the river cuts. Although not exposed, it is clear that the diapir underlies the exposures of Domain 0 in the study area. From diapir contact, across Domain 0 to

the Cutler Fm, the rocks are highly fractured and faulted, bedding orientations change abruptly across faults, and fractures vary between fault blocks.

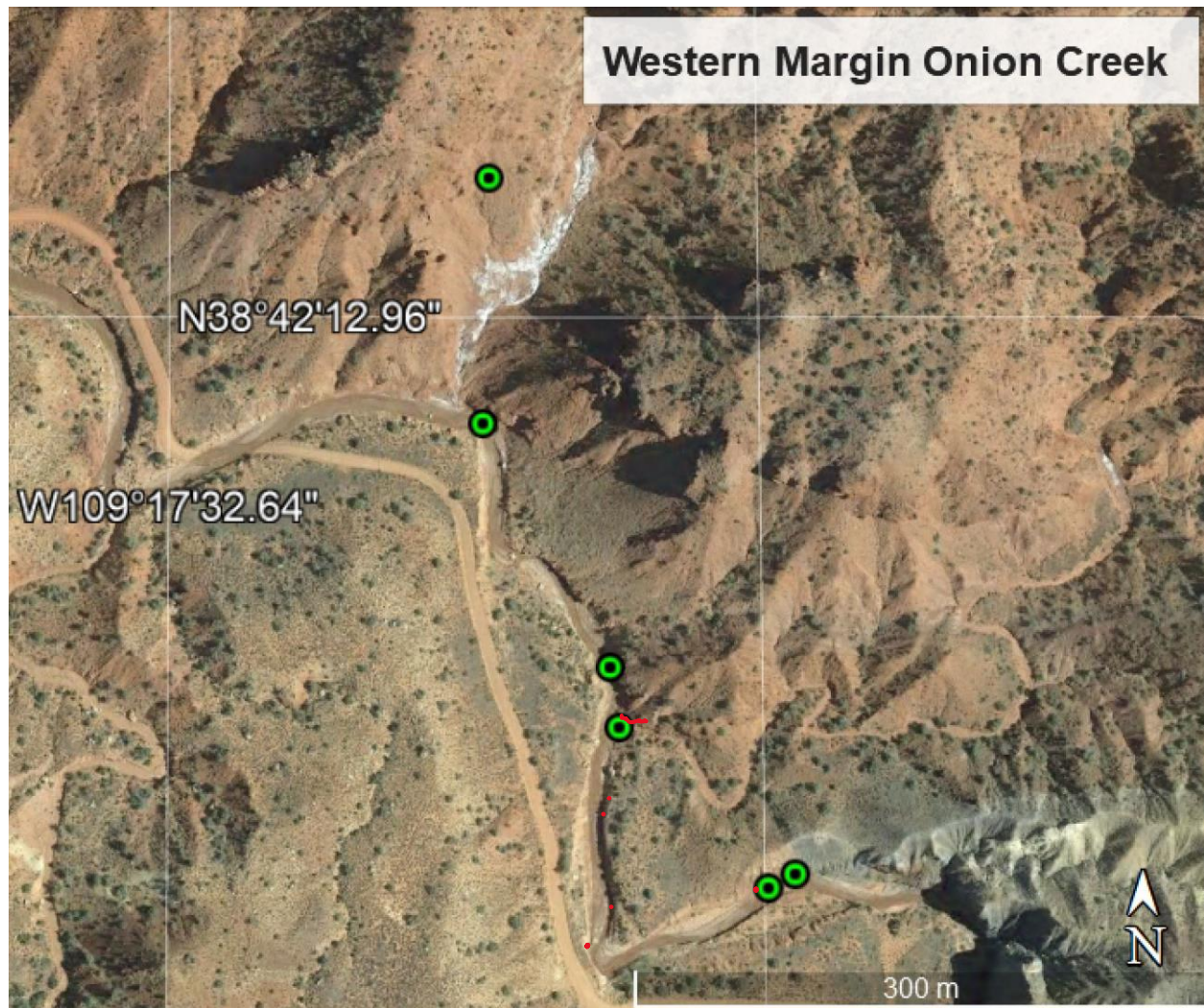


Figure 10: The study area from Google Earth (red dots show the intact fault locations while green show the location of the stations)

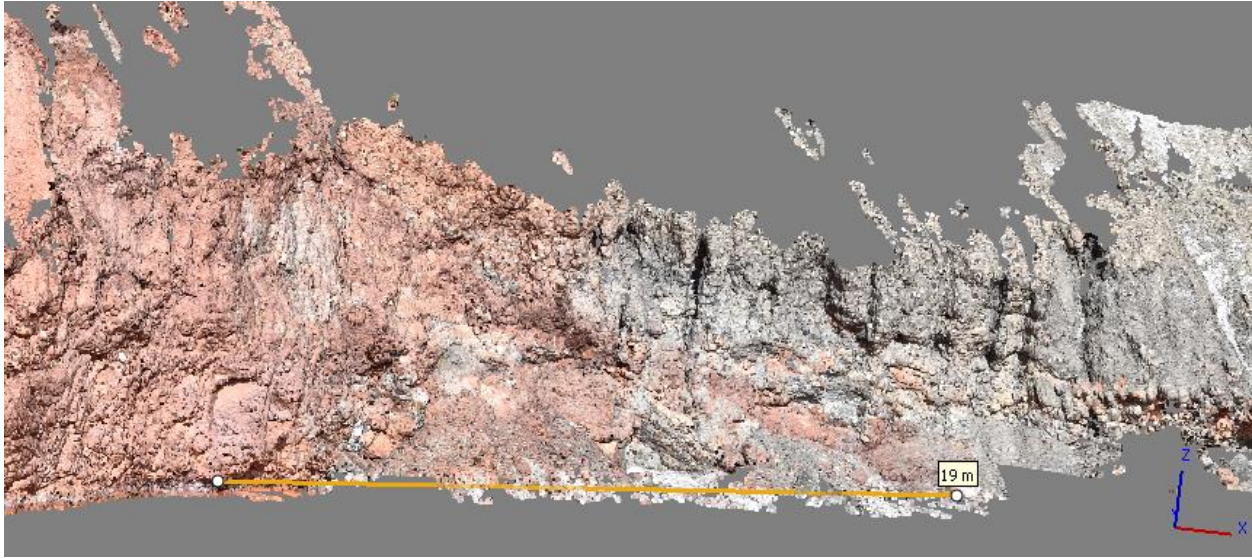


Figure 11: The figure shows the contact of the salt diapir with Domain 0. The orange line joints the two fracture stations “On the Contact” to the right and “Closest to the Contact” on the left, a distance of 19m

Figure 12 shows a part of the 3D model at the diapir contact looking NNW. Here the bedding-strikes NE and dips steeply northwest. Approximately 20 m west of the contact, (Figure 11-Figure 13) bedding changes abruptly to ~EW strikes with south dips. These dip changes indicate distinct faults cutting the section, but in this area bedding is typically indistinct, suggesting extensive cataclastic across this 20-30m side zone along the contact.

Figure 14 shows a view of the 3D model at the third fracture analysis site (labeled Intermediate-Diapir side Figure 10). This site is located just south of the fault contact with the Cutler Fm. At this site a fault is present in the deep shadow part of the model (Figure 14) but the fault is recognized by bedding truncations and abrupt change in strike and dip of bedding. Unfortunately, the field fracture station is not visible in the 3D model due to a gap in the model produced by deep shadow under the big rock on the right hand of the figure. The actual station is under the red circle shown on Figure 14.

Figure 15 shows a view of the 3D model at the fourth station (Intermediate-Cutler Side, Figure 10) located within the Cutler Fm outside the domain 0 zone. At this site bedding strikes ~EW and dips moderately south. Fracture sets can be clearly seen intersecting bedding at this site (Figure 15) with a sub horizontal intersection line.

Figure 16 and Figure 17 show the 3D model at the sites “Cutler, Second Farthest” and “Cutler Farthest” respectively. Both of these sites show similar bedding orientations to that seen in Figure 15Figure 17, with similar prominent fracture systems visible at high angle to bedding planes. These last three sites proved critical for comparisons of fracture orientations determined on the model and in the field because the fractures form prominent 3D surface visible on the model (Figure 15Figure 17).

There were many faults within the domain 0 part of the study area but these faults are difficult to trace beyond the limits of the outcrop due to alluvial cover separating the two cliff-walls along the creek bed. Figure 18, from the western river cut at the edge, there is a remarkable fault (shown on the map, Figure 8) in this study area. This fault is marked by a distinct gouge zone, ~15 cm thick, with the gouge strongly affected by a Sulphur oxidation (Figure 18). The result of this fluid interaction is the development of a very soft and wet clay gouge. Presumably this alteration reflects fluid migration from the diapir.

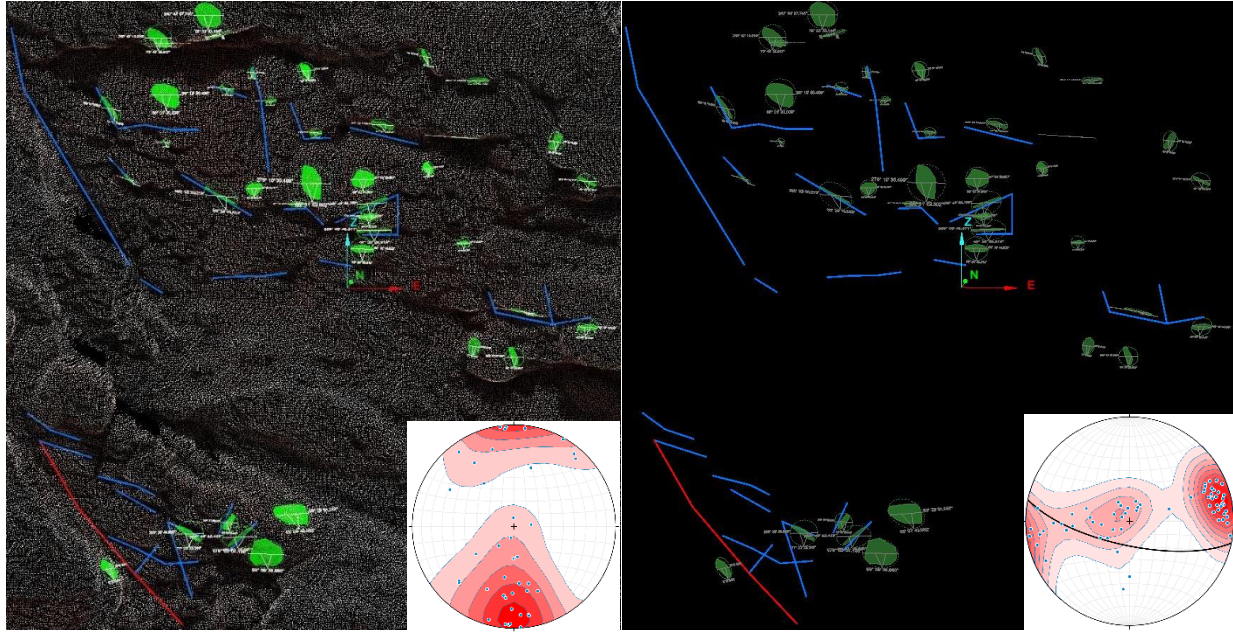


Figure 12: Diapir contact focus site with view zoomed in to the exact area of the field fracture measurements. Left figure shows point cloud with measurements where the right figure shows only patches for orientation estimates. Note the relative sparsity of the point cloud at this zoom level. Figure shows a visualization of remotely sensed fracture orientations developed from 3D model as well as mapped traces of fractures (blue) and faults (red). The stereoplots show the data collected from 3D model (Left) and in situ (right) (see Appendix 1-2 for the data).

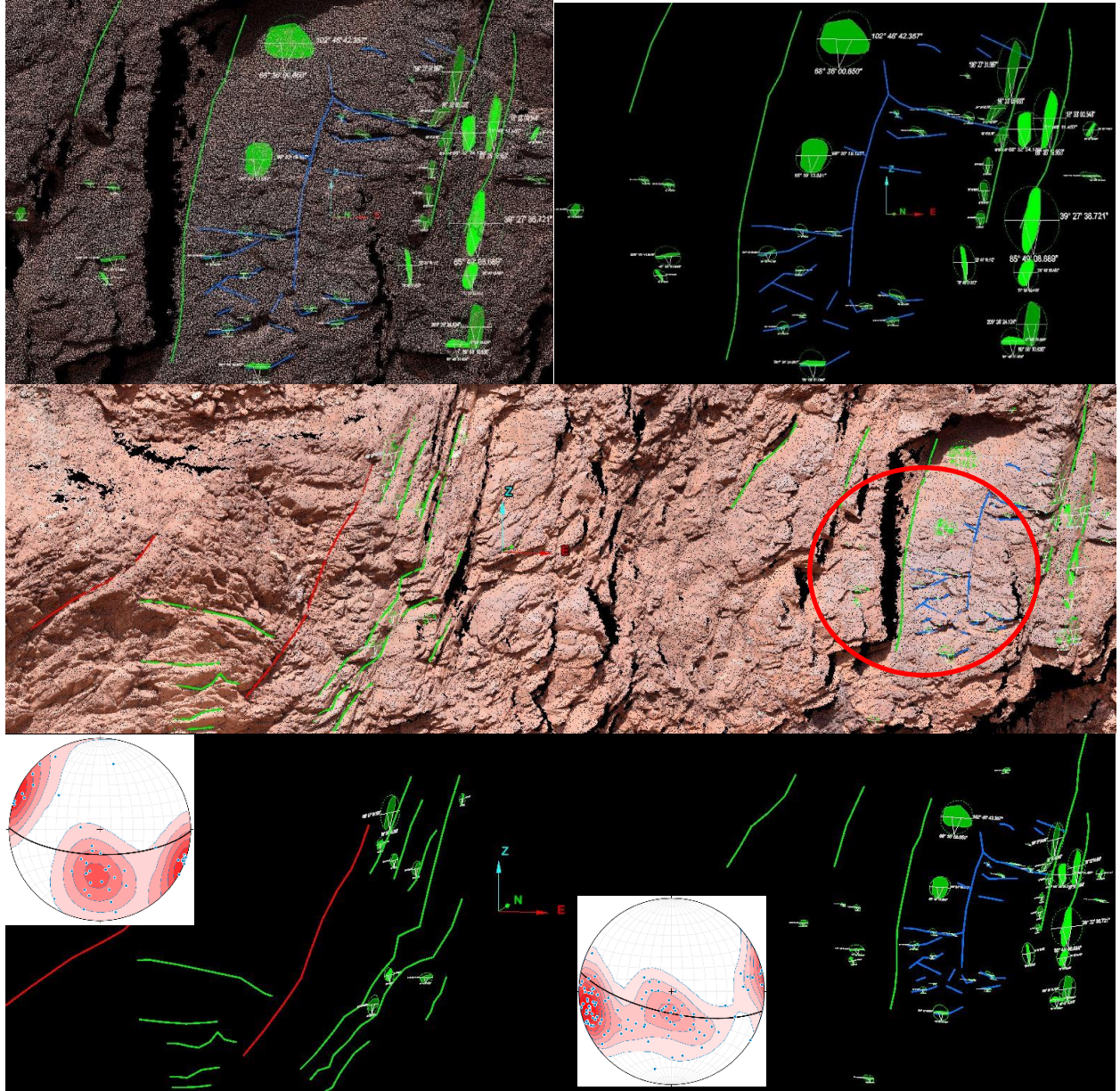
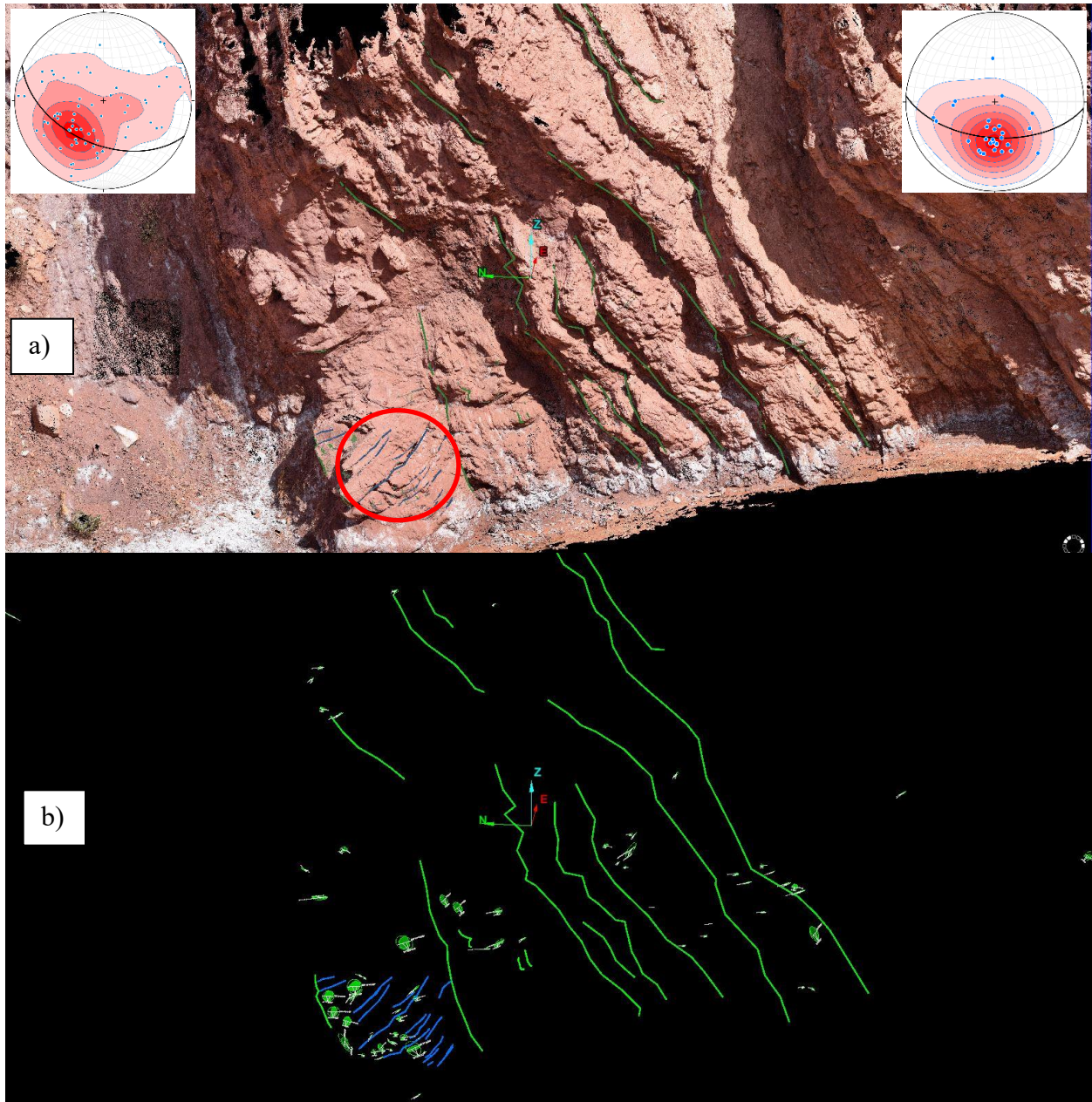


Figure 13: Closest to the Contact focus site (second fracture station from the diapir, about 20m). Center and lower thirds of the figure are a visualization of ~20m of the model near the fracture focus site outlined by the red circle. Central figure shows the model with interpreted traces of bedding (green), fractures (blue) and faults (red, with fractures only mapped in the focus site (red circle). The lower 1/3 of the figure is the same view as the central figure with the point cloud visualization removed and showing patches where fractures were measured on the model. Upper third of the figure is a zoomed in view of the area of the field measurements with the point cloud include on the left and interpretation only on the right (color scheme and green patches are the same as the lower figure). The

stereoplot insets show the fracture data collected from 3D model (left) and in situ (right). All views are looking to the north along an approximately horizontal line. (see Appendix 1-2 for the data)



Figure 14: Intermediate-Diapir side focus site (third station from the diapir). The color is the same as Figure 12 with fracture traces (blue) and beds (green). The red circle shows the approximate fracture station that collected data in situ, but Photoscan left as a gap because of deep shadow (see text for discussion)



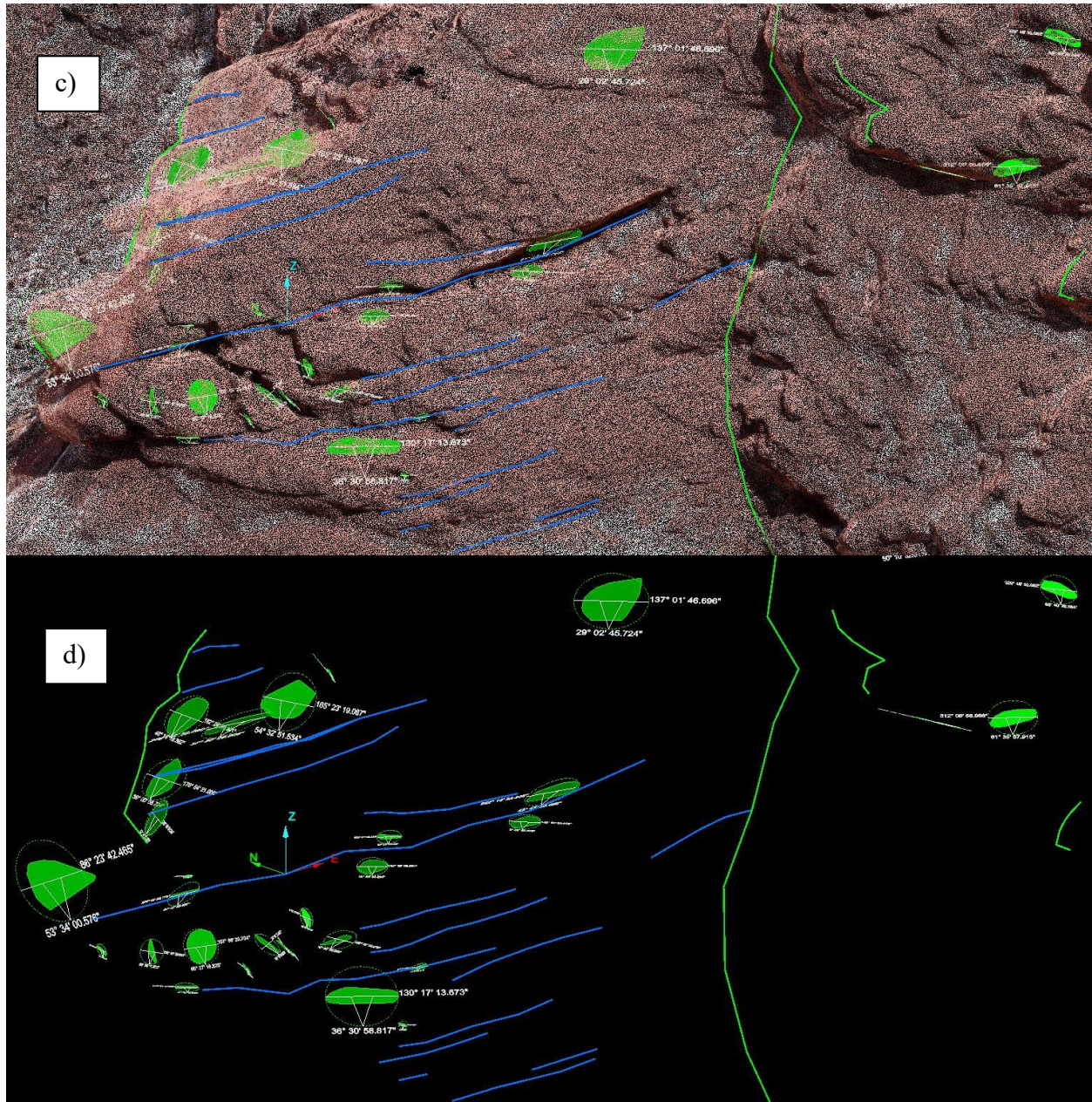
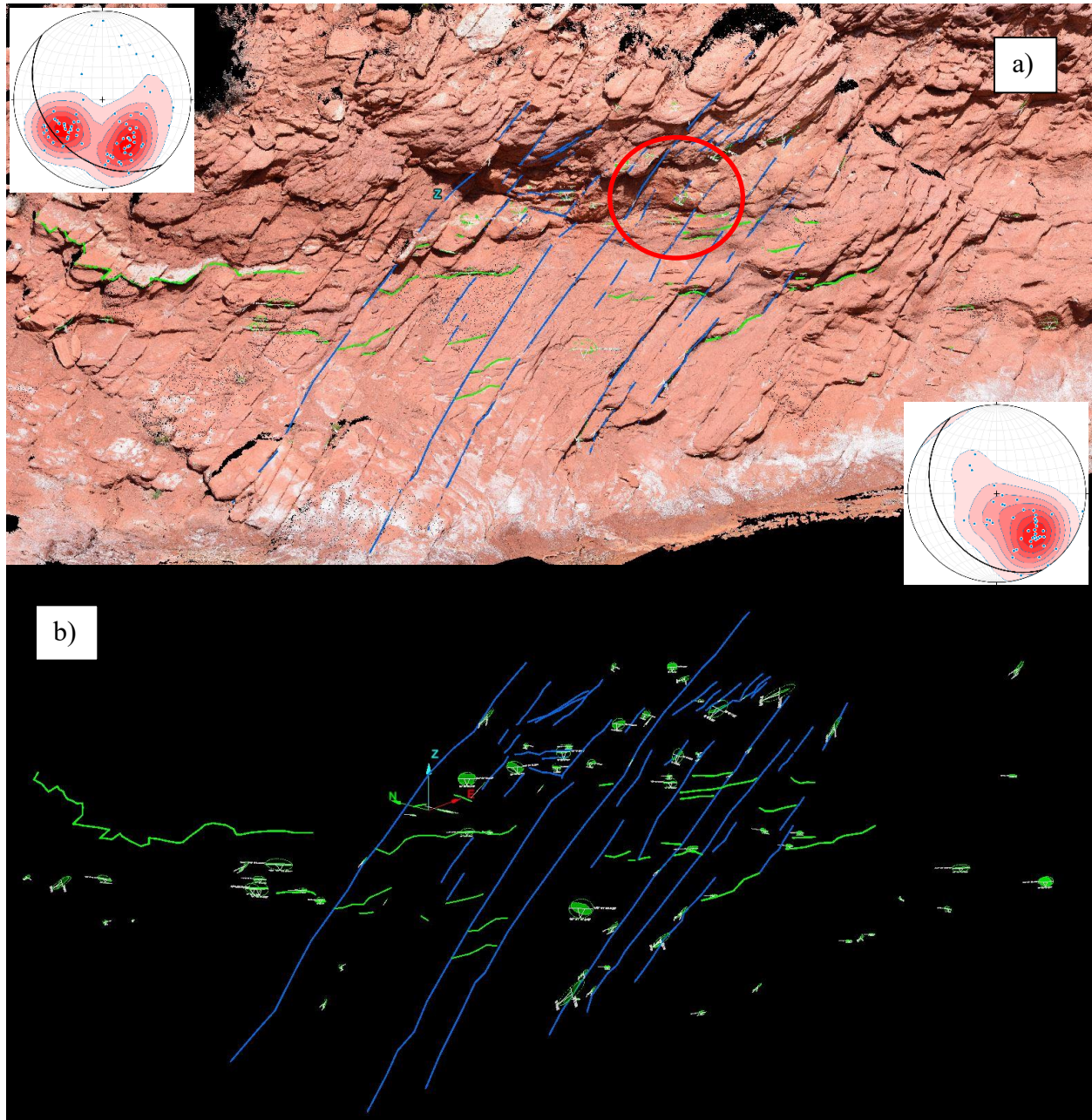


Figure 15: Intermediate Cutler Side focus site (the fourth station from the diapir). Upper pair of figures (a and b) show zoomed out view of ~30m of the outcrop with bedding (green) and fractures (blue) traced on the outcrop (fractures only mapped at the field in-situ site, outlined as the red circle). Upper figure (a) is a visualization of the point cloud with interpretations on it and the lower figure (b) shows the same area with the point cloud turned off. Lower two figures (c and d) show an enlarged view of the area of the field measurements corresponding to the red circle in the upper figures. Color scheme of lower figure is the same as upper figure. Green patches in all figures

are sites of fractures measured on the model using multipoint tools in I-site studio. The stereoplot inserts in a show the fracture orientation data collected from 3D model (left) and in situ (right). (see Appendix 1-2 for the data)



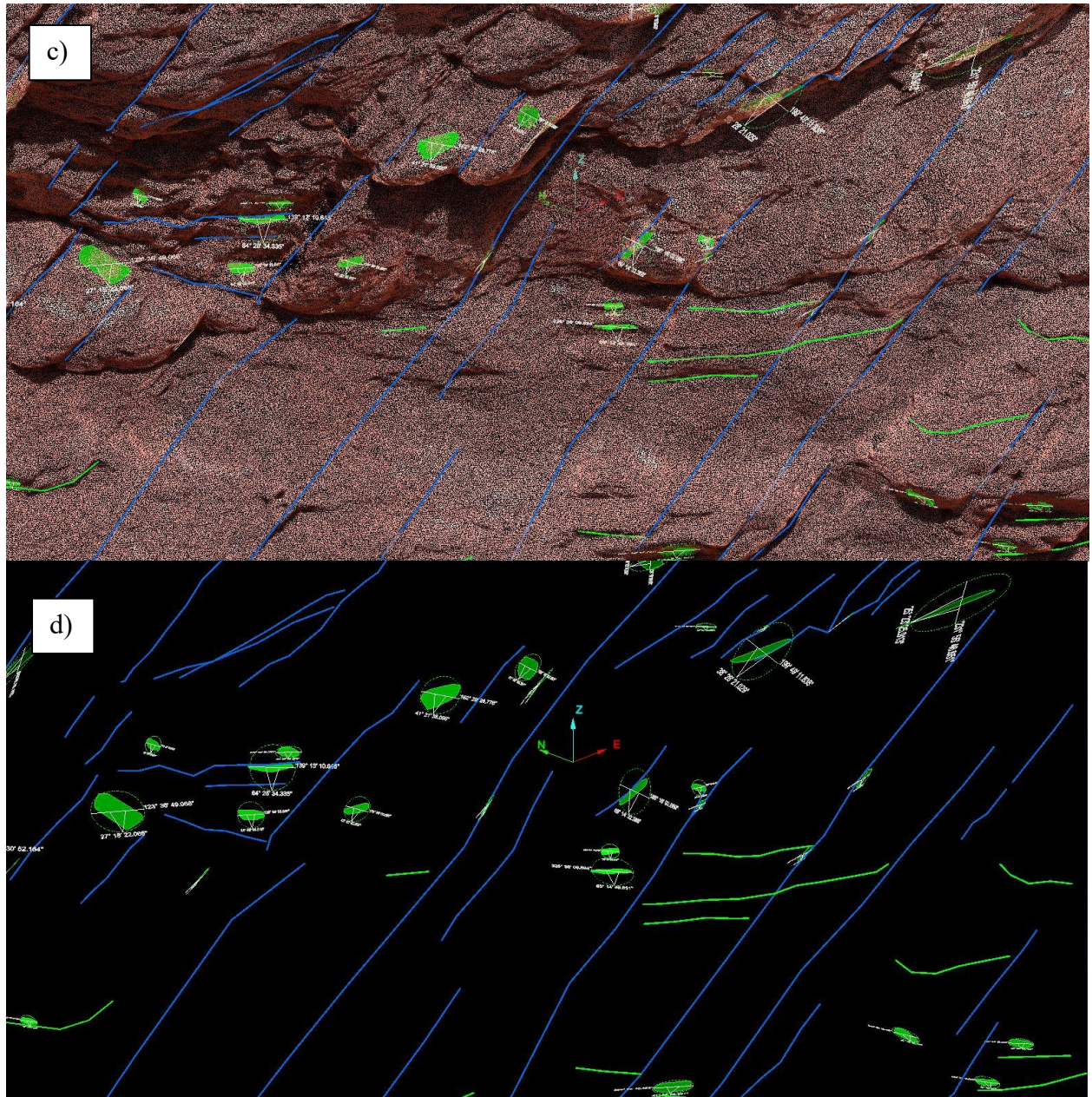
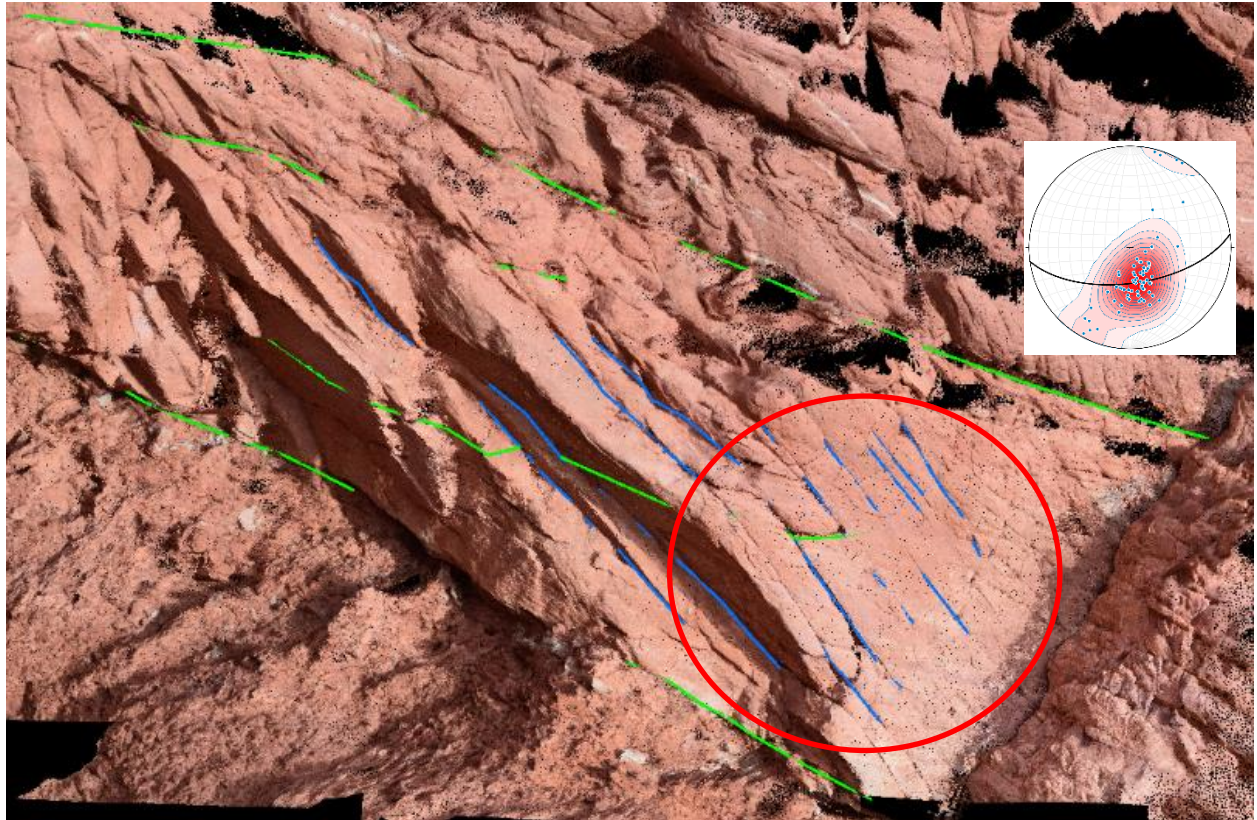


Figure 16: Cutler-Second Farthest focus site (the fifth station from the diapir on the mouth of the Canyon). Upper pair of figures (a and b) show zoomed out view of ~10m of the outcrop with bedding (green) and fractures (blue) traced on the outcrop. Red circle shows the area where field data were acquired and enlarged in the lower figures. Upper figure (a) is a visualization of the point cloud with interpretations on it and the lower figure (b) shows the same area with the point cloud turned off. Lower two figures (c and d) show an enlarged view of the area of the field measurements corresponding to the red circle in the upper figures. Color scheme of lower figure is the same as upper figure. Green patches in all figures are sites of fractures measured on the model using multipoint tools in

I-site studio. The stereoplot inserts in show the fracture orientation data collected from 3D model (left) and in situ (right). (see Appendix 1-2 for the data).



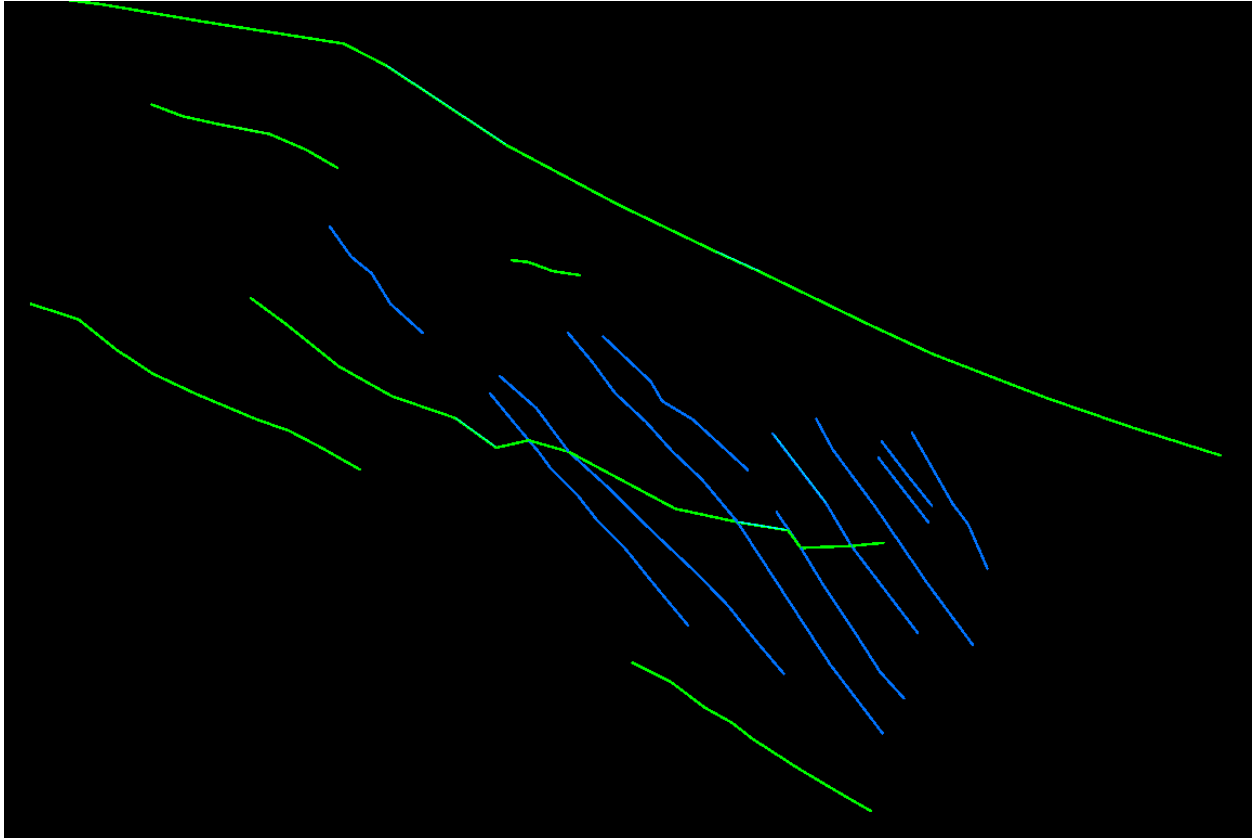


Figure 17: Cutler-Farthest focus site (the sixth station in the canyon on the cliff from the diapir). Upper Figure shows a visualization of the point cloud with interpretations drawn on it and lower figure shows the interpretation without the point cloud. Red circle shows the position of the in-situ field measurements at this site. Stereograms show orientations measured in situ only. This site could not be properly georeferenced for direct comparison between field and model data. (see text for discussion) (see Appendix I for the data).



Figure 18: The fault filled with Sulphur oxidation (see Figure 8)

3.2.3. Analysis of Fracture Orientations

After generating the 3D models in Agisoft Photoscan, I exported the dense point cloud as “.las” format to import them to CloudCompare to merge the data into a single project. I then exported again in the same format for importing to Maptek I-site Studio to digitize and analyze the fracture systems. I-site studio has a “geotech” utility which allows multipoint analysis of a point cloud to determine a best orientation for a plane fitting these points. Thus, this method is an extension of the classic 3-point problem to multi-point analyses. This technique seemed an ideal use of the 3D model to supplement field measurements and I attempted to use the method at several sites within the 3D model.

Although the method seemed ideal, several problems were encountered during this analysis;

1) There was a problem with one of the focus sites from the Intermediate section-*Diapir Side* on the 3D models. At the site, despite care with photo overlap and close spacing, there is a hole in the point cloud, primarily due to shadow (Figure 14), so it left a gap on the model that could not be analyzed as a direct comparison between the field data and model derived data.

2) The 3D model of *Cutler-Farthest* (Figure 10) was run by itself because it was geographically isolated from the main outcrop along the creek. Despite numerous processing schemes to improve this model, it could not be properly georeferenced. The origin of the problem is unknown. A well-defined model was generated relatively easily (Figure 17) and allowed mapping directly on the model, but when ground control points were added to the model they showed large spatial errors and the resultant model was rotated by large amounts from the real world. The origin of this problem is unknown, but it is likely one or more of the GCP's either has a bad position or was somehow mislocated on the model. Thus, I was unable to directly compare field measurements to model measurements at this site.

3) I digitized fractures from the diapir contact site twice, (*On the Contact* in Figure 12 and Figure 20). After digitizing the fractures and using the fracture traces to estimate fracture orientation using the multipoint tool, I made a plot of the data on an equal area stereoplot to compare it with the stereoplot of the data collected for this site in the field (Figure 19). Figure 12 and Figure 19 show the comparison of the stereoplots with a kamb contoured plot and scatter plot respectively. These plots demonstrate that the remote measurement produced a very unsatisfactory result, with apparent fracture orientations nearly 90 degrees from the observed and none of the field fractures recognized. The origin of the problem is discussed below.

Despite these problems, fracture analysis on the 3D model at all other sites produced satisfactory results, and arguably, better results, than field measurements, for reasons discussed

below. Figure 13 and Figure 15 show *Closest to the Contact* and *Intermediate Cutler Side* stations where the results of model based measurements are comparable to the field data, but slightly different. Specifically, at both sites the general fracture patterns are similar in the two methods; in Figure 13, both stereonet plots reveal the pattern seen on the 3D model of two, approximately orthogonal fracture sets perpendicular to bedding and in Figure 15 only one fracture set, perpendicular to bedding, is shown. However, the positions of the maxima in the stereonet plots are clearly shifted slightly between field based and model-based measurements; e.g. in Figure 15, the maxima for poles to fractures in the field based data is located at S (181/64) whereas the maxima from the model is at SW, an angle of $\sim 206/76$. Similarly, Figure 13, the field based maxima in the poles to fracture plots are at W (260/12) and S (176/70) vs corresponding maxima on the model based plots at NW (111/02) and S (185/50), which differ by angles of 149/10 and 351/20 respectively. Note that although the scalar angular differences among these plots are similar, the direction of the shifts is not, suggesting strongly the source of the distinction is a random, presumably measurement based, error in one or both methods. This conclusion on measurement error is supported by bedding measurements in both the field and on the model, which unlike the fracture data are indistinguishable great circles on the plots (Figure 13 and Figure 15).

Perhaps the most intriguing results is at the “Cutler--second farthest site” (Figure 16). Here, I did the same procedure, with the field and model measurements, and the main fracture set (mapped in blue in Figure 16) yielded indistinguishable orientations using the two methods (pole maxima at 137/35 for the field data and 147/43 for the model based measurements). Similarly, bedding orientations match perfectly for field vs model-based measurements (great circles in the stereonet plots, Figure 16). What is different is the model based measurements revealed a clear,

second fracture set at this site, with a maxima in the pole plot at 236/43, or ~90 degrees from the other fracture set. This fracture set is clearly visible in the larger scale view of the outcrop (upper figures, in Figure 16) as the sub-horizontal steps in the bedding surface.

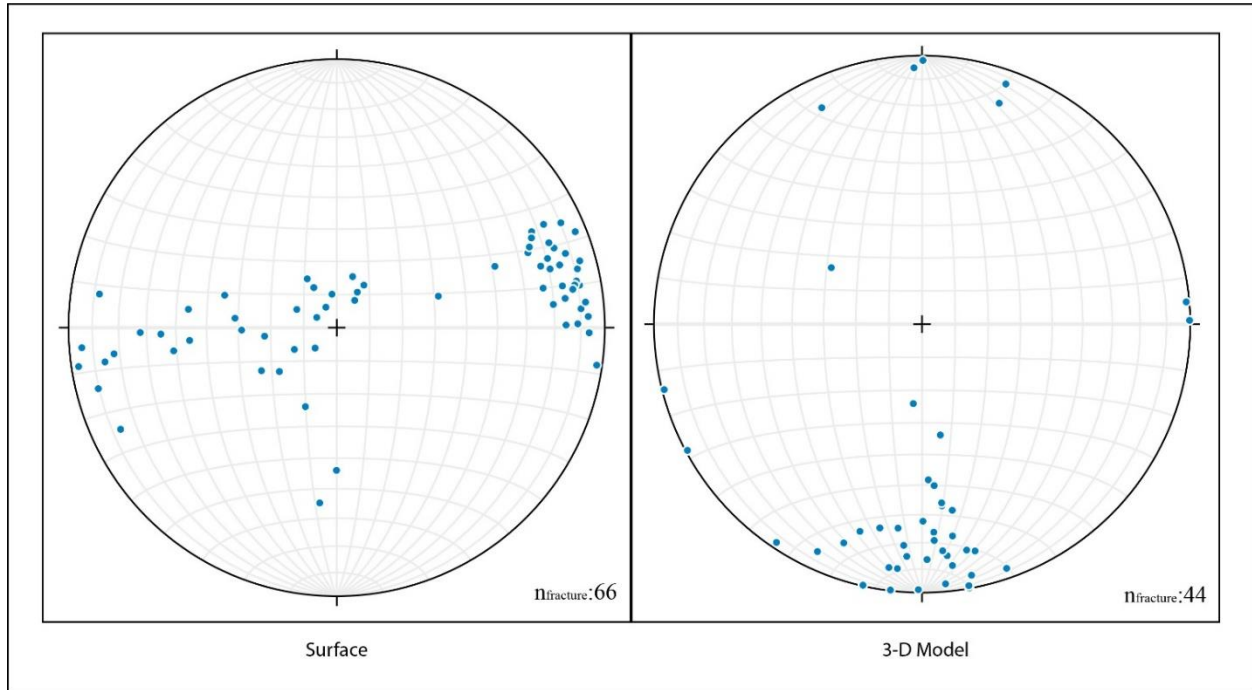


Figure 19: The comparison of surface and 3-D model On the Contact focus site (see Appendix 4 for the data)

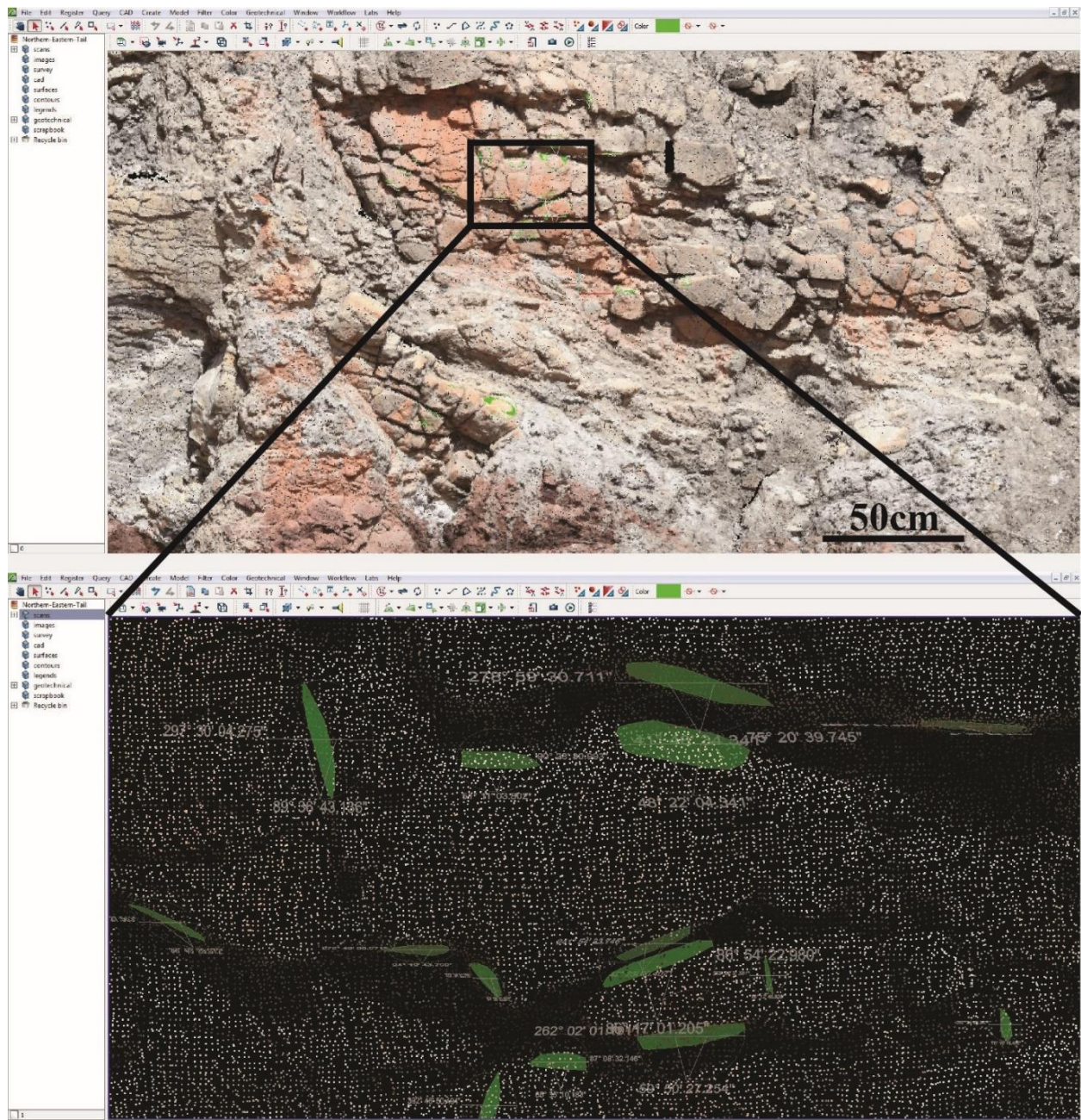


Figure 20: Diapir contact focus site (the first fracture station from the diapir). The images show the strike and dip measurements of the fractures (green) on the 3D model. When zoomed in, there is a pixel issue on the Maptek. Note that the pixels are larger than original (see Appendix 4 for the data)

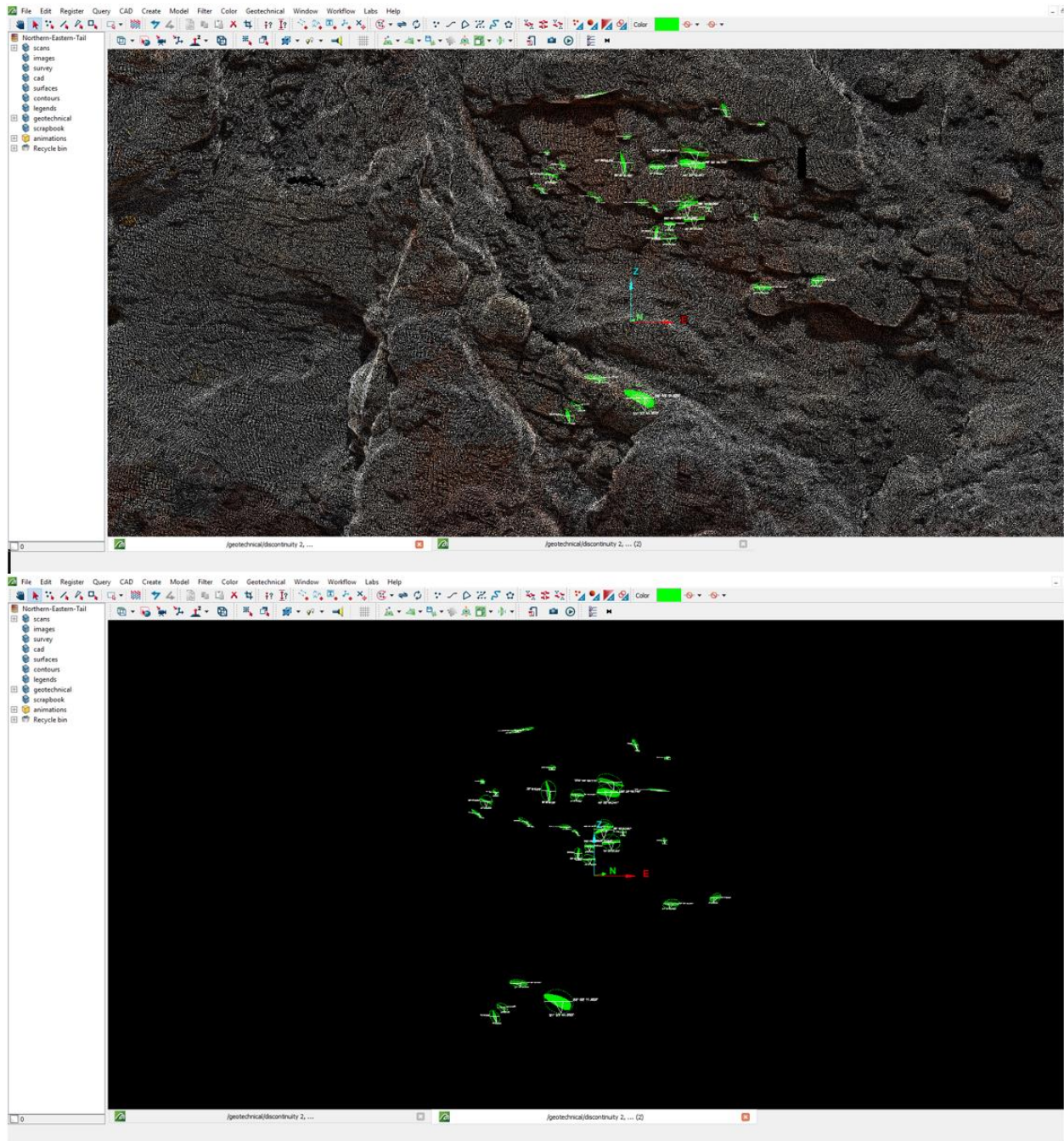


Figure 21: Maptek surface (Upper figure shows **Error! Reference source not found.**'s original figure while lower figure shows the fracture tracing map) (see Appendix 4 for the data)

3.2.4. *Fracture systems*

Figure 22-Figure 23 show stereogram plots of bedding orientations together with fracture orientations for each of the six focus sites. Figure 22 show measurements in modern coordinates while Figure 23 show rotated measurements with bedding restored to horizontal. The purpose of the rotation was to determine if the data were consistent with the fractures forming early in the diapiric history or late. That is, if the fractures formed when the beds were subhorizontal, the orientations should be consistent among the sites when beds are restored to horizontal whereas if they formed late, they should be similar in modern, geographic coordinates. Note that this analysis makes two major assumptions: 1) that all fractures formed at the same time; and 2) that bed rotations occurred along the strike line. Although neither of these assumptions can be fully justified, in the absence of other data, these are the only end members that can be easily evaluated.

Comparing Figure 22 and Figure 23, it is not clear if either end-member assumption is correct. The three inner sites (domain 0) all show the same general pattern of two orthogonal fracture sets approximately perpendicular to bedding, and in modern coordinates these sites show grossly different patterns (Figure 22) but are relatively similar when bedding is restored to horizontal (Figure 23). This suggests that for these sites the fractures either formed early, when beds were near horizontal, or some complex superposition which coincidentally produces this result. In contrast, outer zone sites in the more intact Cutler, are more ambiguous (Figure 22 and Figure 23). That is, in modern coordinates, these sites have similar fracture orientations, but because they also have similar bedding orientations, this pattern remains in restored coordinates. Thus, for these sites, orientation alone is ambiguous on the timing of fracture formation.

One solution to this discrepancy may arise when we incorporate the results analyzing fractures on the 3D model, particularly the *Cutler-Second Farthest* site where this broader scale analysis on the model revealed a second fracture set at ~60 degrees to the main fracture set. This geometry suggests a conjugate shear fracture set, which is consistent with small offsets observed on some fractures in the field (Figure 24). Under that interpretation, the observation that the fractures are generally perpendicular to bedding implies they probably formed early, when the beds were subhorizontal. In addition, if the fractures are a conjugate set in the outer, Cutler exposures, they imply a subhorizontal, bed-parallel maximum stress oriented ~NS to NNE in restored bedding coordinates.

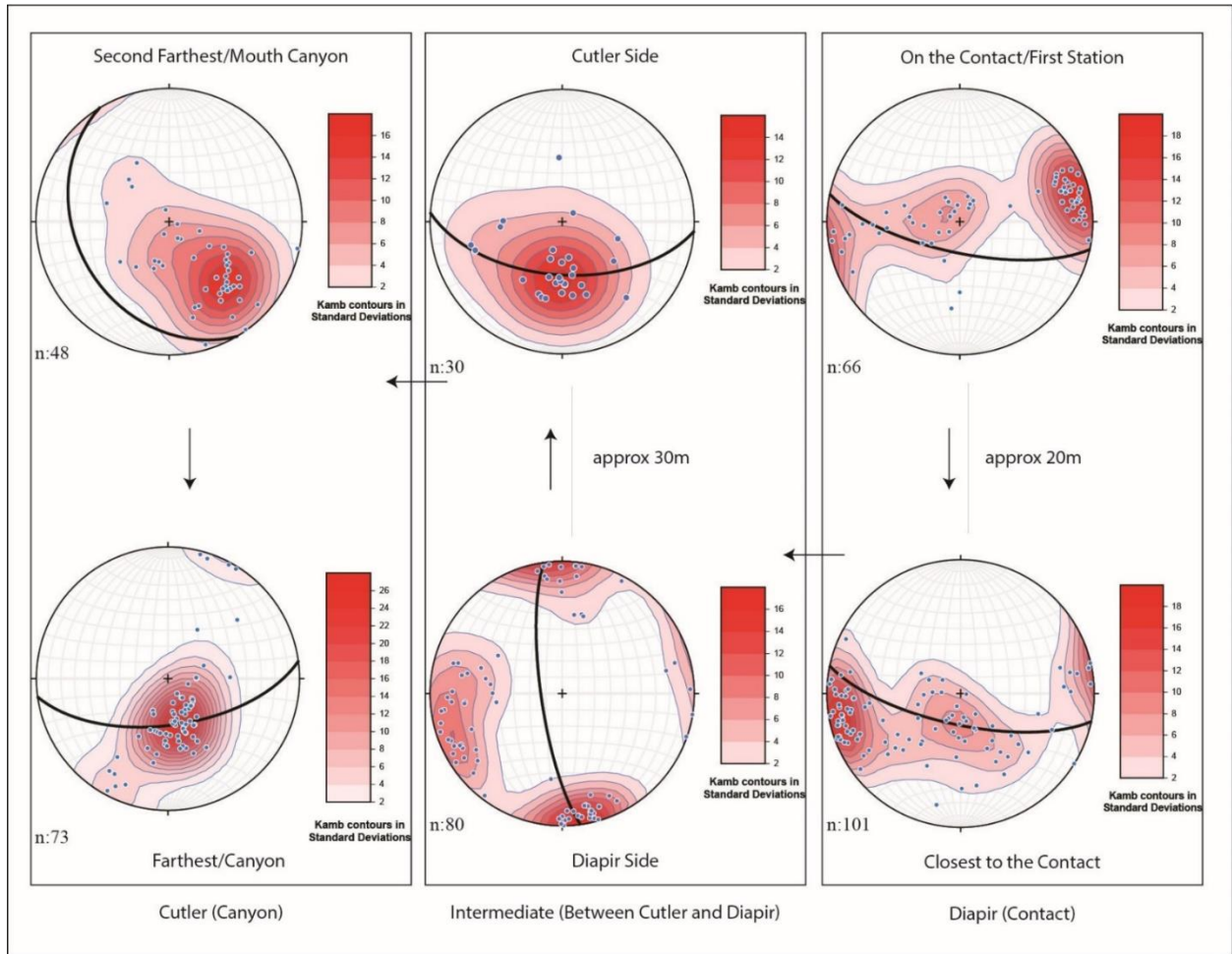


Figure 22: Kamb contoured equal area stereonet plot of poles to fractures with bedding at each site shown as heavy great circle. (The arrows between stereoplots show distance from diapir contact toward the canyon) (see Appendix 1 for the data)

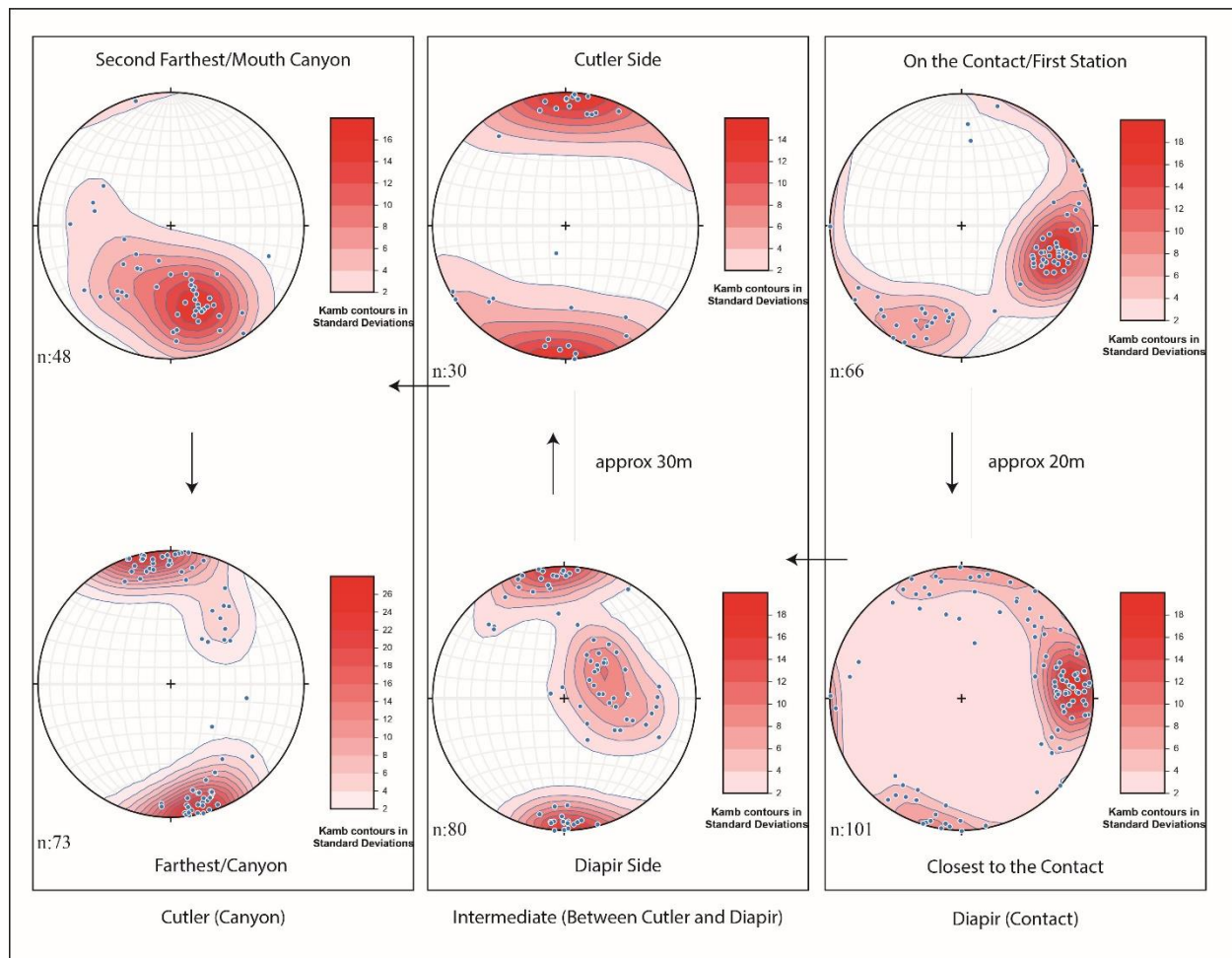


Figure 23: Kamb contours after bedding rotation to horizontal (The arrows between stereoplots show distance from diapir contact toward the canyon) (see Appendix 1 for the data)



Figure 24: Offset around the Cutler-Second Farthest site

4. Discussion

This study was undertaken with a goal of understanding fracture systems surrounding the Onion Creek diapir with a focus on a well exposed deformed zone at the diapir margin. The approach of using a field study combined with new 3D modeling capabilities led to unexpected results with implications for approaches in the use of these field methods. Thus, I discuss the implications for use the technology in this type of study as well as the original intent of the study, the evaluation of the fracture system and its relation to the formation of the Onion Creek diapir and significance for risking hydrocarbon traps adjacent to diapirs.

4.1. Mapping of fracture systems on virtual outcrop models

The 3D models allow spectacular display of the 3D geometry of the study outcrop. Manipulating these models in 3D allowed assessment of structural geometry that would not be possible in the field. For this study, one application of the model was that focus sites, together with adjacent parts of the outcrops could be revisited at will within the visualization of the model. This allowed mapping of surfaces including bedding and fractures by digitizing in 3D, directly on the model as well as acquire measurements for bedding, faults and/or fractures in the models.

An example of an application of using the model to extend field observations is the analysis shown in Figure 16. Here, limitations of access to the outcrop suggested that only one generation of fractures was present, based on field measurements. However, by zooming out to larger areas on the model, and carefully assessing the larger outcrop, it became apparent that a second fracture set was present, displayed particularly prominently on the upper part of the outcrop where it could not be easily reached for field measurements. In addition, the model

allowed routine mapping of bedding and fractures within this exposure, further clarifying the geometry (Figure 16).

4.2. Use of 3D outcrop models for remote sensing of fracture orientations

In addition to allowing detailed mapping of fracture traces, the 3D model also afforded an opportunity to dramatically increase the density of strike and dip data relative to data collected in the field, provided the method gave comparable results. To evaluate that question, I measured fractures on the model from the same focus sites that were examined in the field (Figure 30). The model aided in this analysis by allowing analysis across a broader region, including not only the focus site but adjacent areas as well. The result of this analysis was different at different sites with some sites showing reasonable correlations between the field data and the model measurements (e.g. Figure 13 and Figure 15), one site that faithfully reproduced the field result but also revealed a second fracture set (Figure 16), and one site that gave completely unacceptable results (Figure 12). Each of these cases implies different explanations.

In the case of the unacceptable comparison (Figure 12) my hypothesis is that remote analysis failed because of a unique feature of the outcrop geometry and lighting conditions at the time the images were acquired. Specifically, the fracture apertures are small relative to the model resolution (e.g. note the point spacing in Figure 12). Thus, small error in position of points in the point cloud can produce a significant error when a multipoint analysis is used to determine strike and dip. In addition, most of the fractures intersect a relatively planar surface, bedding, with deep shadow obscuring internal surfaces of the fracture. Thus, when attempting to measure the surfaces with remote analysis, there is insufficient 3D control to accurately measure the fractures. This is clear in Figure 12 where most of the measured surfaces are at low angles for the

model surface, yet the fractures that were being measured are at a high angle to the surface. Some of these surfaces are real fractures of other orientations but they then stand out in the analysis because most of the fractures in the dominant set remain unresolvable in the method.

In contrast, other sites, where the fractures are prominent at outcrop scale and are easily seen on the model (Figure 13, Figure 15 and Figure 16) the results of the measurements on the model are comparable to the field measurements. In fact, it could be argued here that the results of measurements on the model are actually superior to the field measurements (Figure 25 and Figure 27). From six focus site, although the *On the Contact* (Figure 12) has tiny fractures and two unsuccessful results, *Intermediate-Diapir Side* focus site (Figure 14) did not appear on the model, and *Cutler-Farthest* station had georeferenced systematic error that I could not figure out (Figure 17), the rest gave spectacular results (Figure 25 and Figure 27). Therefore, it is simply possible that the model measurements are actually better because it works well in larger areas.

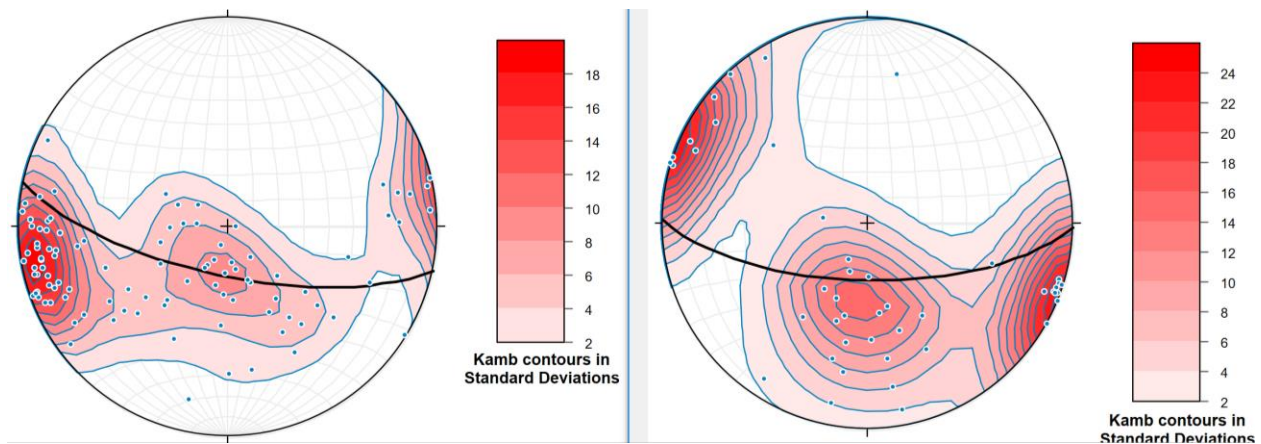


Figure 25: Closest to the Contact focus site. Stereoplot from field data (Left), from the model (Right)

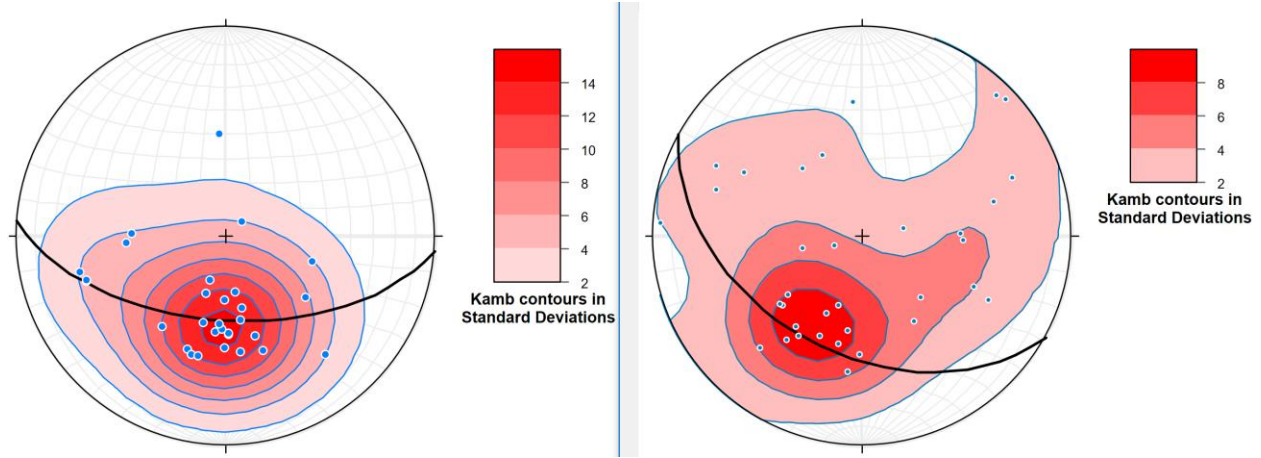


Figure 26: Intermediate-Cutler Side focus site. Stereoplot from field data (Left), from the model (Right)

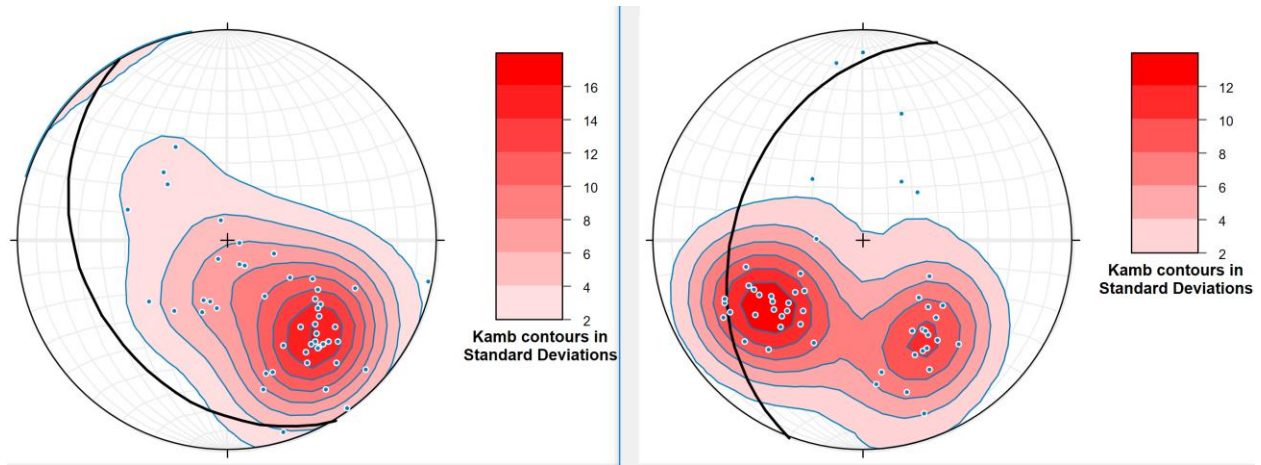


Figure 27: Cutler-Second Farthest focus site. Stereoplot from field data (Left), from the model (Right)

In any case, using 3D models remotely saves time and reduces expenses of fieldwork. In addition, it allows whole outcrop analysis. For example, in this study it would have been possible to map every fracture visible down to cm levels throughout the outcrop, extrapolating orientations from these data as well as connectivity. Ideally the could even be extended to 3D by constructing 3D surface from the model elements, although assumptions in surface construction would likely outweigh benefits of this kind of analysis. Nonetheless, future studies should consider this type of approach.

Although promising, the result here indicates caution is needed in using these data. The failure of the method in the site near the contact of the diapir is discouraging, yet not surprising in retrospect when the measurements are seen in the context of the sparsity of the point cloud where the measurements were made (Figure 12). One solution to this problem is apparent from Figure 12. Had we taken more detailed photos, at this specific site, and made a very high-resolution virtual outcrop at the scale of 1-2m, it is likely the smaller fractures could have been resolved. Thus, future studies should consider multiscale approaches; using these methods to make outcrop models of the entire site, as done here, with focused, more detailed models at key sites.

4.3. Fracture patterns near the Onion Creek diapir

It is difficult to generalize to larger scales from a small exposure like the outcrop examined here, yet the fracture systems observed are consistent with general models of stresses surrounding salt diapirs. Salt intrusion causes localized high stresses and leads to initiation and propagation of fractures near diapirs (NRC, 1996). Quinta et al. (2012), for example, states that the influence of diapir-related stress decreases away from diapir margin, which is consistent with the observed limited range of fault systems and bed rotations near diapirs. The data sets analyzed here indicate fractures are more abundant near the diapir, including rocks that are basically cataclasites within 20m of the diapir contact. In these close in sites there has been intense deformation, but it is difficult to determine how it relates to diapirism when bedding changes abruptly across numerous faults. Nonetheless, it is important to note that in areas near the contact where bedding is recognizable, the faults form a mutually orthogonal set which is itself perpendicular to bedding. This suggests these fractures are extension fractures and record a stretching within the bedding plane or two superimposed orthogonal extension events.

Farther from the diapir contact, however, the fracture systems are distinctly simpler in the study site shear fractures (Figure 24). Salt diapirs typically show radial faults, concentric faults, or both (Davison et al., 2000; Stewart, 2006) (Figure 28 and Figure 29). Figure 22Figure 23 show that in this study area both radial and concentric fractures are present at the six field sites I examined but vary with position relative to the diapir margin. In particular, in modern coordinates two fracture sets are present at sites near the diapir margin (domain 0) but only one is present farther from the diapir. In all cases, however, the fractures are approximately perpendicular to bedding. In modern, coordinates this pattern is indistinct, but when the fractures are restored with bedding horizontal, a clearer picture emerges (Figure 23). That is, in the sites close to the diapir, the two fracture sets include both a radial and concentric fracture set whereas farther from the diapir, the radial set disappears and only the concentric set is present. This implies that well away from the diapir margin radial extension was dominant whereas near the diapir margin the rocks were stretched in two directions.

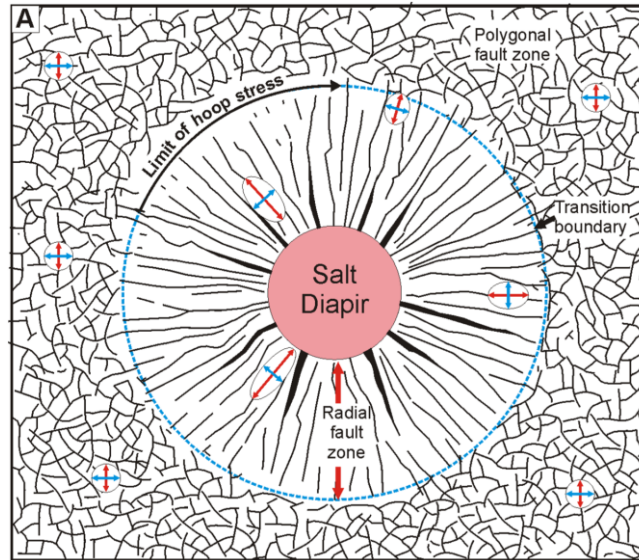


Figure 28: Polygonal fault zone demonstrates the geometry of a hoop stress field around a salt diapir. The blue dashed circle shows the limit of the hoop stress. The red and blue arrows respectively illustrate the ratio of the maximum and minimum horizontal principal stresses (Carruthers, 2012).

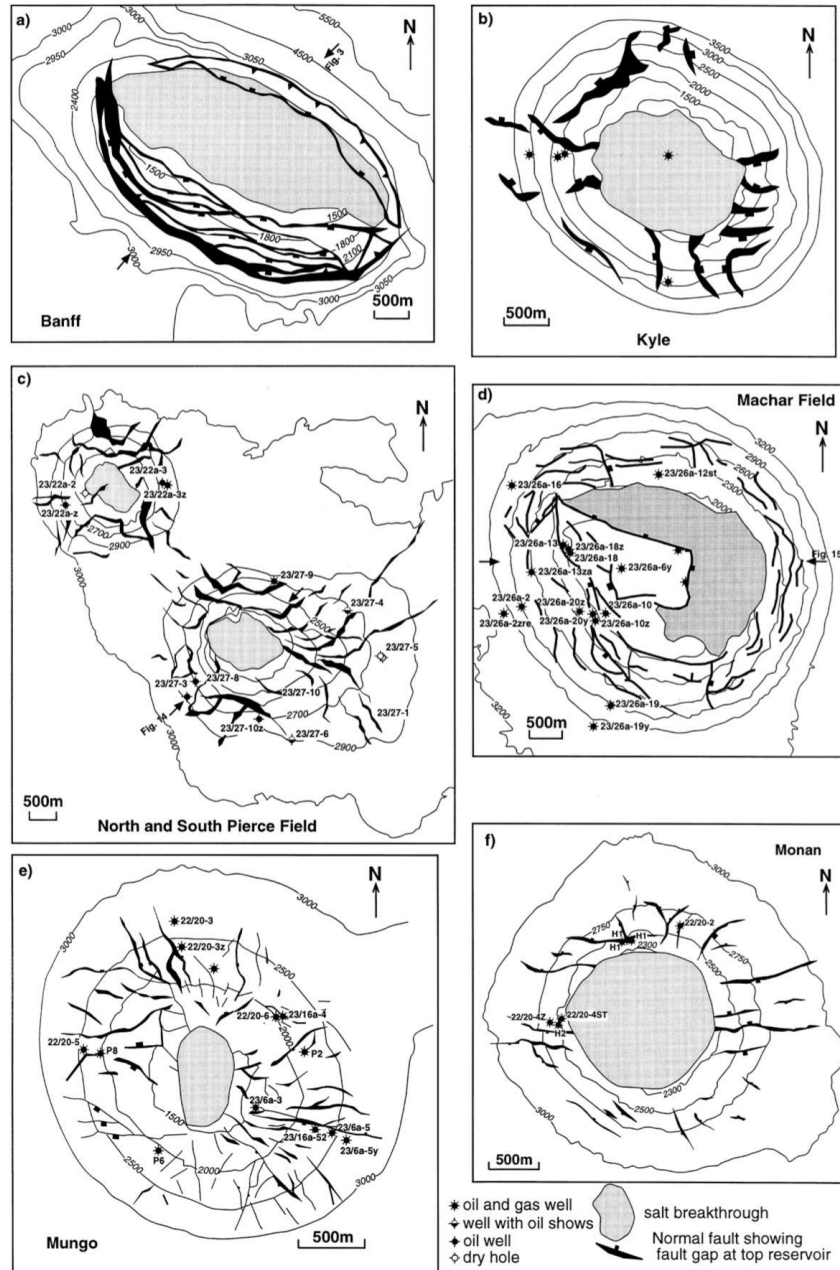


Figure 29: The figure shows some examples of radial and concentric faults around salt diapirs. Interpreted by oilcompanies such as Conoco, Ranger Oil, Enterprise Oil, and BP Amoco (Davison, 2000).

4.4. Use of a Smart Phone for collecting orientation data

In the last few years the availability of software on smartphones that allow use of these devices as a replacement for a conventional geologist's compass has led to their widespread use,

yet evidence has accumulated that not all devices yield accurate results (Novakova and Pavlis, 2017; Allmendinger et al, 2017). However, the results of my device were positive (Figure 30). Specifically, in a field test of similar fracture sets in the same outcrop (Figure 30) the phone and Brunton compass produced comparable results. The device used is an iPhone 7 Plus (IOS 11) with Field Move Clino 2.3.3 and thus, this result is consistent with the conclusions of Allmendinger et al. (2017) that iPhones, in general, produce good results in these types of analyses.

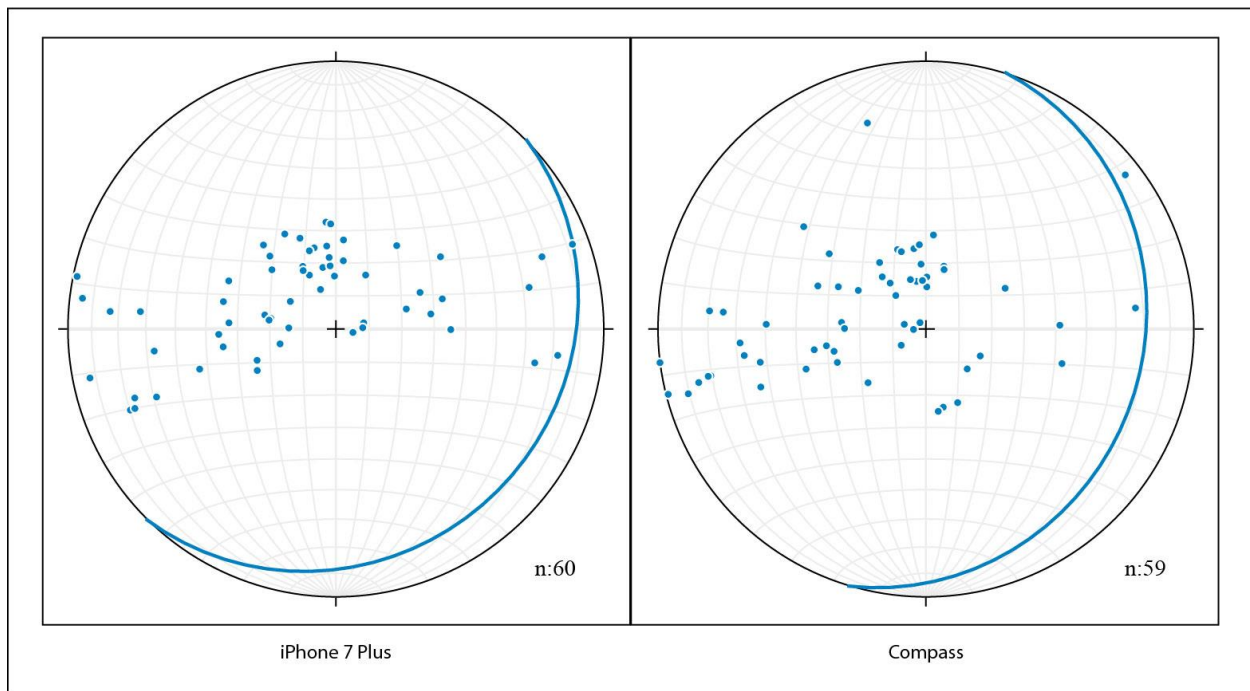


Figure 30: The comparison of phone data and compass data (the declination of the compass was set up for El Paso which is 8.1° although the phone was set up for Utah that is 10.1°) (Blue circles show the overall fracture orientation while dots show fracture) (see Appendix 3 for the data)

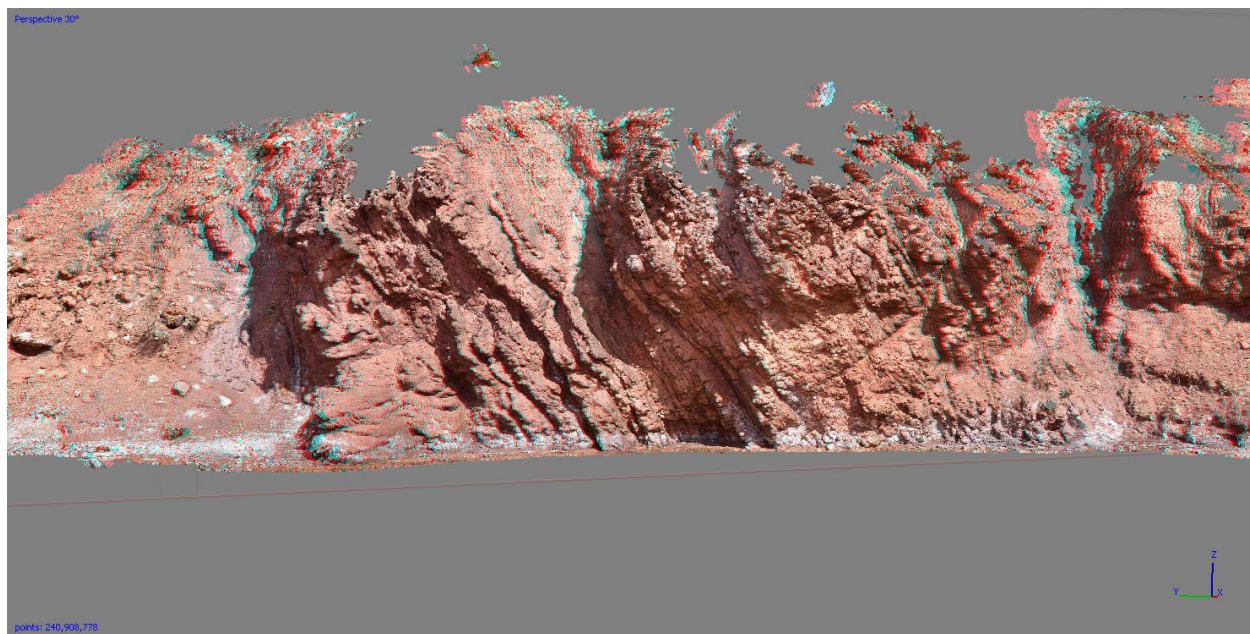


Figure 31: Cyan-magenta stereo view of the data station of Cutler side in the intermediate section of the walls from Photoscan stereo view (Y shows north, X shows east)

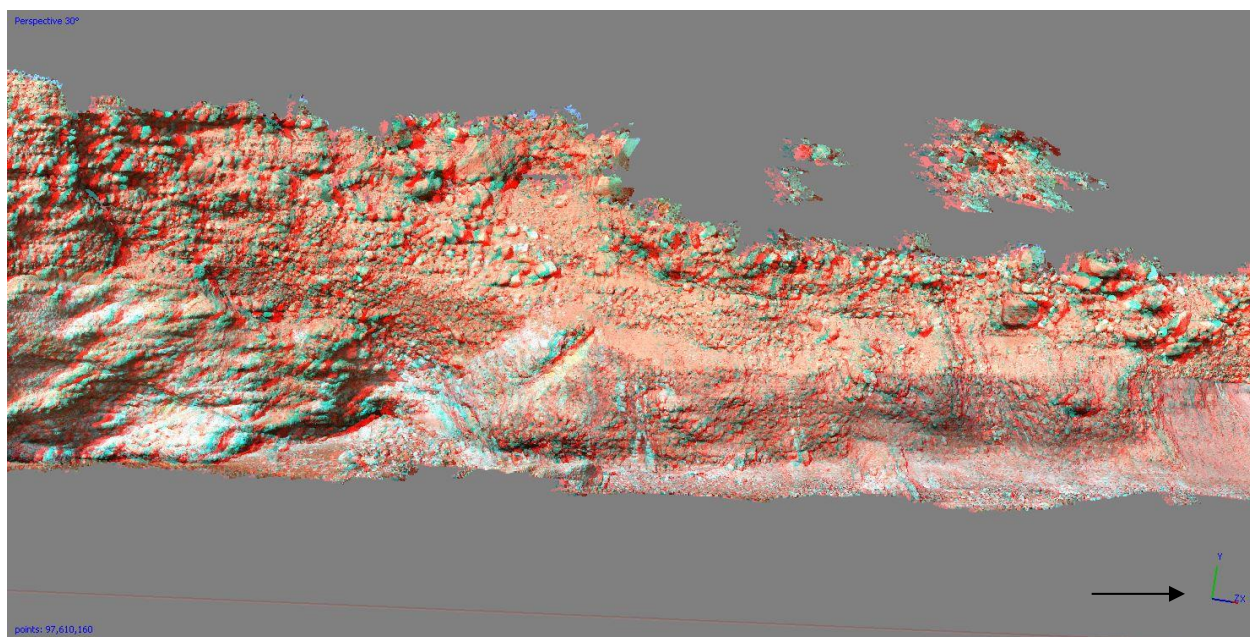


Figure 32: Southern west section of the walls in cyan-magenta Photoscan stereo view (Black arrow shows north)

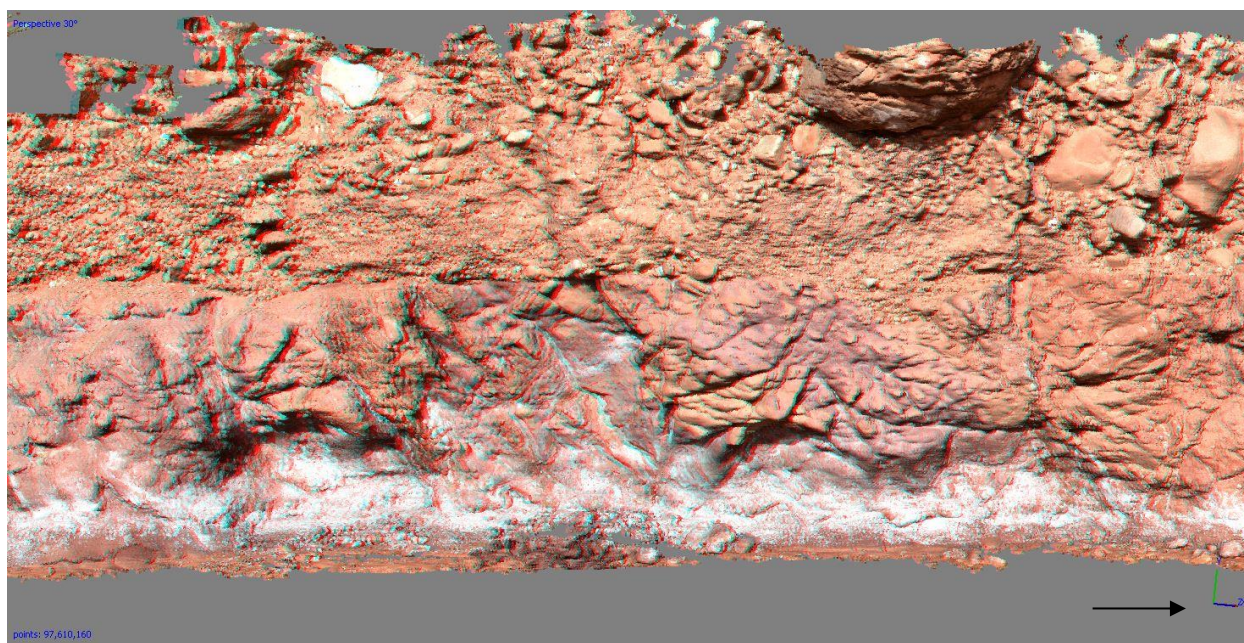


Figure 33: Northern west section of the walls in cyan-magenta Photoscan stereo view (Black arrow shows north)

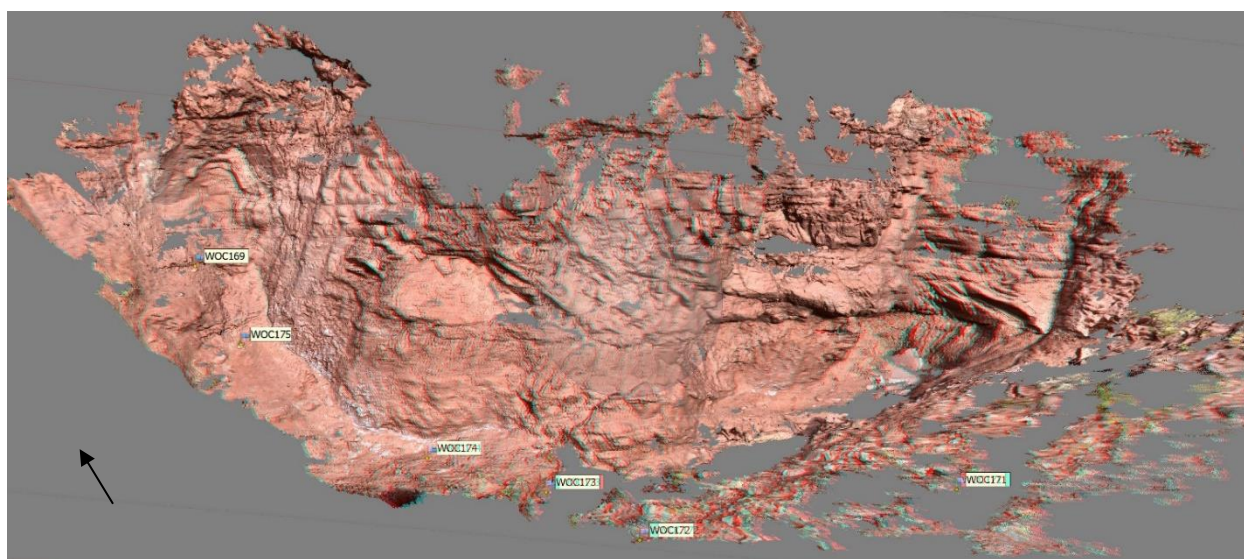


Figure 34: Canyon in cyan-magenta Photoscan stereo view (Black arrow shows north)

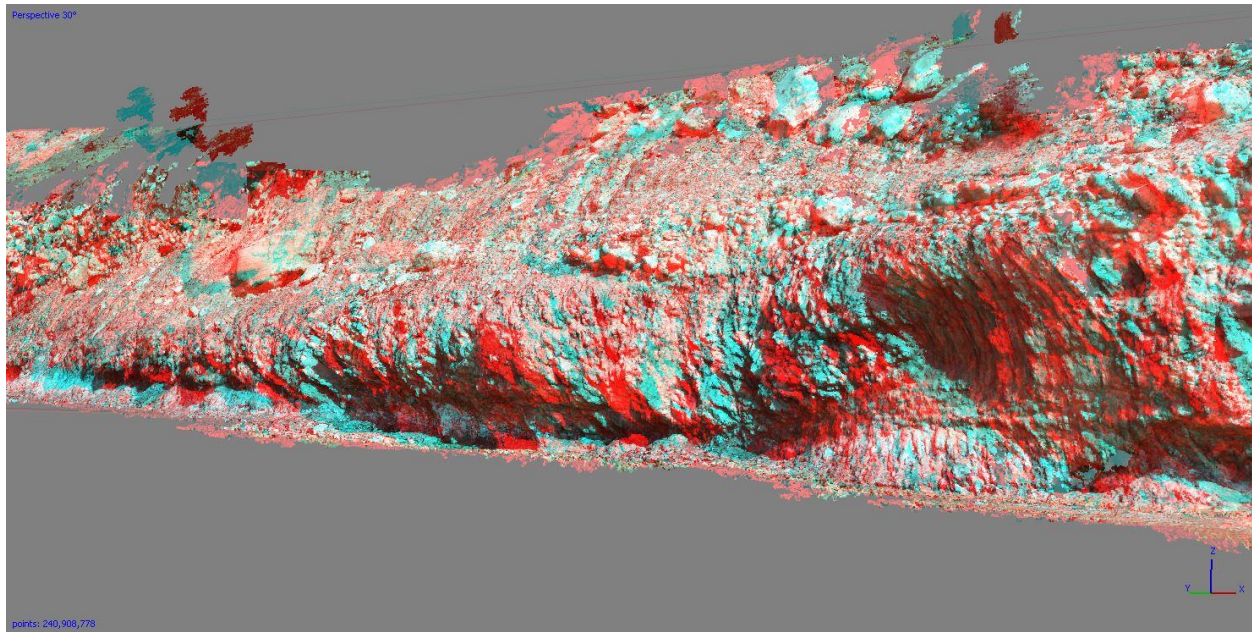


Figure 35: Southern east side of the walls in cyan-magenta Photoscan stereo view (Y shows north, X shows east)

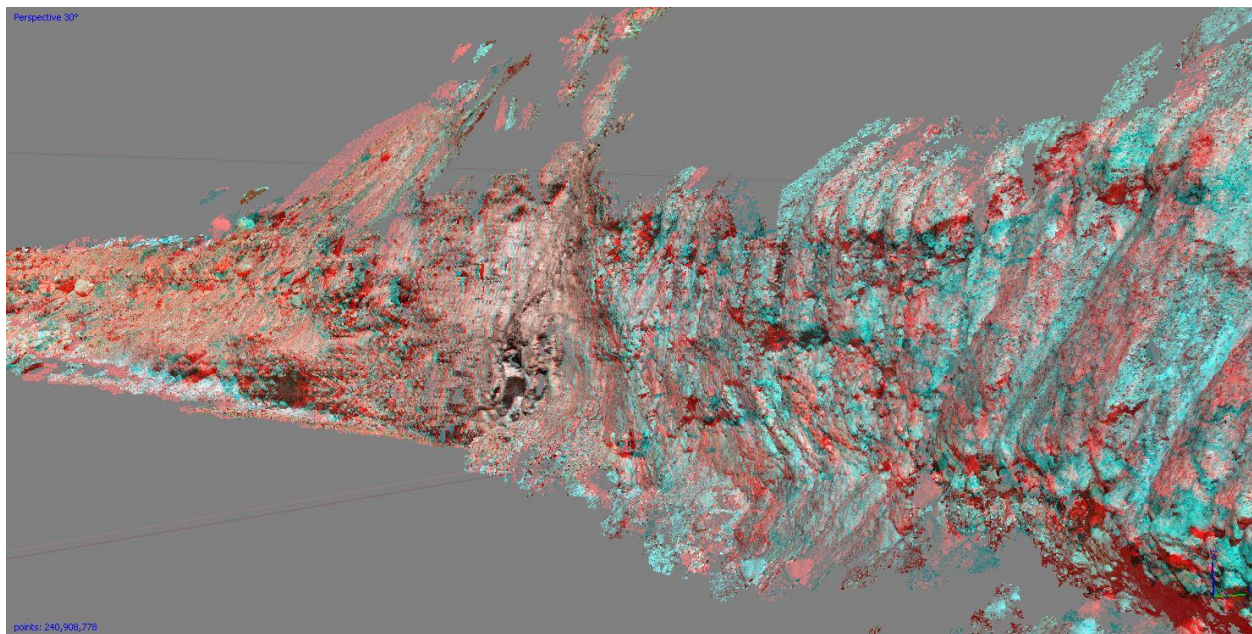


Figure 36: The data stations of the diapir section of the wall in cyan-magenta Photoscan stereo view (Y shows north, X shows east)

5. Conclusion

Field exposures in the Paradox Basin provide unprecedented detail of the fracture systems and stratigraphic relationships that developed along the margins of salt diapirs, providing potential to understand fracture permeability development in other salt provinces. This study used a new method, 3D terrain models using SfM/MvS, as a base for detailed mapping and fracture analysis to the mm level. Use of these methods produced new insights with important cautionary notes.

Significant results of this study include:

- 1) Collecting ground control data with only a handheld GPS is -generally not reliable if submeter accuracy is needed. Although horizontal errors were in line with known errors in WAAS corrected differential GPS, a systematic error of more than 8m was seen in the vertical. The origin of this error is unknown, but likely related to weather conditions or local geography disturbing the differential correction in the WAAS signal.
- 2) Measuring orientations with a cell phone provides reasonable data, provided the cell phone passes simple tests.
- 3) Remotely measured fracture orientations using multipoint analyses on the model produced mixed results. Where fractures were large and easily observed on the model, the model-based measurements were broadly consistent with the field measurements. Indeed, in one site the model-based measurements were arguably superior because they resolved a second fracture set that was clear on unreachable cliff faces above the field site, but difficult to see at the level of a field observer. In contrast, at one site near the diapir contact the field measurements and model-based

- measurements were completely at odds. In this case, however, the origin of the discrepancy is clear from 3D assessment of the model where relatively 2D outcrop geometry and small scale of the fractures prevented routine analysis of 3D geometry.
- 4) In cases like the site where model measurements failed a solution to the problem is either use standard field measurement techniques for the analysis or to use a camera to second set of very high-resolution photographs of the site of interest, allowing development of a more high resolution model where the fracture systems would be resolved. Which method is used would depend on field time vs lab time.
 - 5) Although the 3D model allowed rapid mapping of fracture traces and bedding traces to show 3D trajectories, fracture connectivity and fracture network assessments proved difficult in this study site because although fractures and faults could be traced through the cliff faces of the model, a large flat area between the cliffs is covered, handicapping tracing feature between the two complete exposures.

Suggestion for Future work

The salt diapir at Onion Creek needs more fracture sites around it better understand the relationships between faults, fracturing and bedding. Similar techniques could be used elsewhere along the diapir margin from domain 0 sites into intact Cutler to further clarify the origin of the fracture systems. Particularly important is a more regional assessment of the fracture systems, beyond the limited range of this small outcrop.

References

- Allmendinger, W. A., Siron, C. R., & Scott, C. P. (2017) Structural data collection with mobile devices: Accuracy, redundancy, and best practices. *Journal of Structural Geology*. 102, 98-112.
- Baars, D.L. (1966) Pre-Pennsylvanian paleotectonics - key to basin evolution and petroleum occurrences in Paradox basin, Utah and Colorado. *Am. Assoc. Petrol. Geol. Bull.*, 50, 2082-2111.
- Baars, D.L. & Stevenson, G.M. (1981) Tectonic evolution of the Paradox basin, Utah and Colorado. In: *Geology of the Paradox Basin* (Ed. by D.L. Wiegand), pp. 23-31. Association of Geologists, Rocky Mount.
- Baars, D. L., & Stevenson, G. M., 1982, Subtle stratigraphic traps in Paleozoic rocks of Paradox Basin; in *The Deliberate Search for the Subtle Trap*, M. T. Halbouty, ed.: American Association of Petroleum Geologists, Memoir 32, p. 131-158.
- Banbury, N.J. (2005) The role of salt mobility in the development of supra-salt sedimentary depocentres and structural styles. PhD Thesis, University of Edinburgh, Edinburgh, UK.
- Barbeau, D.L. (2003) A flexural model for the Paradox basin: implications for the tectonics of the Ancestral Rocky Mountains. *Basin Res.*, 15, 97-115.
- Bromley, M.H. (1991) Architectural features of the Kayenta Formation (Lower Jurassic), Colorado Plateau, USA: relationship to salt tectonics in the Paradox Basin. *Sediment. Geol.*, 73, 77-99.
- Carrivick, J.L., Smith, M.W., Quincey, D.J., 2016, *Structure from Motion in the Geosciences*, John Wiley and Sons, West Sussex, UK, 197p.

- Carruthers, T. D. (2012) Interaction of Polygonal Fault Systems with Salt Diapirs. PhD Thesis, Cardiff University, Cardiff, UK.
- Condon, S.M. (1997) Geology of the Pennsylvanian and Permian Cutler Group and Permian Kaibab Limestone in the Paradox Basin, Southeastern Utah and Southwestern Colorado. *US Department of the Interior, US Geological Survey Bulletin*, 2000-P, 46pp.
- Cook, E. L. (2017) Near Salt Stratal Geometries and Implications for the Evolution of the Onion Creek Diapir Moab, UT. MS Thesis, Brigham Young University, UT.
- Davison, I., Alsop, I., Birch, P., Elders, C., Evans, N., Nicholson, H., Rorison, P., Wade, D., Woodward, J., Young, M., 2000. Geometry and late-stage structural evolution of Central Graben salt diapirs, North Sea. *Marine and Petroleum Geology* 17(4), 499-522.
- Dershowitz, W. S., and P. R. LaPointe, 1994, Discrete fracture approaches for oil and gas applications, in R. A. Nelson and S. Laubach, eds., *Rock mechanics*: Rotterdam, Balkema, p. 19– 30.
- Doelling, H. H. (1985) Geologic map of Arches National Park and vicinity, Grand County, Utah. Utah Geological and Mineral Survey Map 74. Utah Geological Survey.
- Doelling, H. H. (1988) Geology of Salt Valley anticline and Arches National Park, Grand County, Utah. In: *Salt Deformation in the Paradox Region* (Ed. By H.H. Deolling, C.G. Oviatt & P.W. Huntoon), *Utah Geol. Survey Bull.*, 122, 1-60.
- Doelling, H.H. (2001) Geologic map of the Moab and eastern part of the San Rafael Desert 30'x60' Quadrangles, Grand and Emery Counties, Utah and Mesa County, Colorado. Utah Geological Survey Map 180.
- Doelling, H.H. (2002) Geologic map of the Fisher Towers 7.5' quadrangle, Grand County, Utah. Utah Geological Survey Map 183, p.22.

- Elston, D.P. & Shoemaker, E.M. (1960) Late Paleozoic and early Mesozoic structural history of the Uncompahgre front. In: *Geology of the Paradox basin fold and fault belt - Four Corners Geol. Soc. 3rd Field Conference* (Ed. by K.G. Smith), pp. 47-55 Four Corners Geological Society, Durango, Colorado.
- Elston, D.P., Shoemaker, E.M. & Landis, E.R. (1962) Uncompahgre front and salt anticline region of Paradox Basin, Colorado and Utah. *Am. Assoc. Petrol. Geol. Bull.*, 46, 1857-1878.
- Elston, D.P. & Shoemaker, E.M. (1963) Salt anticlines of the Paradox Basin, Colorado and Utah. In: *Symposium on Salt - Northern Ohio Geological Society* (Ed. by A.C. Bersticker), pp. 131-146. Northern Ohio Geological Society, Cleveland, Ohio.
- Friedman, J.D., Case, J.E. & Simpson, S.L. (1994) Tectonic trends of the northern part of the Paradox Basin, southeastern Utah and southwestern Colorado, as derived from landsat multispectral scanner imaging and geophysical and geologic mapping. *US Department of the Interior, US Geological Survey Bulletin*, 2000-C, p. 34.
- Green M. (2014) Own work, CC BY-SA 3.0, <https://commons.wikimedia.org/w/index.php?curid=33572534>
- Goldhammer, R.K., Oswald, E.J. & Dunn, P.A. (1994) High frequency, glacio-eustatic cyclicity in the Middle Pennsylvanian of the Paradox Basin-an evaluation of Milankovitch forcing. In: *Orbital Forcing and Cyclic Sequences. Special publication of International Association of Sedimentologists*, 19 (Ed. by P.L. de Boer & D.G. Smith), 243-283. Blackwell, Oxford.
- Grisi, K. C. (2018) Attributes of Fisher Valley Megaflap and Comparison to the Gypsum Valley Megaflap, Paradox Basin, Utah and Colorado: Implications for Controls on Megaflap Formation. MS Thesis, The University of Texas at El Paso, TX.

- Hanke, J.R. (2017) *Journal of Structural Geology*. <https://doi.org/10.1016/j.jsg.2017.11.012>
- Hazel, J.E. (1994) Sedimentary response to intrabasinal salt tectonism in the Upper Triassic Chinle Formation, Paradox basin, Utah. US Department of the Interior, US Geological Survey Bulletin, 2000-F, 34pp.
- Hite, R.J. & Buckner, D.H. (1981) Stratigraphic correlations, facies concepts, and cyclicity in Pennsylvanian rocks of the Paradox basin. In: *Geology of the Paradox Basin* (Ed. By D.L. Wiegand), pp.147-159. Association of Geologists, Rocky Mount.
- Hite, R.J., Anders, D.E., and Ging, T.G., 1984, Organic-rich source rocks of Pennsylvanian age in the Paradox Basin of Utah and Colorado: in *Hydrocarbon Source Rocks of the Greater Rocky Mountain Region*, Rocky Mountain Association of Geologists, p. 255-271.
- Hudec, M.R. (1995) The Onion Creek salt diapir: an exposed diapir fall structure in the Paradox basin, Utah. In: *Salt, Sediment and Hydrocarbons. Gulf Coast Section Society of Economic Paleontologists and Mineralogists 16th Annual Research Conference* (Ed. by C.J. Travis, H. Harrison, M.R. Hudec, B.C. Vendeville, F.J. Peel & B.F. Perkins), pp. 125-134. Gulf Coast Section SEPM Foundation, Houston, Texas.
- Kelley, V.C., 1958, Tectonics of the region of the Paradox Basin, in Sanborn, A.F., ed., *Guidebook to the Geology of the Paradox Basin*: Intermountain Association of Petroleum Geologists, p. 31–38.
- Kenroy, P. R. (2013) Fracture-Controlled Paleohydrologic Systems in the Vicinity of Salt Diapirs. MS Thesis, Northern Illinois University.
- Kluth, C.F. & Duchene, H.R. (2009) Late Pennsylvanian and early Permian structural geology and tectonic history of the Paradox Basin and Uncompahgre uplift, Colorado and Utah. In: *The Paradox Basin Revisited - New Developments in Petroleum Systems and Basin*

- Analysis* (Ed. by W.S. Houston, L.L. Wray & P.G. Moreland), pp. 178-197. Association of Geologists, Rocky Mount.
- Lawton, T.F. & Buck, B.J. (2006) Implications of diapir-derived detritus and gypsic paleosols in Lower Triassic strata near the Castle Valley salt wall, Paradox Basin, Utah. *Geology*, 34, 885-888.
- Matthews, N.A., Noble, T.A., and Breithaupt, B.H., 2016, Close-range photogrammetry for 3D ichnology: The basics of photogrammetric ichnology, in Falkingham, P.L., Marty, D., and Richter, A, eds., *Dinosaur Tracks – The Next Steps*: Bloomington, Indiana, Indiana University Press, 520 p.
- Moore, K.D., Soreghan, G.S. & Sweet, D.E. (2008) Stratigraphic and structural relations in the proximal Cutler Formation of the Paradox Basin: implications for timing of movement on the Uncompahgre front. *Mountain Geol.*, 45, 49-68.
- Narr, W., Schechter D.S. & Thompson L.B. (2006) Naturally Fractured Reservoir Characterization. Society of Petroleum Engineers, 112.
- Nelson, R. A., 1985, *Geologic analysis of naturally fractured reservoirs*: Houston, Texas, Gulf Publishing, 320 p.
- Nicholas, I.P.W., Alves T.M., & Blenkinsop T.G. (2016) Reservoir leakage along concentric faults in the Southern North Sea: Implications for the deployment of CCS and EOR techniques. *Tectonophysics*. 690. 97-116.
- NRC (National Research Council) (1996) *Rock Fractures and Fluid Flow*. National Academy Press, Washington D.C. 551.

- Novakova, L. & Pavlis T. L. (2017) Assessment of the precision of smart phones and tablets for measurement of planar orientations: A case study. *Journal of Structural Geology*. 97, 93-103.
- Nuccio, V.F. & Condon, M. (1996) Burial and thermal history of the Paradox basin, Utah and Colorado, and petroleum potential of the Middle Pennsylvanian Paradox Formation, US Department of the Interior, US Geological Survey Bulletin, 2000-O, p. 41.
- Ohlen, H.R. & McIntyre, L.B. (1965) Stratigraphy and tectonic features of the Paradox Basin, four corners area. *Am. Assoc. Petrol. Geol. Bull.*, 49, 2020-2040.
- Online Reference 1, <https://www.nikonusa.com/en/nikon-products/product/dslr-cameras/d5300.html>
- Online Reference 2, http://kb.unavco.org/kb/assets/24/Gb1000_om.pdf
- Online Reference 3, www.geo.cornell.edu/geology/faculty/RWA/programs/stereonet.html
- Online Reference 4, https://www.faa.gov/about/office_org/headquarters_offices/ato/service_units/techops/navservices/gnss/waas/benefits/
- Online Reference 5, <http://www.esri.com/news/arcuser/0703/geoid1of3.html>
- Peacock, D.C.P., Sanderson, D.J. & Rotevatn A. (2018) Relationships between fractures. *Journal of Structural Geology*. 106, 41-53.
- Quinta, A., Tavani, S. & Roca E. (2012) Fracture pattern analysis as a tool for constraining the interaction between regional and diapir-related stress fields: Poza de la Sal Diapir (Basque Pyrenees Spain). *Salt Tectonics, Sediments and Prospectivity*. Geological Society of London, Special Publications. 363, 521-532.

- Rasmussen, D.L. (2006) Escalation of Pennsylvanian-Permian Uncompahgre uplift revealed by progression of siliciclastic wedges in the Paradox basin, Colorado and Utah. In: *Geological Society of America 38th Annual Meeting - Abstracts with Programs Geol. Soc. Am.*, p. 29.
- Rasmussen, L. & Rasmussen, D.L. (2009) Burial history analysis of the Pennsylvanian petroleum system in the deep Paradox Basin fold and fault belt, Colorado and Utah. In: *The Paradox Basin Revisited - New Developments in Petroleum Systems and Basin Analysis* (Ed. By W.S. Houston, L.L. Wray & P.G. Moreland), pp. 24-94. Association of Geologists, Rocky Mount.
- Raup, O.B. & Hite, R.J. (1992) Lithology of evaporite cycles and cycle boundaries in the upper part of the Paradox Formation of the Hermosa Group of Pennsylvanian age in the Paradox Basin, Utah and Colorado. US Department of the Interior, US Geological Survey Bulletin, 2000-B, p. 37.
- Reiss, L. H., 1982, The reservoir engineering aspects of fractured formations: Houston, Texas, Gulf Publishing, 186 p.
- Rohrbaugh Jr., M. B., Dunne, W. M., & Mauldon, M. (2002). Estimating fracture trace intensity, density, and mean length using circular scan lines and windows. *AAPG Bulletin*. v. 86, no. 12, p. 2089–2104.
- Stewart, S.A., 2006. Implications of passive salt diapir kinematics for reservoir segmentation by radial and concentric faults. *Marine and Petroleum Geology* 23(8), 843-853.
- Trudgill, B.D. & Paz, M. (2009) Restoration of mountain front and salt structures in the northern Paradox Basin, SE Utah. In: *The Paradox Basin Revisited - New Developments in Petroleum Systems and Basin Analysis* (Ed. By W.S. Houston, L.L. Wray & P.G. Moreland), pp. 132-177. Association of Geologists, Rocky Mount.

- Trudgill B.D., 2011, Evolution of salt structures in the northern Paradox Basin: controls on evaporite deposition, salt wall growth and supra-salt stratigraphic architecture. *Basin Research*, **23**, 208-238.
- Venus, J. H., Mountney, N. P., and McCaffrey, W.D., 2015, Syn-sedimentary salt diapirism as a control on fluvial-system evolution: an example from the proximal Permian Cutler Group, SE Utah, USA. *Basin Research*, **27**, 152-182.
- Wengerd, S.A. (1962) Pennsylvanian sedimentation in Paradox Basin, Four Corners region. In: *Pennsylvanian system in the United States - A symposium* (Ed. by C.C. Branson), pp. 264-330. Am. Assoc. Petrol. Geol., Tulsa, Oklahoma.
- White, M.A. & Jacobson, M.I. (1983) Structures associated with the southwest margin of the ancestral Uncompahgre Uplift. In: *Northern Paradox Basin - Uncompahgre Uplift* (Ed. By W.R. Averett), pp. 33-39. Grand Junction Geological Society, Grand Junction, Colorado.

Appendix

Note: Right Arm (Hand) rule is applied for all the data collected.

Appendix 1

Location Name	Plane Type	Dip	Strike
On the Contact	Joint	86.93184	188.0206
On the Contact	Joint	78.43345	8.237022
On the Contact	Joint	9.942538	121.3026
On the Contact	Joint	74.6913	153.1973
On the Contact	Joint	46.5691	7.359779
On the Contact	Joint	74.72725	159.6113
On the Contact	Joint	81.03471	166.0405
On the Contact	Joint	22.23876	354.0491
On the Contact	Joint	17.58585	58.61084
On the Contact	Joint	79.88213	175.3539
On the Contact	Joint	74.75798	164.0471
On the Contact	Joint	52.78524	158.4556
On the Contact	Joint	76.27741	351.847
On the Contact	Joint	12.49043	118.5608
On the Contact	Joint	84.75092	355.7199
On the Contact	Joint	65.16581	158.2955
On the Contact	Joint	72.48669	353.5581
On the Contact	Joint	74.23563	172.4541
On the Contact	Joint	80.51147	169.8446
On the Contact	Joint	79.60308	168.7347
On the Contact	Joint	84.81478	157.8499
On the Contact	Joint	80.82881	345.9218
On the Contact	Joint	81.91376	173.9259
On the Contact	Joint	32.30918	162.1965
On the Contact	Joint	14.27886	59.86685
On the Contact	Joint	10.48571	80.75381
On the Contact	Joint	7.325335	61.34114
On the Contact	Joint	66.43555	156.9999
On the Contact	Joint	73.95701	169.2721
On the Contact	Joint	78.32726	178.8651
On the Contact	Joint	54.78507	275.8413
On the Contact	Joint	43.69757	270.4276
On the Contact	Joint	21.8984	323.5857
On the Contact	Joint	26.42608	331.0451

On the Contact	Joint	45.90946	355.4197
On the Contact	Joint	15.33203	121.2265
On the Contact	Joint	55.38374	358.2066
On the Contact	Joint	62.48687	358.8654
On the Contact	Joint	7.113765	28.24515
On the Contact	Joint	70.89864	164.3468
On the Contact	Joint	78.63577	169.5912
On the Contact	Joint	8.956587	319.066
On the Contact	Joint	29.20736	359.0963
On the Contact	Joint	31.43262	5.780457
On the Contact	Joint	77.86917	335.1301
On the Contact	Joint	13.58762	25.10345
On the Contact	Joint	16.33319	106.1944
On the Contact	Joint	80.95529	154.6776
On the Contact	Joint	36.0332	16.44467
On the Contact	Joint	51.52423	352.2583
On the Contact	Joint	66.83234	168.8799
On the Contact	Joint	73.8407	179.1206
On the Contact	Joint	82.81068	180.9504
On the Contact	Joint	69.58707	173.6232
On the Contact	Joint	82.55633	177.2373
On the Contact	Joint	25.61653	292.4061
On the Contact	Joint	14.48073	334.2122
On the Contact	Joint	77.61931	170.5388
On the Contact	Joint	71.12614	161.5626
On the Contact	Joint	78.17944	161.7873
On the Contact	Joint	82.74725	164.4059
On the Contact	Joint	69.49313	153.517
On the Contact	Joint	68.43951	154.971
On the Contact	Joint	73.63914	157.9221
On the Contact	Joint	67.89931	162.9387
On the Contact	Joint	86.92583	351.6616
Closest to the Contact	Joint	89	211.603
Closest to the Contact	Joint	74.38967	1.758194
Closest to the Contact	Joint	73.10008	11.20316
Closest to the Contact	Joint	83.43986	2.11496
Closest to the Contact	Joint	69.0992	357.7215
Closest to the Contact	Joint	68.77864	357.746
Closest to the Contact	Joint	17.33647	3.521133
Closest to the Contact	Joint	18.13592	272.7873
Closest to the Contact	Joint	78.12278	344.585
Closest to the Contact	Joint	16.83441	259.0238

Closest to the Contact	Joint	23.37774	248.0939
Closest to the Contact	Joint	79.36705	347.5238
Closest to the Contact	Joint	71.68233	351.8221
Closest to the Contact	Joint	83.19112	345.8605
Closest to the Contact	Joint	89	166.6301
Closest to the Contact	Joint	73.4901	282.4685
Closest to the Contact	Joint	22.22611	249.0123
Closest to the Contact	Joint	65.47832	165.2102
Closest to the Contact	Joint	65.94545	176.0228
Closest to the Contact	Joint	49.74899	194.5033
Closest to the Contact	Joint	51.47725	341.3108
Closest to the Contact	Joint	47.41932	221.1824
Closest to the Contact	Joint	40.20081	219.7929
Closest to the Contact	Joint	73.21349	336.3447
Closest to the Contact	Joint	39.00541	274.2021
Closest to the Contact	Joint	83.21065	347.6457
Closest to the Contact	Joint	86.77541	168.5401
Closest to the Contact	Joint	37.63688	309.3112
Closest to the Contact	Joint	50.22749	315.883
Closest to the Contact	Joint	26.74869	273.1642
Closest to the Contact	Joint	34.44733	236.178
Closest to the Contact	Joint	69.30959	328.5171
Closest to the Contact	Joint	53.20942	320.5961
Closest to the Contact	Joint	59.89247	320.5243
Closest to the Contact	Joint	85.77098	352.2072
Closest to the Contact	Joint	15.03163	234.0423
Closest to the Contact	Joint	80.35434	353.7656
Closest to the Contact	Joint	73.12293	2.447838
Closest to the Contact	Joint	86.50249	175.625
Closest to the Contact	Joint	27.30751	27.41409
Closest to the Contact	Joint	7.813634	288.2897
Closest to the Contact	Joint	21.22074	23.85633
Closest to the Contact	Joint	27.00196	346.241
Closest to the Contact	Joint	11.79474	4.802628
Closest to the Contact	Joint	16.68446	297.0478
Closest to the Contact	Joint	86.85181	6.430305
Closest to the Contact	Joint	85.49409	340.7939
Closest to the Contact	Joint	88.01561	4.266548
Closest to the Contact	Joint	56.07086	220.8775
Closest to the Contact	Joint	59.7705	269.5155
Closest to the Contact	Joint	74.87854	327.5915
Closest to the Contact	Joint	71.70718	168.8557

Closest to the Contact	Joint	57.92996	354.288
Closest to the Contact	Joint	82.84657	323.2185
Closest to the Contact	Joint	47.35606	242.7017
Closest to the Contact	Joint	62.01241	201.4467
Closest to the Contact	Joint	69.9558	338.3342
Closest to the Contact	Joint	70.73573	178.7818
Closest to the Contact	Joint	61.54294	352.1941
Closest to the Contact	Joint	78.20707	359.2153
Closest to the Contact	Joint	79.69637	8.993164
Closest to the Contact	Joint	74.34502	352.5594
Closest to the Contact	Joint	76.38618	340.2262
Closest to the Contact	Joint	74.93105	358.8836
Closest to the Contact	Joint	88.09032	340.1355
Closest to the Contact	Joint	14.02628	37.28598
Closest to the Contact	Joint	22.62389	358.7997
Closest to the Contact	Joint	3.236838	181.2697
Closest to the Contact	Joint	14.2286	263.506
Closest to the Contact	Joint	84.44711	25.51603
Closest to the Contact	Joint	18.40886	299.1305
Closest to the Contact	Joint	30.22222	330.6487
Closest to the Contact	Joint	79.06699	349.7332
Closest to the Contact	Joint	84.26547	345.7835
Closest to the Contact	Joint	23.5894	281.293
Closest to the Contact	Joint	58.55408	260.3865
Closest to the Contact	Joint	42.66602	318.7894
Closest to the Contact	Joint	29.00778	265.3836
Closest to the Contact	Joint	88.60589	350.4198
Closest to the Contact	Joint	57.46929	241.7545
Closest to the Contact	Joint	76.95856	170.0732
Closest to the Contact	Joint	82.5064	359.8226
Closest to the Contact	Joint	80.35164	354.9059
Closest to the Contact	Joint	13.93991	292.0927
Closest to the Contact	Joint	26.37685	306.2669
Closest to the Contact	Joint	49.56755	295.3436
Closest to the Contact	Joint	48.79955	233.2788
Closest to the Contact	Joint	37.15688	244.3333
Closest to the Contact	Joint	43.47303	235.7245
Closest to the Contact	Joint	46.64288	327.3669
Closest to the Contact	Joint	27.41118	225.4261
Closest to the Contact	Joint	85.65697	339.3006
Closest to the Contact	Joint	72.7849	350.3296
Closest to the Contact	Joint	56.51685	327.5405

Closest to the Contact	Joint	39.63931	308.5441
Closest to the Contact	Joint	81.20385	336.6258
Closest to the Contact	Joint	83.47166	337.4212
Closest to the Contact	Joint	86.67258	340.8282
Closest to the Contact	Joint	78.38163	341.7267
Closest to the Contact	Joint	73.83311	341.385
Closest to the Contact	Joint	79.49355	350.0431
Closest to the Contact	Bedding	66.3919	107.4412
Closest to the Contact	Bedding	71.84866	96.96304
Closest to the Contact	Bedding	74.49491	101.5121
Intermediate-Diapir Side	Joint	78.90175	81.24908
Intermediate-Diapir Side	Joint	85.51373	117.6112
Intermediate-Diapir Side	Joint	89	269.5485
Intermediate-Diapir Side	Joint	80.33014	256.8501
Intermediate-Diapir Side	Joint	45.66456	7.611664
Intermediate-Diapir Side	Joint	41.72137	352.8376
Intermediate-Diapir Side	Joint	43.91322	0.681587
Intermediate-Diapir Side	Joint	81.49525	270.1371
Intermediate-Diapir Side	Joint	84.65822	268.7297
Intermediate-Diapir Side	Joint	74.32809	258.1646
Intermediate-Diapir Side	Joint	86.19447	342.7787
Intermediate-Diapir Side	Joint	85.07203	257.6231
Intermediate-Diapir Side	Joint	85.86127	177.8627
Intermediate-Diapir Side	Joint	74.66614	331.6125
Intermediate-Diapir Side	Joint	77.8469	264.6998
Intermediate-Diapir Side	Joint	89	189.0116
Intermediate-Diapir Side	Joint	83.94551	346.4731
Intermediate-Diapir Side	Joint	74.3252	333.5776
Intermediate-Diapir Side	Joint	79.20492	352.3483
Intermediate-Diapir Side	Joint	62.99045	355.2895
Intermediate-Diapir Side	Joint	83.62292	267.0508
Intermediate-Diapir Side	Joint	66.56654	327.1881
Intermediate-Diapir Side	Joint	78.9804	318.4383
Intermediate-Diapir Side	Joint	86.1389	307.5062
Intermediate-Diapir Side	Joint	86.65659	95.81642
Intermediate-Diapir Side	Joint	81.60965	352.0054
Intermediate-Diapir Side	Joint	67.89848	254.5378
Intermediate-Diapir Side	Joint	84.99313	327.3287
Intermediate-Diapir Side	Joint	83.38138	331.4136
Intermediate-Diapir Side	Joint	75.85271	358.3398
Intermediate-Diapir Side	Joint	78.29082	96.75488
Intermediate-Diapir Side	Joint	86.69809	81.51573

Intermediate-Diapor Side	Joint	86.4196	255.6288
Intermediate-Diapor Side	Joint	80.37498	337.3206
Intermediate-Diapor Side	Joint	83.90166	78.44525
Intermediate-Diapor Side	Joint	68.97829	167.3194
Intermediate-Diapor Side	Joint	66.25835	10.49249
Intermediate-Diapor Side	Joint	75.66156	333.4166
Intermediate-Diapor Side	Joint	62.8471	338.801
Intermediate-Diapor Side	Joint	82.07809	255.4357
Intermediate-Diapor Side	Joint	88.11795	254.1507
Intermediate-Diapor Side	Joint	63.5294	354.4421
Intermediate-Diapor Side	Joint	73.93779	88.57458
Intermediate-Diapor Side	Joint	65.76099	89.49855
Intermediate-Diapor Side	Joint	51.79741	103.3762
Intermediate-Diapor Side	Joint	50.46798	98.30371
Intermediate-Diapor Side	Joint	50.72488	104.9934
Intermediate-Diapor Side	Joint	47.23248	30.76091
Intermediate-Diapor Side	Joint	49.52287	11.26071
Intermediate-Diapor Side	Joint	54.34226	17.67897
Intermediate-Diapor Side	Joint	79.86938	334.3727
Intermediate-Diapor Side	Joint	72.81657	355.897
Intermediate-Diapor Side	Joint	87.74935	84.00899
Intermediate-Diapor Side	Joint	75.64494	257.2453
Intermediate-Diapor Side	Joint	86.48572	98.57056
Intermediate-Diapor Side	Joint	74.73204	258.9835
Intermediate-Diapor Side	Joint	82.06952	256.9012
Intermediate-Diapor Side	Joint	79.46844	243.0808
Intermediate-Diapor Side	Joint	73.38313	245.3608
Intermediate-Diapor Side	Joint	70.40596	321.8748
Intermediate-Diapor Side	Joint	87.66964	270.6133
Intermediate-Diapor Side	Joint	81.72563	262.4959
Intermediate-Diapor Side	Joint	85.17966	89.25388
Intermediate-Diapor Side	Joint	76.22958	82.27763
Intermediate-Diapor Side	Joint	78.30201	252.0404
Intermediate-Diapor Side	Joint	54.52566	277.8014
Intermediate-Diapor Side	Joint	62.6143	12.19942
Intermediate-Diapor Side	Joint	81.76582	262.8541
Intermediate-Diapor Side	Joint	72.17008	1.240196
Intermediate-Diapor Side	Joint	71.06689	16.43361
Intermediate-Diapor Side	Joint	86.52545	95.89961
Intermediate-Diapor Side	Joint	76.69234	164.5347
Intermediate-Diapor Side	Joint	82.11655	338.2299
Intermediate-Diapor Side	Joint	74.531	15.64977

Intermediate-Diapir Side	Joint	77.21186	267.7764
Intermediate-Diapir Side	Joint	83.24171	261.2152
Intermediate-Diapir Side	Joint	77.57741	327.089
Intermediate-Diapir Side	Joint	88.52773	199.0355
Intermediate-Diapir Side	Joint	79.75354	256.1294
Intermediate-Diapir Side	Joint	83.46064	271.6438
Intermediate-Diapir Side	Bedding	78.6042	267.5342
Intermediate-Diapir Side	Bedding	75.96758	76.21291
Intermediate-Cutler Side	Joint	40.27813	86.82098
Intermediate-Cutler Side	Joint	43.90811	304.5864
Intermediate-Cutler Side	Joint	47.67201	288.32
Intermediate-Cutler Side	Joint	49.16245	285.6515
Intermediate-Cutler Side	Joint	49.25606	282.5595
Intermediate-Cutler Side	Joint	29.13193	256.674
Intermediate-Cutler Side	Joint	22.0894	260.1385
Intermediate-Cutler Side	Joint	33.38575	259.8147
Intermediate-Cutler Side	Joint	59.10919	341.6437
Intermediate-Cutler Side	Joint	40.14435	217.3633
Intermediate-Cutler Side	Joint	23.85224	288.3937
Intermediate-Cutler Side	Joint	58.63402	342.1318
Intermediate-Cutler Side	Joint	46.45857	262.2452
Intermediate-Cutler Side	Joint	41.11604	253.384
Intermediate-Cutler Side	Joint	39.01506	268.0264
Intermediate-Cutler Side	Joint	48.05645	251.8519
Intermediate-Cutler Side	Joint	44.39362	270.6625
Intermediate-Cutler Side	Joint	35.17915	283.8887
Intermediate-Cutler Side	Joint	35.82637	196.759
Intermediate-Cutler Side	Joint	63.57046	229.8501
Intermediate-Cutler Side	Joint	38.95605	355.5377
Intermediate-Cutler Side	Joint	36.98277	1.192863
Intermediate-Cutler Side	Joint	18.14006	289.2512
Intermediate-Cutler Side	Joint	24.8288	271.1459
Intermediate-Cutler Side	Joint	8.508178	139.1251
Intermediate-Cutler Side	Joint	60.53428	345.9532
Intermediate-Cutler Side	Joint	38.09842	276.2594
Intermediate-Cutler Side	Joint	38.4075	267.999
Intermediate-Cutler Side	Joint	36.61438	272.174
Intermediate-Cutler Side	Joint	34.45826	273.6054
Intermediate-Cutler Side	Bedding	57.29176	109.1294
Intermediate-Cutler Side	Bedding	57.49846	79.40802
Canyon-Second Farthest	Bedding	29.35419	147.2179
Canyon-Second Farthest	Joint	19.10887	196.5447

Canyon-Second Farthest	Joint	26.39067	235.7036
Canyon-Second Farthest	Joint	28.45381	210.9834
Canyon-Second Farthest	Joint	67.42216	228.0133
Canyon-Second Farthest	Joint	47.70707	219.292
Canyon-Second Farthest	Joint	44.69303	214.5069
Canyon-Second Farthest	Joint	44.78344	216.8981
Canyon-Second Farthest	Joint	45.24349	229.7761
Canyon-Second Farthest	Joint	52.16948	226.2637
Canyon-Second Farthest	Joint	53.849	228.7015
Canyon-Second Farthest	Joint	36.93391	204.6988
Canyon-Second Farthest	Joint	12.34341	236.3306
Canyon-Second Farthest	Joint	47.60762	241.6208
Canyon-Second Farthest	Joint	55.13077	234.8273
Canyon-Second Farthest	Joint	60.50264	222.4768
Canyon-Second Farthest	Joint	56.85497	229.4628
Canyon-Second Farthest	Joint	55.14408	200.3871
Canyon-Second Farthest	Joint	53.65765	231.1853
Canyon-Second Farthest	Joint	27.00332	278.7956
Canyon-Second Farthest	Joint	56.22416	253.4966
Canyon-Second Farthest	Joint	62.70671	256.2444
Canyon-Second Farthest	Joint	39.12247	321.106
Canyon-Second Farthest	Joint	30.06549	288.6241
Canyon-Second Farthest	Joint	10.92845	242.9874
Canyon-Second Farthest	Joint	87.46092	191.682
Canyon-Second Farthest	Joint	48.84378	223.7761
Canyon-Second Farthest	Joint	7.784497	72.7395
Canyon-Second Farthest	Joint	8.033218	294.463
Canyon-Second Farthest	Joint	85.02206	253.5732
Canyon-Second Farthest	Joint	79.49735	223.0456
Canyon-Second Farthest	Joint	56.96312	227.4567
Canyon-Second Farthest	Joint	56.03085	218.7663
Canyon-Second Farthest	Joint	40.56883	16.71011
Canyon-Second Farthest	Joint	41.13172	208.8376
Canyon-Second Farthest	Joint	34.45322	306.7082
Canyon-Second Farthest	Joint	59.48649	236.8348
Canyon-Second Farthest	Joint	56.03232	250.663
Canyon-Second Farthest	Joint	25.1353	285.008
Canyon-Second Farthest	Joint	88.29825	234.6545
Canyon-Second Farthest	Joint	74.28445	236.3994
Canyon-Second Farthest	Joint	25.38513	291.0316
Canyon-Second Farthest	Joint	41.8953	61.10728
Canyon-Second Farthest	Joint	36.25359	46.34106

Canyon-Second Farthest	Joint	4.96365	194.0903
Canyon-Second Farthest	Joint	31.66067	43.28693
Canyon-Second Farthest	Joint	42.13549	213.4944
Canyon-Second Farthest	Joint	57.60256	224.9394
Canyon-Second Farthest	Joint	56.71742	229.771
Canyon-Second Farthest	Bedding	31.87736	151.3948
Canyon-Farthest	Joint	21.57814	296.0702
Canyon-Farthest	Joint	23.41538	158.9892
Canyon-Farthest	Joint	17.36576	178.7625
Canyon-Farthest	Joint	33.23157	284.1706
Canyon-Farthest	Joint	25.02855	258.0387
Canyon-Farthest	Joint	33.66731	281.5063
Canyon-Farthest	Joint	21.58548	241.1544
Canyon-Farthest	Joint	79.49654	296.273
Canyon-Farthest	Joint	34.88399	268.159
Canyon-Farthest	Joint	21.45983	227.6596
Canyon-Farthest	Joint	26.36443	265.1959
Canyon-Farthest	Joint	35.2357	251.9789
Canyon-Farthest	Joint	32.67936	239.8276
Canyon-Farthest	Joint	45.41766	286.7774
Canyon-Farthest	Joint	40.42323	297.4745
Canyon-Farthest	Joint	31.07386	251.6371
Canyon-Farthest	Joint	30.02276	248.8743
Canyon-Farthest	Joint	39.30988	260.5481
Canyon-Farthest	Joint	44.81803	255.5728
Canyon-Farthest	Joint	85.30583	107.9164
Canyon-Farthest	Joint	86.99842	104.3452
Canyon-Farthest	Joint	43.06992	244.9853
Canyon-Farthest	Joint	49.55643	240.8705
Canyon-Farthest	Joint	40.91513	272.924
Canyon-Farthest	Joint	29.15745	233.087
Canyon-Farthest	Joint	28.23899	253.272
Canyon-Farthest	Joint	38.4818	272.9633
Canyon-Farthest	Joint	20.29613	221.8706
Canyon-Farthest	Joint	24.01116	244.3437
Canyon-Farthest	Joint	21.90931	248.2742
Canyon-Farthest	Joint	20.06557	262.3645
Canyon-Farthest	Joint	33.53745	289.715
Canyon-Farthest	Joint	23.52654	292.8325
Canyon-Farthest	Joint	12.51727	193.547
Canyon-Farthest	Joint	17.19106	176.8939
Canyon-Farthest	Joint	18.72604	219.4823

Canyon-Farthest	Joint	71.45715	299.5462
Canyon-Farthest	Joint	75.09755	292.1397
Canyon-Farthest	Joint	71.50311	302.7888
Canyon-Farthest	Joint	59.81118	303.2517
Canyon-Farthest	Joint	10.35397	238.9887
Canyon-Farthest	Joint	13.80909	235.1027
Canyon-Farthest	Joint	19.25791	235.7335
Canyon-Farthest	Joint	14.7776	258.813
Canyon-Farthest	Joint	14.29234	233.1546
Canyon-Farthest	Joint	28.8089	261.3554
Canyon-Farthest	Joint	27.7242	263.9151
Canyon-Farthest	Joint	21.96915	260.8155
Canyon-Farthest	Joint	58.2325	139.3162
Canyon-Farthest	Joint	35.98254	260.4971
Canyon-Farthest	Joint	32.74182	241.2232
Canyon-Farthest	Joint	33.30564	253.5287
Canyon-Farthest	Joint	35.68517	120.4176
Canyon-Farthest	Joint	39.19192	246.8119
Canyon-Farthest	Joint	38.63824	178.2515
Canyon-Farthest	Joint	85.75241	298.6631
Canyon-Farthest	Joint	42.98409	257.8492
Canyon-Farthest	Joint	49.59664	251.2373
Canyon-Farthest	Joint	29.20035	242.7045
Canyon-Farthest	Joint	41.96989	271.2112
Canyon-Farthest	Joint	22.27531	234.9883
Canyon-Farthest	Joint	19.4533	219.8324
Canyon-Farthest	Joint	34.05413	278.66
Canyon-Farthest	Joint	21.58517	229.8192
Canyon-Farthest	Joint	87.53558	121.5181
Canyon-Farthest	Joint	44.28306	260.3294
Canyon-Farthest	Joint	54.13665	280.1805
Canyon-Farthest	Joint	47.9461	278.8083
Canyon-Farthest	Joint	26.35475	261.2547
Canyon-Farthest	Joint	34.00388	272.3958
Canyon-Farthest	Joint	32.11048	253.3575
Canyon-Farthest	Joint	34.61721	252.6123
Canyon-Farthest	Joint	87.56047	117.8605
Canyon-Farthest	Bedding	33.35604	172.8015
Canyon-Farthest	Bedding	17.48084	183.6611
Canyon-Farthest	Bedding	28.81878	140.2048
Canyon-Farthest	Bedding	36.24105	161.766
Canyon-Farthest	Bedding	28.41006	201.5909

Appendix 2

Location Name	Plane Type	Dip	Strike
On the Contact	Joint	76.559	260.916
On the Contact	Joint	13.65	172.2
On the Contact	Joint	85.532	275.177
On the Contact	Joint	88.335	119.292
On the Contact	Joint	69.662	271.842
On the Contact	Joint	41.86	280.986
On the Contact	Joint	54.967	271.62
On the Contact	Joint	66.138	314.681
On the Contact	Joint	7.209	84.204
On the Contact	Joint	75.678	265.789
On the Contact	Joint	88.022	85.327
On the Contact	Joint	56.757	249.827
On the Contact	Joint	46.444	265.013
On the Contact	Joint	56.403	293.547
On the Contact	Joint	38.204	324.698
On the Contact	Joint	86.087	273.083
On the Contact	Joint	54.005	258.559
On the Contact	Joint	86.044	84.579
On the Contact	Joint	89.014	81.944
On the Contact	Joint	86.846	96.255
On the Contact	Joint	84.956	272.435
On the Contact	Joint	68.064	74.29
On the Contact	Joint	43.129	55.384
On the Contact	Joint	59.657	276.885
On the Contact	Joint	47.855	255.809
On the Contact	Joint	83.956	285.395
On the Contact	Joint	77.373	260.514
On the Contact	Joint	62.141	30.122
On the Contact	Joint	25.458	271.278
On the Contact	Joint	9.043	286.984
On the Contact	Joint	82.158	129.345
On the Contact	Joint	24.189	264.196
On the Contact	Joint	64.799	315.769
On the Contact	Joint	63.26	58.123
On the Contact	Joint	50.39	105.757
On the Contact	Joint	19.648	309.163
On the Contact	Joint	79.3	132.111
On the Contact	Joint	88.741	266.291
On the Contact	Joint	57.802	278.856

On the Contact	Joint	80.4	53.429
Closest to the Contact	Bedding	65.321	98.505
Closest to the Contact	Bedding	68.6	102.778
Closest to the Contact	Bedding	68.54	71.82
Closest to the Contact	Joint	45.574	253.105
Closest to the Contact	Joint	38.482	284.159
Closest to the Contact	Joint	62.328	279.435
Closest to the Contact	Joint	31.509	303.415
Closest to the Contact	Joint	55.993	261.309
Closest to the Contact	Joint	86.551	196.459
Closest to the Contact	Joint	89.589	18.55
Closest to the Contact	Joint	88.565	199.955
Closest to the Contact	Joint	47.542	268.261
Closest to the Contact	Joint	16.269	297.175
Closest to the Contact	Joint	78.098	281.576
Closest to the Contact	Joint	21.11	269.159
Closest to the Contact	Joint	69.031	261.329
Closest to the Contact	Joint	33.893	256.564
Closest to the Contact	Joint	87.587	197.411
Closest to the Contact	Joint	85.819	39.46
Closest to the Contact	Joint	81.432	201.064
Closest to the Contact	Joint	83.884	24.825
Closest to the Contact	Joint	87.083	198.89
Closest to the Contact	Joint	89.912	202.081
Closest to the Contact	Joint	57.297	244.988
Closest to the Contact	Joint	88.376	16.338
Closest to the Contact	Joint	89.798	16.888
Closest to the Contact	Joint	32.512	292.559
Closest to the Contact	Joint	36.358	262.133
Closest to the Contact	Joint	49.379	280.469
Closest to the Contact	Joint	89.97	209.44
Closest to the Contact	Joint	57.461	284.478
Closest to the Contact	Joint	45.827	304.562
Closest to the Contact	Joint	17.639	6.888
Closest to the Contact	Joint	43.363	238.944
Closest to the Contact	Joint	81.04	259.307
Closest to the Contact	Joint	62.705	101.464
Closest to the Contact	Joint	79.866	303.661
Closest to the Contact	Joint	41.557	287.692
Closest to the Contact	Joint	79.756	22.689
Closest to the Contact	Joint	86.139	200.227
Closest to the Contact	Joint	86.842	200.552

Closest to the Contact	Joint	53.795	197.945
Closest to the Contact	Joint	88.977	197.758
Closest to the Contact	Joint	77.515	33.678
Closest to the Contact	Joint	49.342	39.757
Closest to the Contact	Joint	87.775	47.343
Closest to the Contact	Joint	20.144	290.347
Closest to the Contact	Joint	83.374	58.37
Intermediate-Cutler Side	Bedding	50.176	110.301
Intermediate-Cutler Side	Bedding	51.791	110.051
Intermediate-Cutler Side	Bedding	29.046	137.03
Intermediate-Cutler Side	Bedding	55.622	107.94
Intermediate-Cutler Side	Bedding	59.001	103.256
Intermediate-Cutler Side	Joint	37.502	278.209
Intermediate-Cutler Side	Joint	43.81	282.309
Intermediate-Cutler Side	Joint	85.705	3.454
Intermediate-Cutler Side	Joint	40.221	182.416
Intermediate-Cutler Side	Joint	53.567	86.395
Intermediate-Cutler Side	Joint	61.914	17.52
Intermediate-Cutler Side	Joint	53.474	28.108
Intermediate-Cutler Side	Joint	65.603	159.141
Intermediate-Cutler Side	Joint	39.016	178.906
Intermediate-Cutler Side	Joint	65.747	25.624
Intermediate-Cutler Side	Joint	35.319	64.123
Intermediate-Cutler Side	Joint	51.394	305.517
Intermediate-Cutler Side	Joint	16.647	169.77
Intermediate-Cutler Side	Joint	54.548	165.389
Intermediate-Cutler Side	Joint	57.239	207.093
Intermediate-Cutler Side	Joint	43.268	292.794
Intermediate-Cutler Side	Joint	84.013	136.649
Intermediate-Cutler Side	Joint	47.059	302.384
Intermediate-Cutler Side	Joint	54.797	275.831
Intermediate-Cutler Side	Joint	41.701	318.594
Intermediate-Cutler Side	Joint	28.582	288.207
Intermediate-Cutler Side	Joint	33.393	226.471
Intermediate-Cutler Side	Joint	11.102	340.801
Intermediate-Cutler Side	Joint	44.617	305.711
Intermediate-Cutler Side	Joint	35.095	48.562
Intermediate-Cutler Side	Joint	61.666	312.15
Intermediate-Cutler Side	Joint	47.495	270.892
Intermediate-Cutler Side	Joint	81.932	133.884
Intermediate-Cutler Side	Joint	39.828	238.588
Intermediate-Cutler Side	Joint	37.067	320.874

Intermediate-Cutler Side	Joint	42.07	320.008
Intermediate-Cutler Side	Joint	49.285	204.436
Intermediate-Cutler Side	Joint	33.384	293.952
Intermediate-Cutler Side	Joint	23.507	348.459
Intermediate-Cutler Side	Joint	66.677	329.809
Intermediate-Cutler Side	Joint	42.45	342.968
Intermediate-Cutler Side	Joint	29.637	335.794
Intermediate-Cutler Side	Joint	59.87	340.416
Intermediate-Cutler Side	Joint	56.88	30.517
Intermediate-Cutler Side	Joint	57.063	338.326
Intermediate-Cutler Side	Joint	49.788	276.766
Intermediate-Cutler Side	Joint	68.11	295.683
Intermediate-Cutler Side	Joint	43.097	30.497
Intermediate-Cutler Side	Joint	78.504	3.543
Intermediate-Cutler Side	Joint	8.498	142.058
Intermediate-Cutler Side	Joint	16.977	134.469
Intermediate-Cutler Side	Joint	36.516	130.287
Intermediate-Cutler Side	Joint	33.758	308.042
Intermediate-Cutler Side	Joint	28.79	65.206
Intermediate-Cutler Side	Joint	61.592	204.23
Intermediate-Cutler Side	Joint	81.564	293.083
Intermediate-Cutler Side	Joint	14.085	278.306
Intermediate-Cutler Side	Joint	46.143	339.075
Intermediate-Cutler Side	Joint	69.458	296.883
Intermediate-Cutler Side	Joint	71.521	336.022
Intermediate-Cutler Side	Joint	22.544	208.142
Intermediate-Cutler Side	Joint	23.042	190.745
Intermediate-Cutler Side	Joint	21.904	4.879
Intermediate-Cutler Side	Joint	47.977	342.95
Intermediate-Cutler Side	Joint	33.655	327.79
Intermediate-Cutler Side	Joint	44.934	344.578
Intermediate-Cutler Side	Joint	77.286	169.583
Intermediate-Cutler Side	Joint	36.54	313.598
Second Farthest	Bedding	33.53	145.032
Second Farthest	Bedding	27.306	123.614
Second Farthest	Bedding	28.577	139.514
Second Farthest	Bedding	45.576	140.639
Second Farthest	Joint	45.645	229.577
Second Farthest	Joint	37.178	226.595
Second Farthest	Joint	58.648	264.622
Second Farthest	Joint	50.464	245.519
Second Farthest	Joint	53.347	261.91

Second Farthest	Joint	31.389	225.352
Second Farthest	Joint	29.917	208.681
Second Farthest	Joint	78.245	90.268
Second Farthest	Joint	52.604	107.194
Second Farthest	Joint	76.809	250.25
Second Farthest	Joint	73.851	81.515
Second Farthest	Joint	42.314	226.37
Second Farthest	Joint	36.951	225.938
Second Farthest	Joint	42.799	236.081
Second Farthest	Joint	40.949	219.494
Second Farthest	Joint	50.021	233.158
Second Farthest	Joint	50.282	241.519
Second Farthest	Joint	40.834	316.387
Second Farthest	Joint	60.919	336.632
Second Farthest	Joint	47.267	335.522
Second Farthest	Joint	44.348	321.807
Second Farthest	Joint	38.14	319.246
Second Farthest	Joint	57.92	310.364
Second Farthest	Joint	43.273	325.44
Second Farthest	Joint	45.383	295.28
Second Farthest	Joint	52.292	324.864
Second Farthest	Joint	36.034	309.582
Second Farthest	Joint	47.331	346.972
Second Farthest	Joint	61.138	331.527
Second Farthest	Joint	30.145	318.176
Second Farthest	Joint	44.099	316.512
Second Farthest	Joint	61.191	335.442
Second Farthest	Joint	62.768	319.14
Second Farthest	Joint	46.31	331.45
Second Farthest	Joint	42.32	327.981
Second Farthest	Joint	64.939	330.592
Second Farthest	Joint	27.668	331.205
Second Farthest	Joint	17.763	0.235
Second Farthest	Joint	47.714	313.238
Second Farthest	Joint	41.141	306.236
Second Farthest	Joint	58.821	242.735
Second Farthest	Joint	50.321	239.324
Second Farthest	Joint	48.602	338.051
Second Farthest	Joint	45.794	235.358
Second Farthest	Joint	43.513	235.41
Second Farthest	Joint	41.018	241.67
Second Farthest	Joint	33.568	322.214

Second Farthest	Joint	30.709	50.94
Second Farthest	Joint	57.144	227.636
Second Farthest	Joint	64.393	253.831
Second Farthest	Joint	69.242	186.267
Second Farthest	Joint	64.739	244.451
Second Farthest	Joint	62.418	267.506
Second Farthest	Joint	64.691	106.229
Second Farthest	Joint	63.247	325.603
Second Farthest	Joint	38.473	199.82
Second Farthest	Joint	40.834	317.336
Second Farthest	Joint	51.478	264.815
Second Farthest	Joint	42.848	251.746
Second Farthest	Joint	61.877	255.136
Second Farthest	Joint	25.391	231.93
Second Farthest	Joint	57.929	229.951
Second Farthest	Joint	64.476	139.22
Second Farthest	Joint	54.152	126.443
Second Farthest	Joint	16.926	228.258
Second Farthest	Joint	41.729	190.309
Second Farthest	Joint	69.496	201.094
Second Farthest	Joint	41.361	162.491
Second Farthest	Joint	75.915	319.086
Second Farthest	Joint	55.557	259.911
Second Farthest	Joint	57.915	171.127
Second Farthest	Joint	54.44	246.785
Second Farthest	Joint	45.348	170.601
Second Farthest	Joint	49.619	312.1

Appendix 3

Tool	Dip	Strike
Phone	4.948235	188.8965
Phone	71.62321	186.5846
Phone	30.13729	68.16112
Phone	18.54878	63.49782
Phone	22.27633	11.7323
Phone	36.3609	24.42918
Phone	27.80125	155.783
Phone	33.36781	61.55971
Phone	34.31878	49.22037
Phone	26.90308	42.90192
Phone	30.36902	47.8577
Phone	20.25322	9.875694
Phone	25.91556	339.083
Phone	27.16047	332.9358
Phone	14.5024	2.386642
Phone	84.75029	7.09938
Phone	63.73704	189.4639
Phone	18.79341	117.8147
Phone	22.02708	163.2516
Phone	33.64728	163.5974
Phone	34.92154	179.8282
Phone	25.67582	83.19408
Phone	31.58885	125.4419
Phone	39.01783	144.8111
Phone	62.42533	167.5327
Phone	16.45887	31.51407
Phone	29.07506	170.4008
Phone	21.84569	61.87312
Phone	19.38584	77.45442
Phone	71.22604	338.7667
Phone	19.5227	84.13745
Phone	33.11543	84.29301
Phone	32.43503	86.71559
Phone	20.73262	8.234489
Phone	73.44069	4.649513
Phone	16.302	87.56659
Phone	8.412468	165.382
Phone	25.88165	74.62738
Phone	7.87179	175.1519

Phone	68.20924	341.3586
Phone	60.76585	339.5681
Phone	20.66308	60.50365
Phone	13.16547	68.11871
Phone	35.78185	14.02145
Phone	69.44563	338.7291
Phone	69.93477	160.384
Phone	17.65869	346.136
Phone	82.54185	160.0206
Phone	82.79329	348.9674
Phone	44.05102	344.017
Phone	62.4877	5.330338
Phone	21.0619	95.44962
Phone	25.46165	70.90918
Phone	27.45781	94.2907
Phone	22.10174	84.05116
Phone	36.19484	357.8251
Phone	33.01919	3.743591
Phone	35.15741	351.5049
Phone	57.80047	353.3727
Phone	88.40904	11.64263
Compass	90	353
Compass	59	356
Compass	39	342
Compass	35	350
Compass	16	90
Compass	50	2
Compass	20	105
Compass	25	81
Compass	29	94
Compass	53	349
Compass	65	5
Compass	89	346
Compass	76	347
Compass	15	78
Compass	69	74
Compass	18	52
Compass	55	341
Compass	19	106
Compass	26	70
Compass	13	90
Compass	7	14

Compass	71	348
Compass	3	47
Compass	26	85
Compass	27	152
Compass	81	345
Compass	31	351
Compass	24	258
Compass	72	348
Compass	15	85
Compass	16	72
Compass	83	142
Compass	25	72
Compass	20	85
Compass	14	48
Compass	58	352
Compass	29	339
Compass	24	30
Compass	4	2
Compass	25	55
Compass	29	347
Compass	41	178
Compass	67	174
Compass	17	224
Compass	43	194
Compass	18	206
Compass	25	262
Compass	29	340
Compass	26	5
Compass	50	40
Compass	38	38
Compass	21	50
Compass	24	318
Compass	30	26
Compass	36	22
Compass	24	247
Compass	25	1
Compass	70	5
Compass	9	329

Appendix 4

Dip	Strike
86.93183899	188.0206
78.43344879	8.237022
9.94253826	121.3026
74.69129944	153.1973
46.56910324	7.359779
74.72724915	159.6113
81.03470612	166.0405
22.23875618	354.0491
17.58584785	58.61084
79.88212585	175.3539
74.75798035	164.0471
52.78524017	158.4556
76.27741241	351.847
12.49042988	118.5608
84.75091553	355.7199
65.16580963	158.2955
72.48669434	353.5581
74.23563385	172.4541
80.51147461	169.8446
79.60308075	168.7347
84.81478119	157.8499
80.82881165	345.9218
81.91376495	173.9259
32.30918121	162.1965
14.27886486	59.86685
10.48571396	80.75381
7.32533503	61.34114
66.4355545	156.9999
73.95700836	169.2721
78.32726288	178.8651
54.78507233	275.8413
43.6975708	270.4276
21.89840126	323.5857
26.42608261	331.0451
45.90945816	355.4197
15.33202553	121.2265
55.38374329	358.2066
62.48686981	358.8654
7.11376524	28.24515

70.89863586	164.3468
78.63576508	169.5912
8.95658684	319.066
29.20736122	359.0963
31.432621	5.780457
77.86917114	335.1301
13.58761501	25.10345
16.3331871	106.1944
80.95529175	154.6776
36.0331955	16.44467
51.52423477	352.2583
66.83234406	168.8799
73.84069824	179.1206
82.8106842	180.9504
69.58707428	173.6232
82.55632782	177.2373
25.61652756	292.4061
14.48073006	334.2122
77.61930847	170.5388
71.12614441	161.5626
78.17943573	161.7873
82.74724579	164.4059
69.49313354	153.517
68.43950653	154.971
73.63913727	157.9221
67.89931488	162.9387
86.92583466	351.6616
89	277
68	262
65	277
57	264
62	270
80	276
89	332
34	261
88	271
82	295
80	278
66	282
75	264
86	265
56	264

71	275
73	265
76	109
66	267
79	263
88	90
89	260
59	261
89	346
75	290
69	287
84	259
89	283
48	268
75	274
88	260
85	251
89	179
75	257
84	109
69	267
85	88
74	259
50	266
76	269
88	175
33	32
78	65
24	277

Vita

Muhammed Tarik Iraz, the oldest of four children by Faruk and Rakiye, was born in 1991 in Istanbul, Turkey. He was always curious of man-made items and this made him interested in engineering. Because of his father's job, he went to school in many different cities, and he finally completed his high school, Midyat Aziz Onen Lisesi, Mardin, in 2009. Later, he began studying Geological Engineering in Akdeniz University, Antalya and graduated in 2013. Once he was a senior, he wrote his seminar study and completing thesis on Remote Sensing.

To start working in national companies, he prepared for the national exam that is written once in two years and placed 94th from over 12000 people in his major, in 2014. While he was waiting for the results, he started his graduate studies in Akdeniz University and applied for a study abroad scholarship sponsored by the Ministry of National Education. After he placed 3rd from 57 people, he started studying English in Akdeniz University, Yildiz Technical University and University of Texas at Arlington from March 2015 to December 2016.

In the spring of 2017, he started his master's degree in University of Texas at El Paso. Based on his agreement with Turkish Petroleum Corporation, he focused on fractures and completed his thesis on fractures of Western Onion Creek, Moab, UT.

Contact information: mtiraz@miners.utep.edu

+90 531 855 47 47

This thesis/dissertation was typed by Muhammed Tarik Iraz.